

Triplet de changements de phases

György HETÉNYI

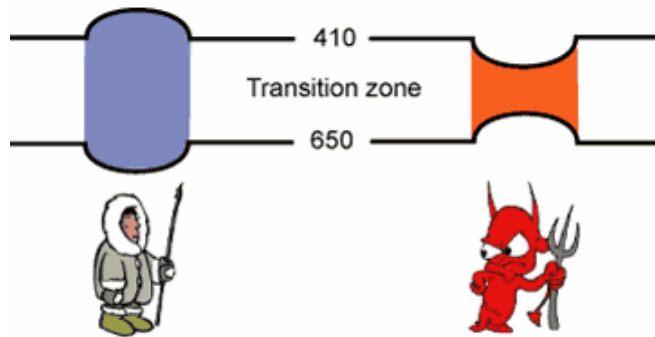
ETH Zürich

`gyorgy.hetenyi@erdw.ethz.ch`

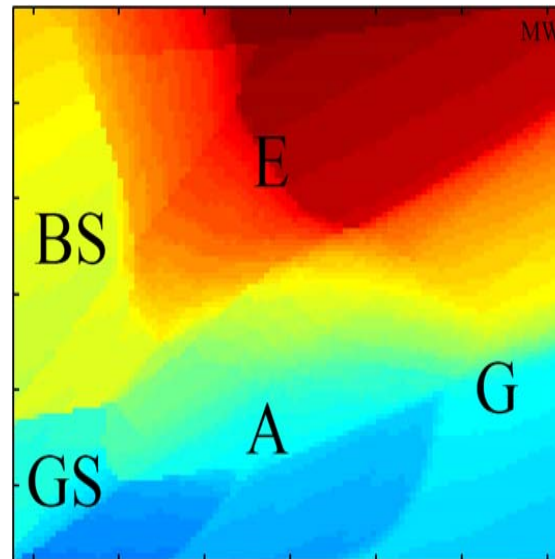


Triplet de changements de phases

zone de transition du manteau



modélisation géodynamique

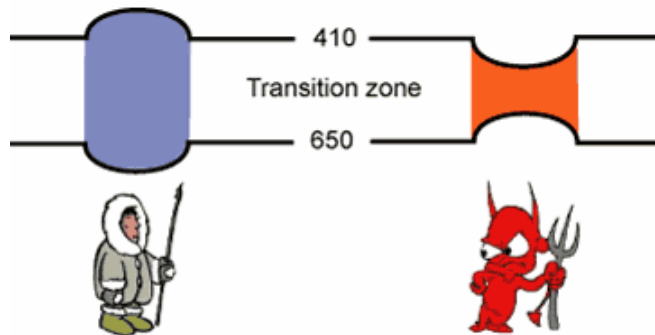


prismation des orgues de lave

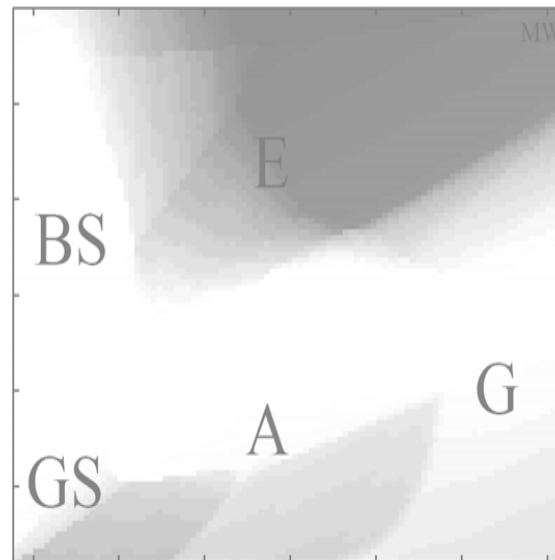


Triplet de changements de phases

zone de transition du manteau



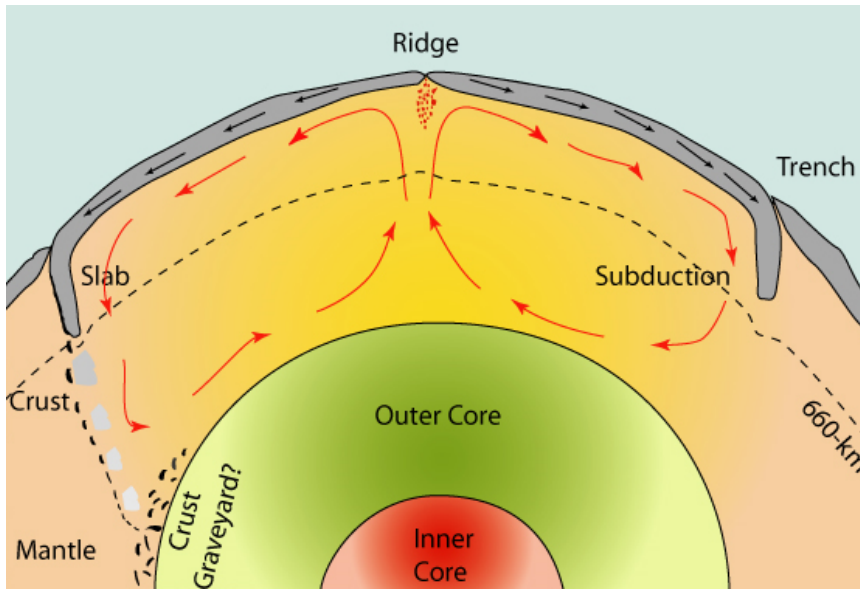
modélisation géodynamique



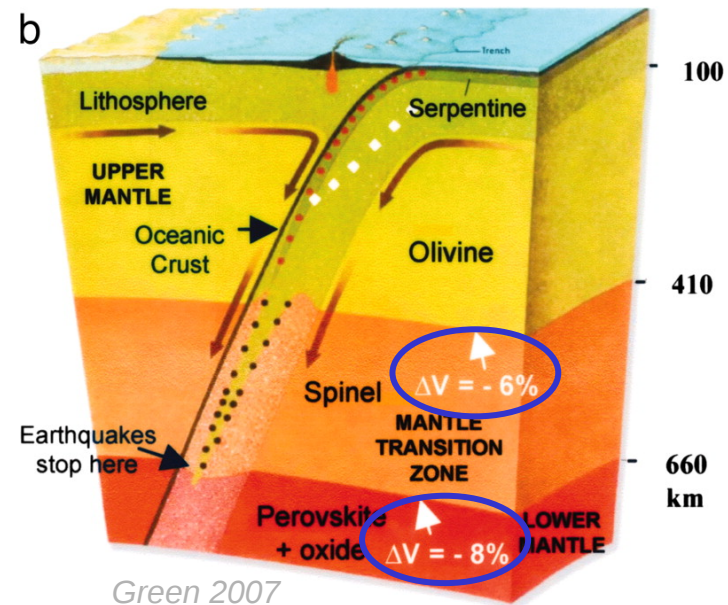
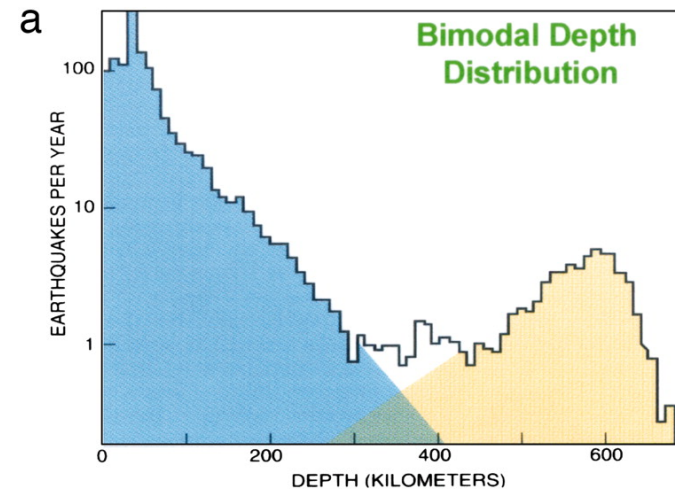
prismation des orgues de lave



Mantle structure and transition zone



S. Rost

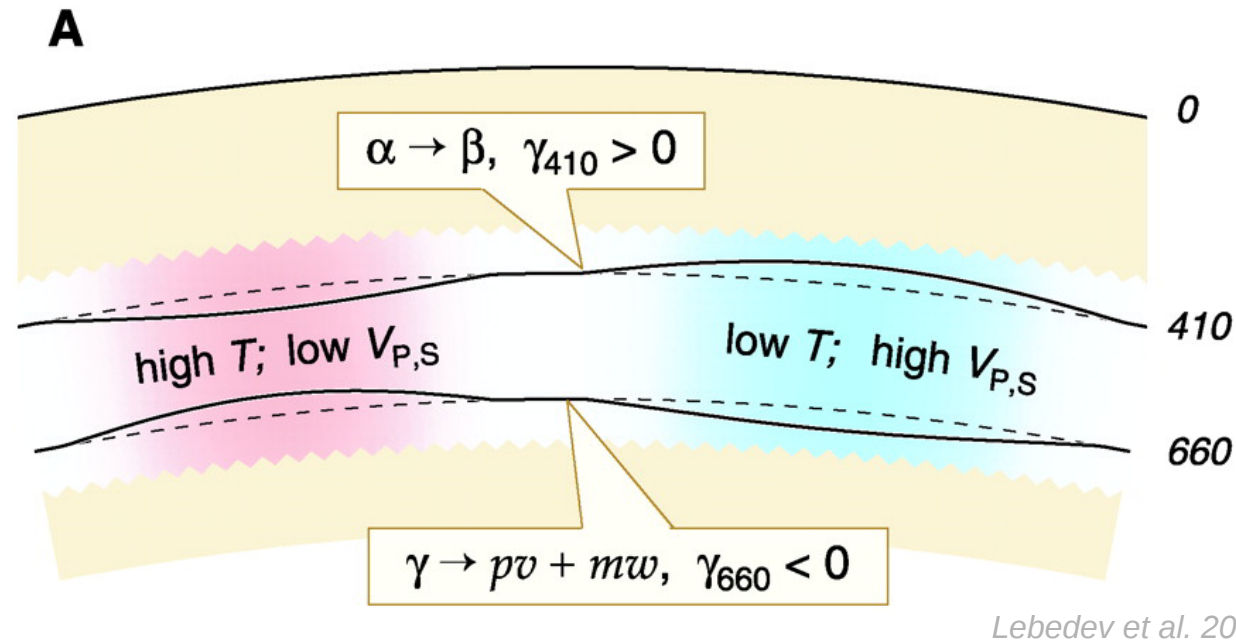


Green 2007

<i>iasp91</i>	Δv_p	Δv_s
410 km:	+3.65%	+4.11%
660 km:	+5.78%	+6.25%

Kennett & Engdahl 1991

Transition zone in an olivine-dominant mantle



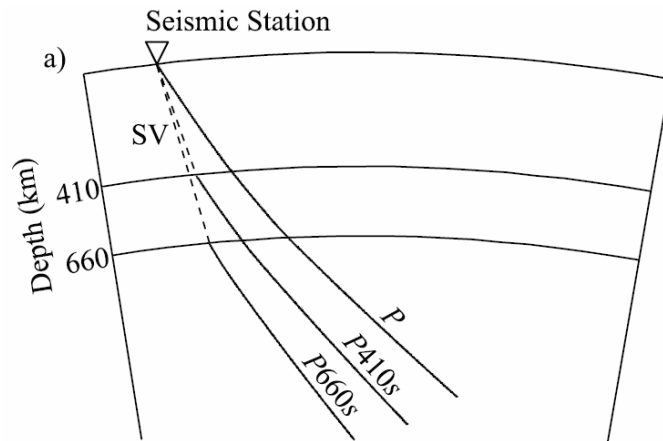
410 km: olivine (α) $\rightarrow$ wadsleyite (β)

520 km: wadsleyite (β) $\rightarrow$ ringwoodite (γ)

660 km: ringwoodite (γ) $\rightarrow$ perovskite + magnesiowüstite

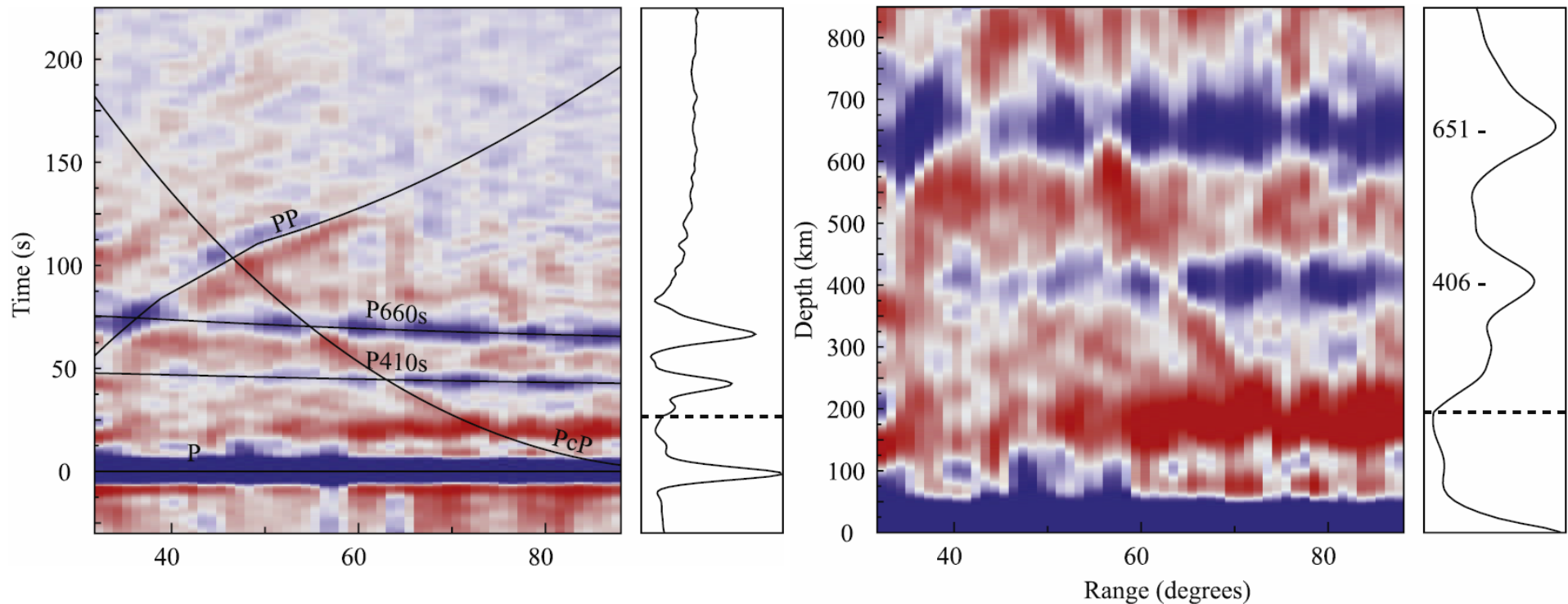
Temperature effect: opposite Clapeyron-slope ($\partial P/\partial T$) signs and deflections

Observations by seismology

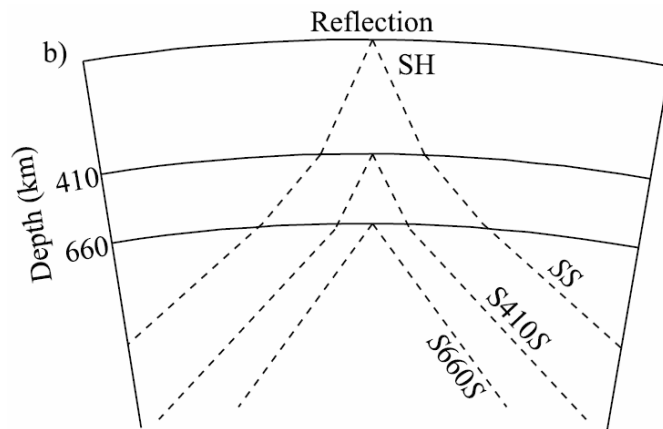


Receiver functions

- P-to-S wave conversion
- delay proportional to depth
(and velocity anomalies)

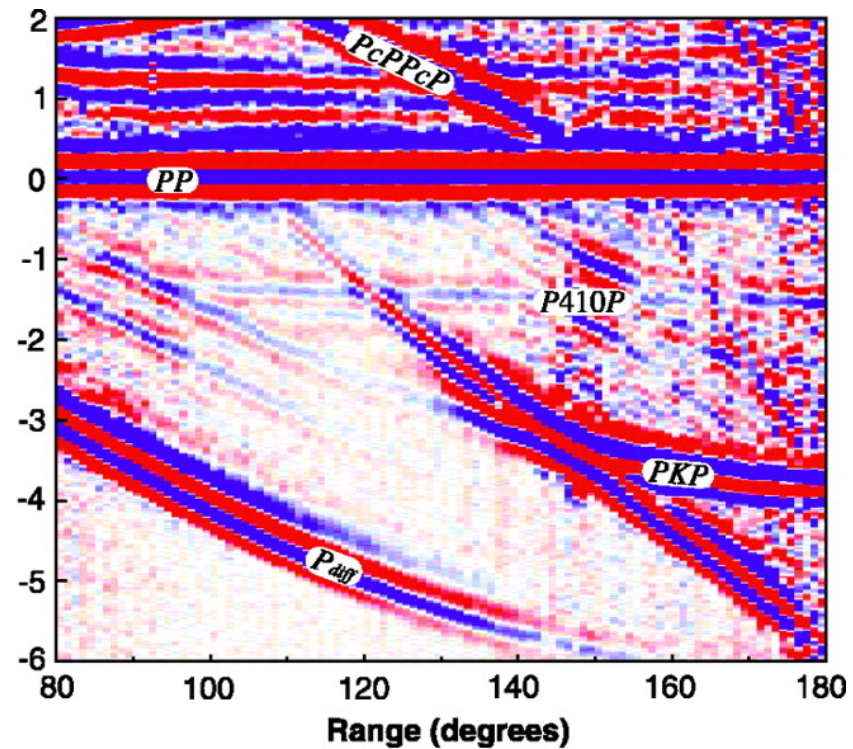
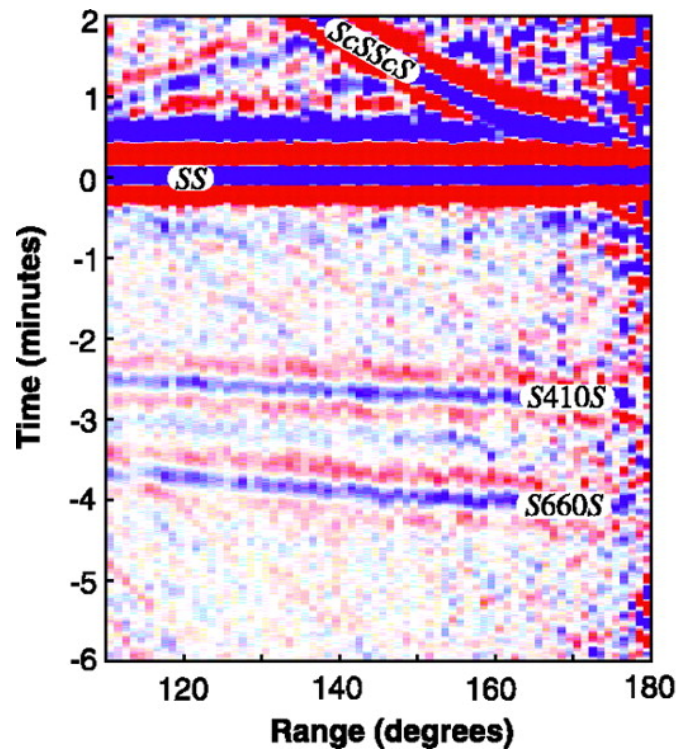


Observations by seismology

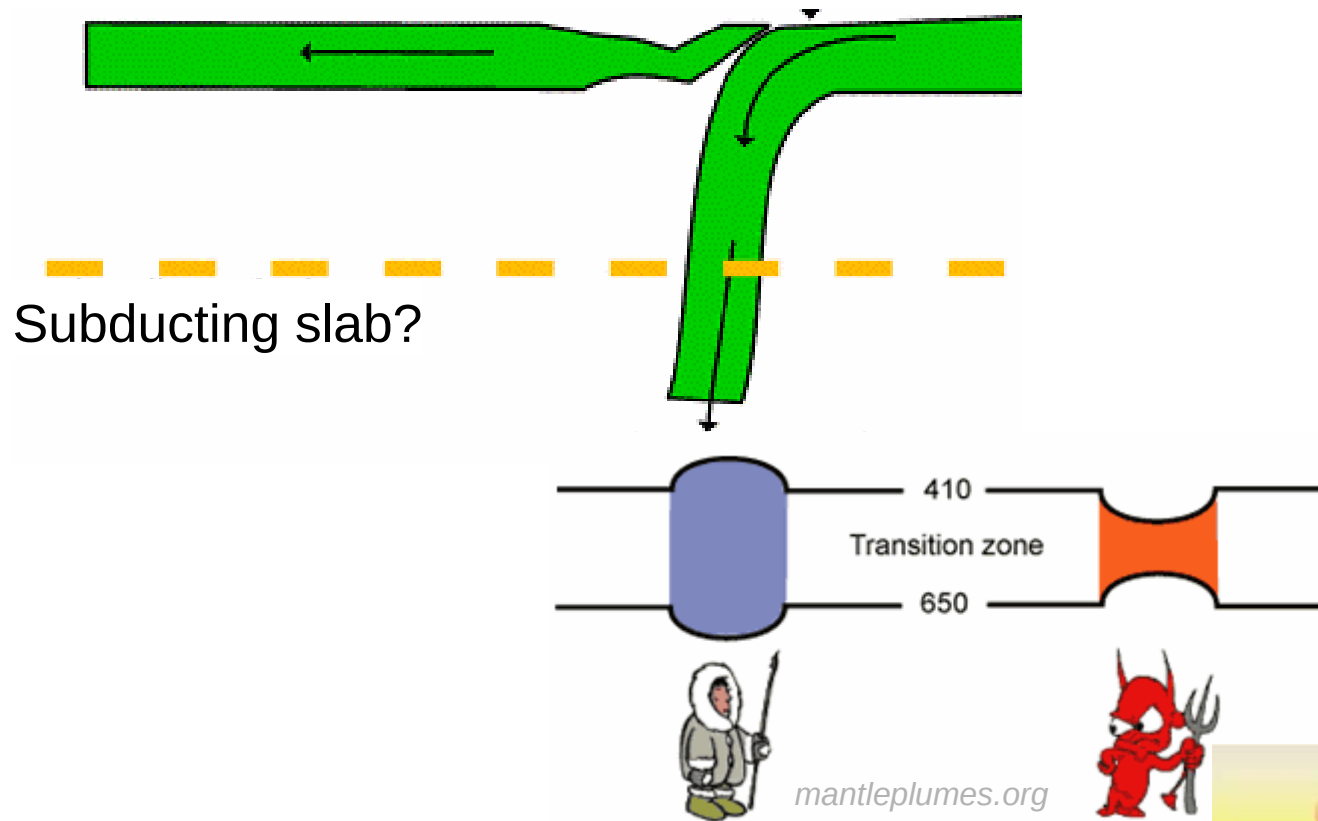


Underside reflections / pre-cursors

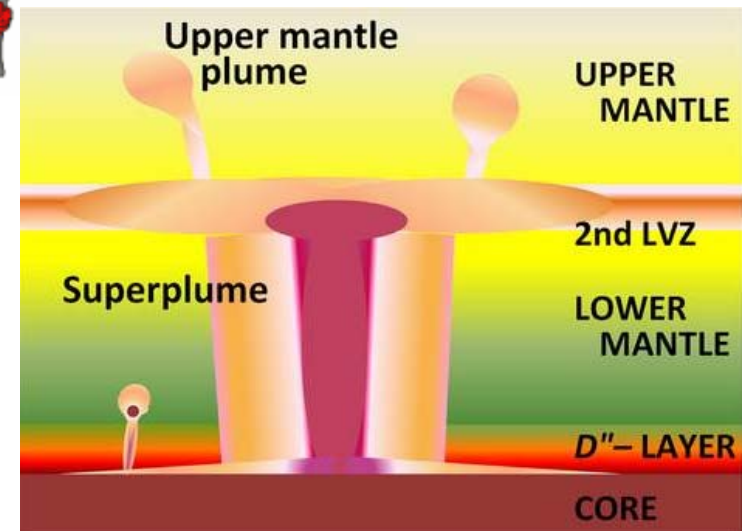
- with S- or P-waves
- long-wavelength variations



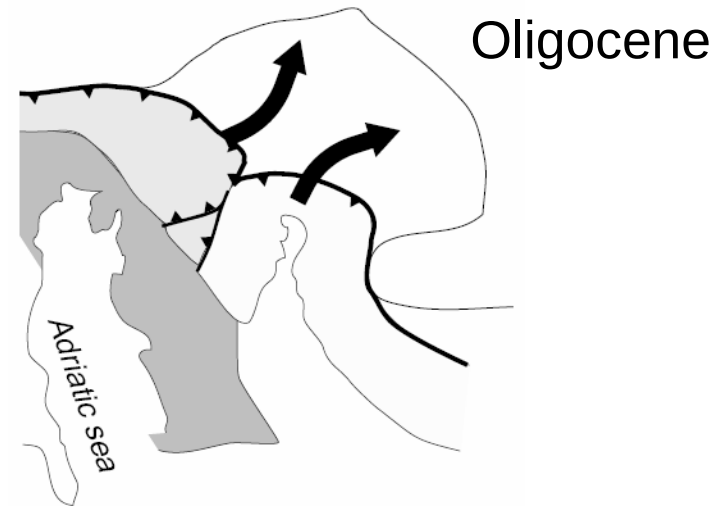
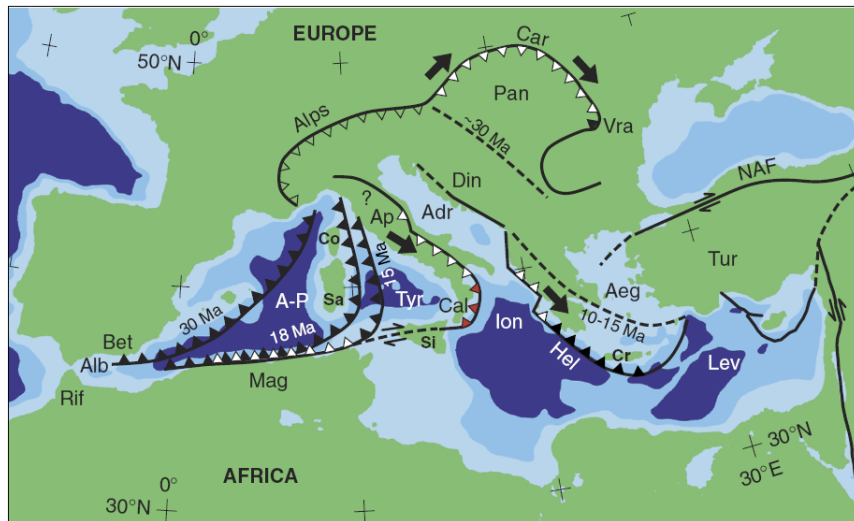
Geodynamic interpretation



Plume?

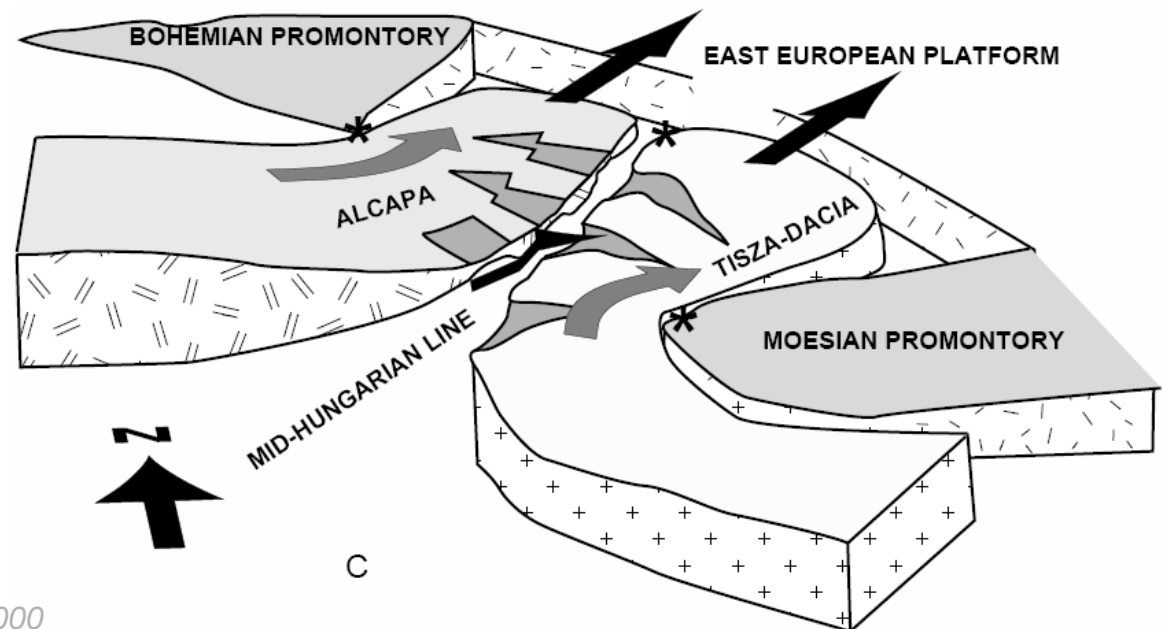


Pannonian Basin



Today:

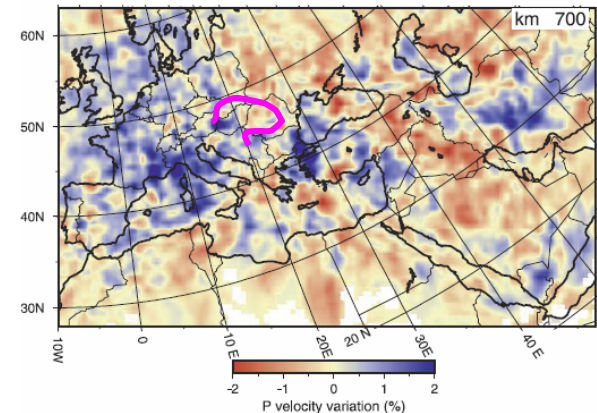
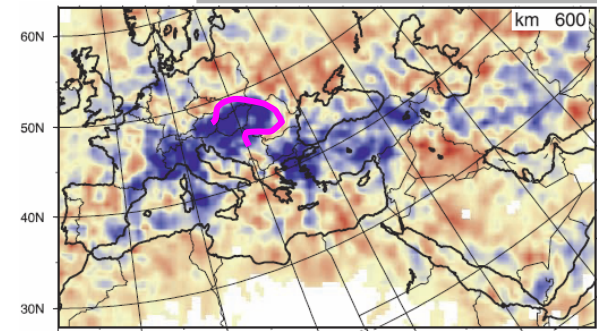
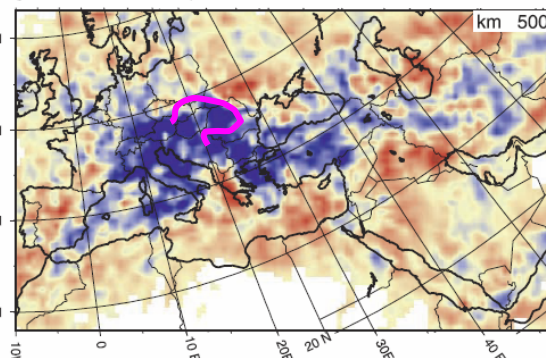
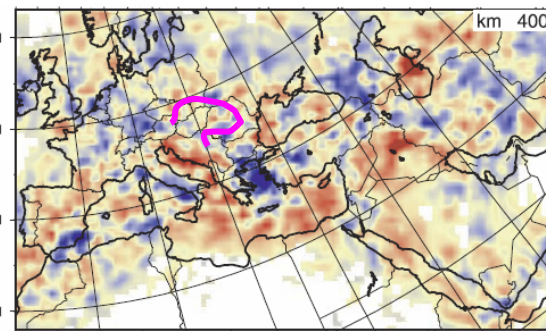
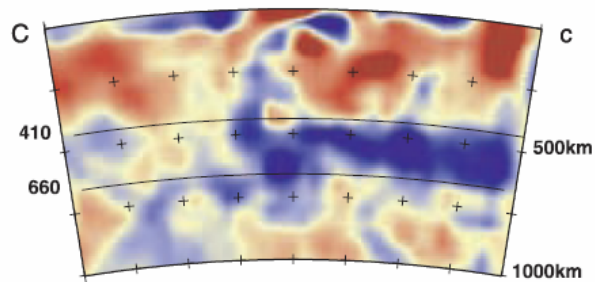
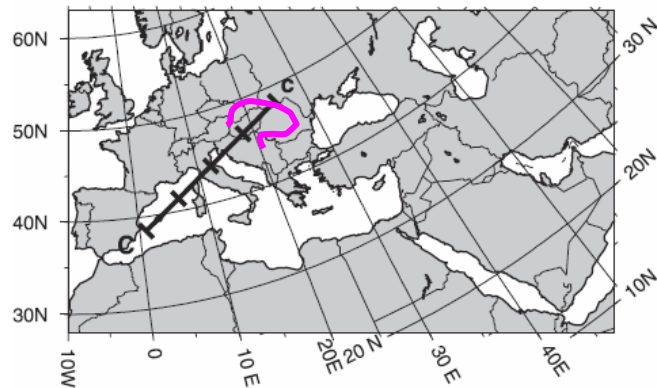
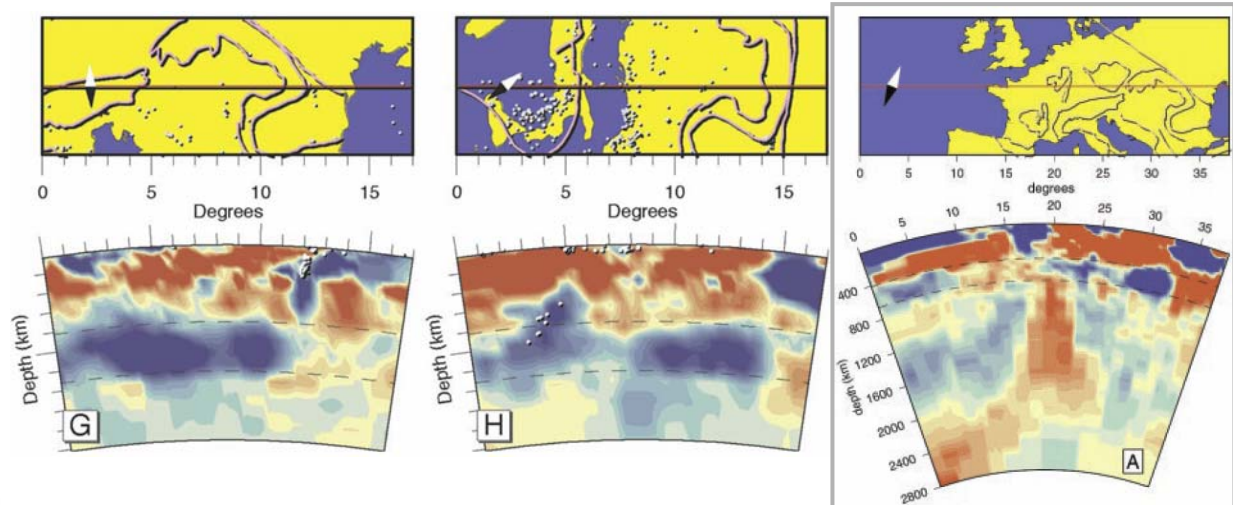
- thin crust (25km) and lithosphere (55km)
- high heat flow (90mW/m^2)
- minor surface deformation (GPS $\sim\text{mm/yr}$)
- minor seismic activity (M4-5-...)



Pannonian Basin

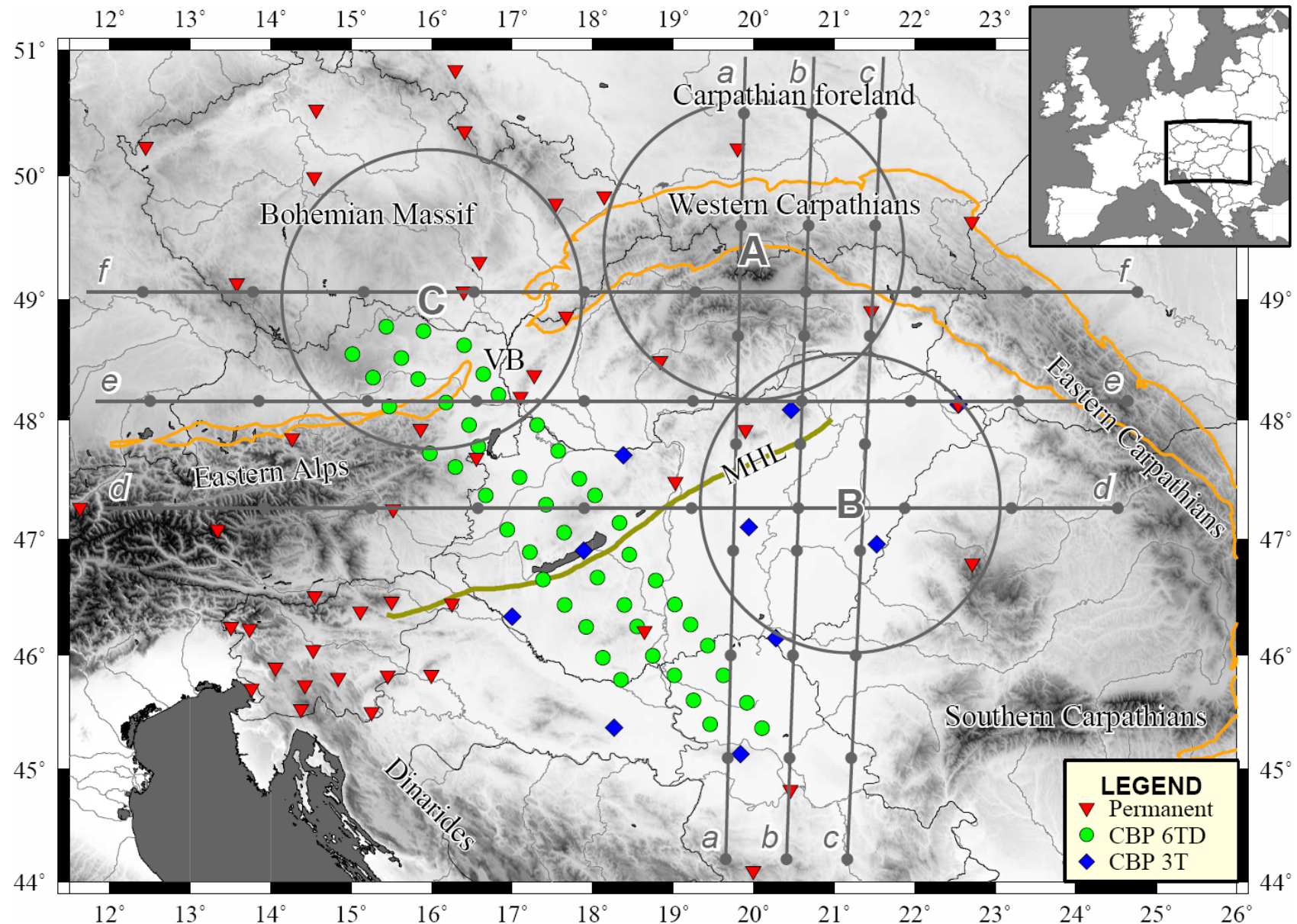
...today:

- fast anomaly in mantle transition zone
- European 'superplume'



Goes et al. 1999; Wortel & Spakman 2000; Piromallo & Morelli 2003

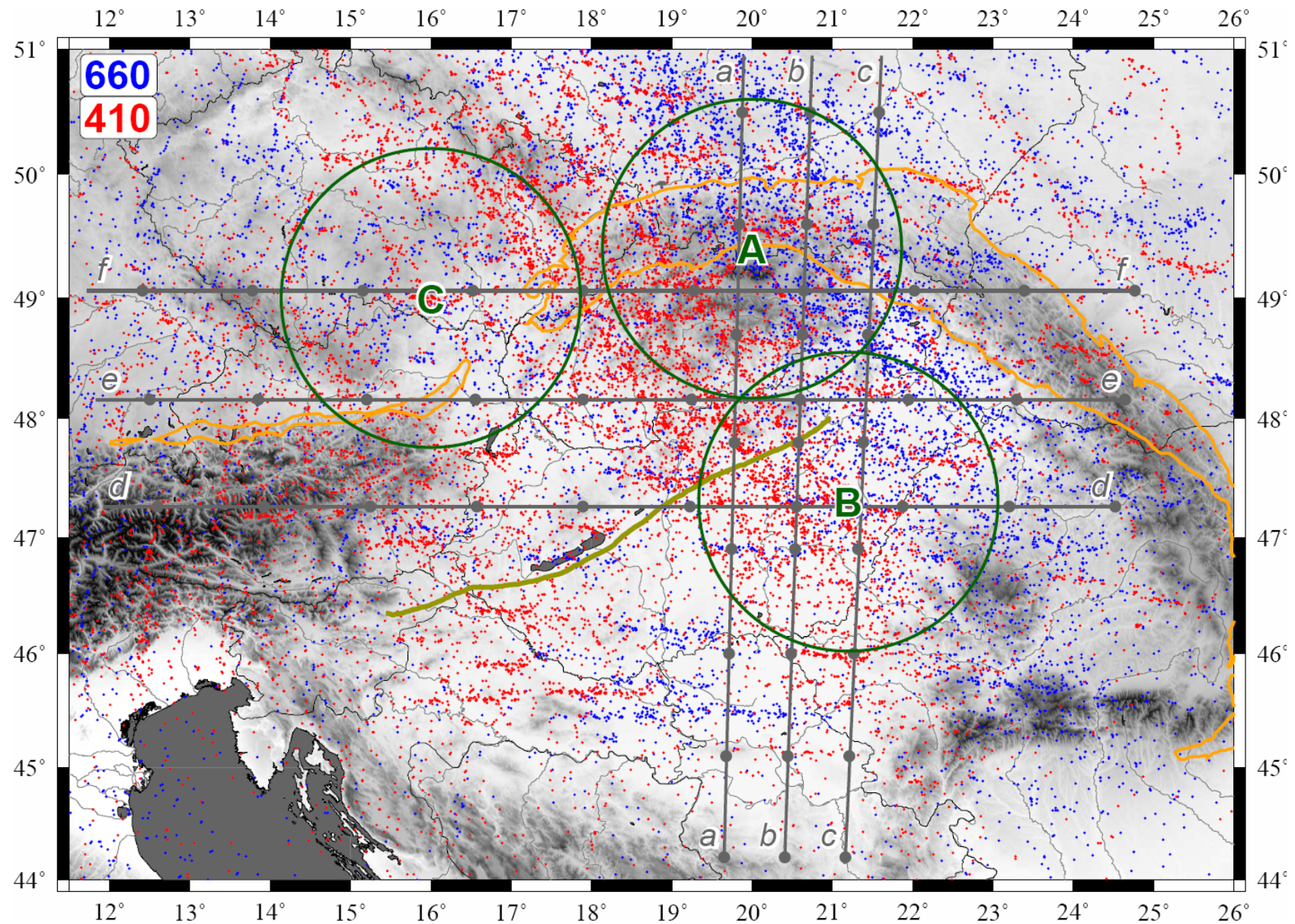
Pannonian Basin



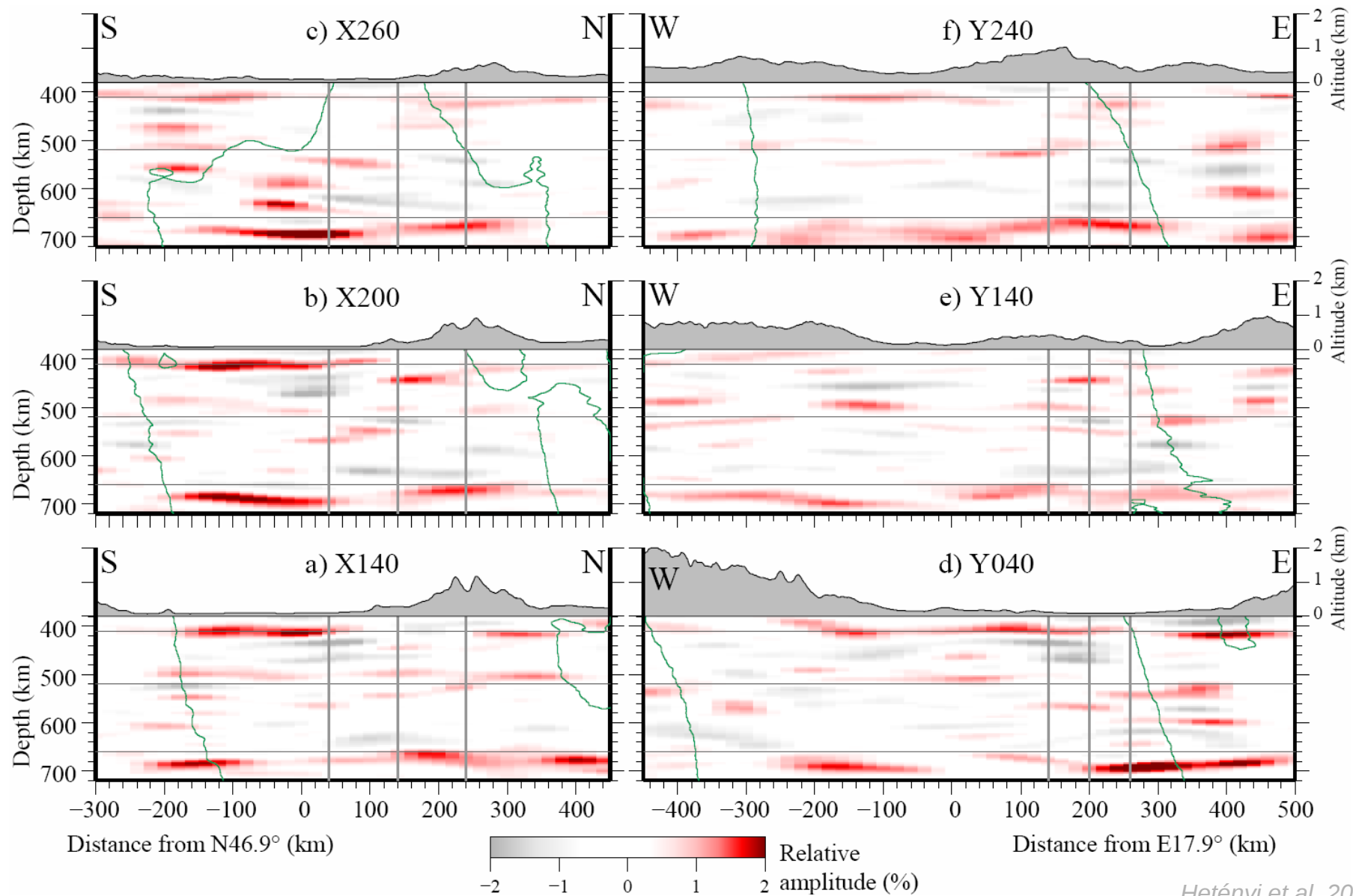
Collaboration: Greg Houseman, Graham Stuart (Univ. Leeds) & CBP

Hetényi et al. 2009

Pannonian Basin



Pannonian Basin



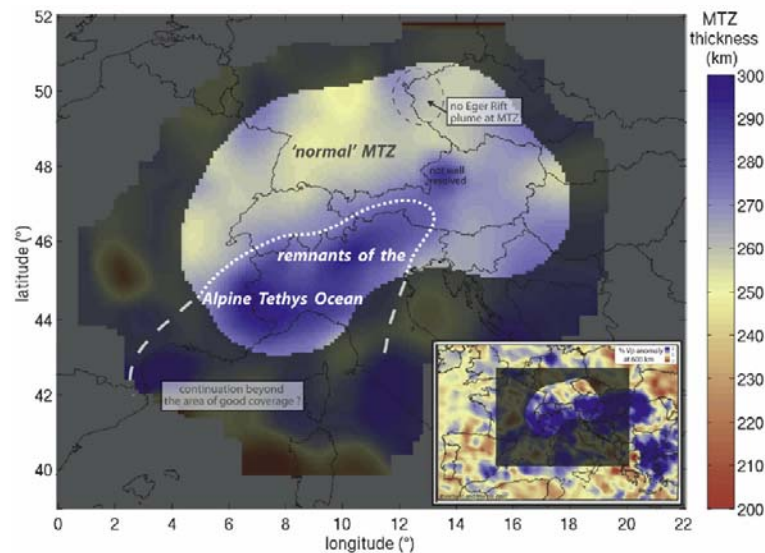
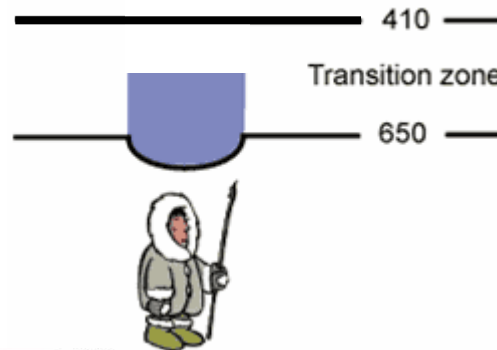
Hetényi et al. 2009

RF, 30-4 sec, CCP migration with regional tomographic model, 20% of max. resolution

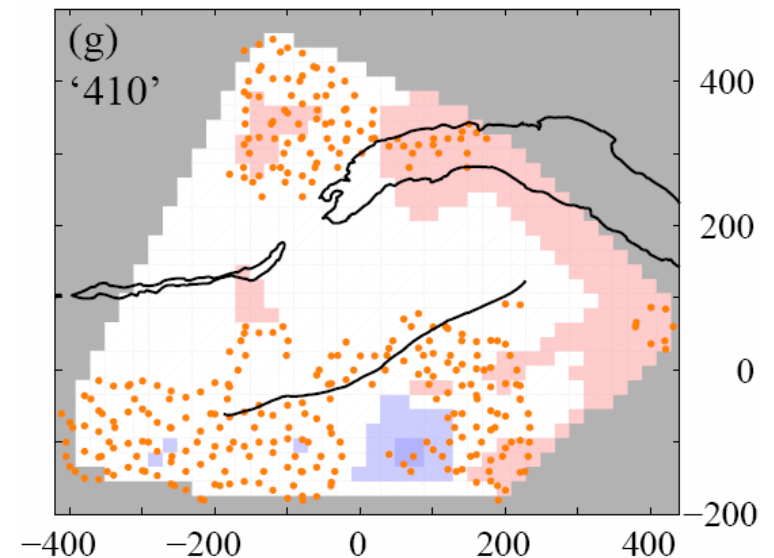
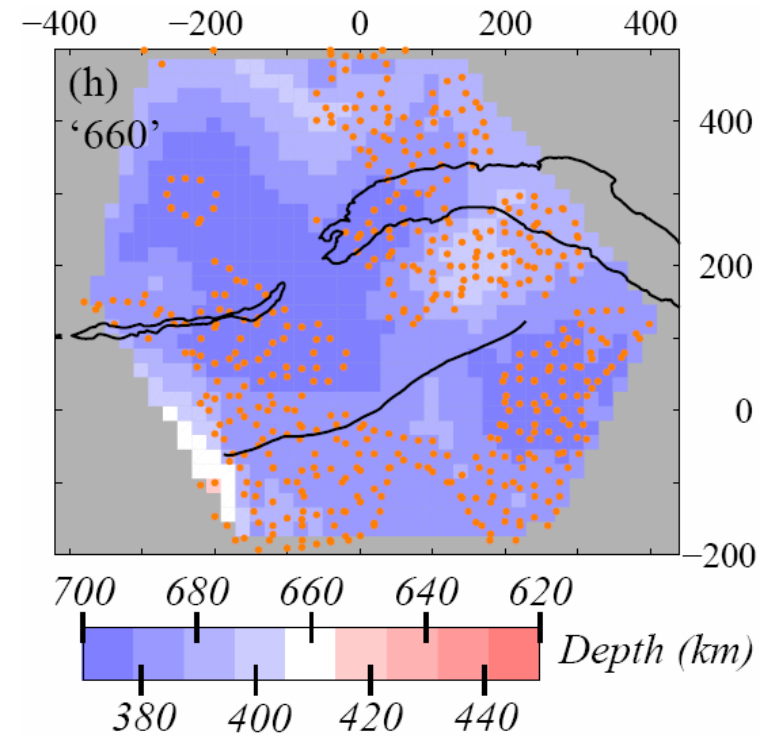
Pannonian Basin

Hetényi et al. 2009

- '410' horizontal, normal depth
- '520' no clear signal
- '660' deepened, significant depth-variation

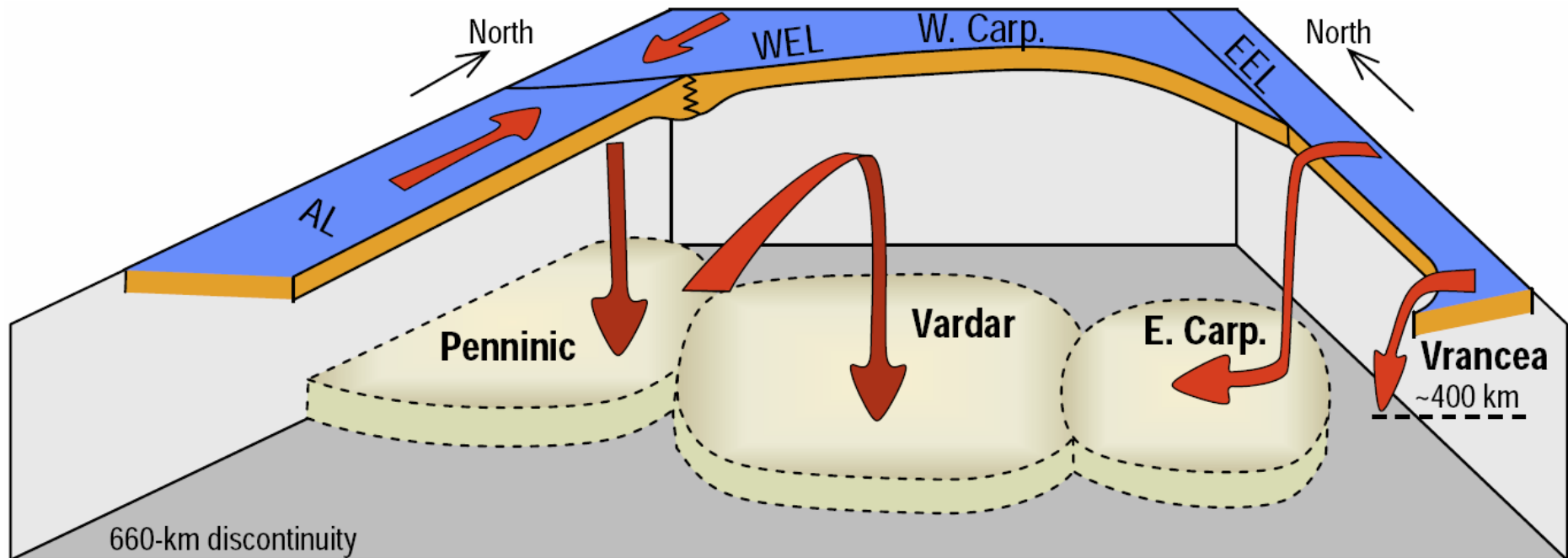


Lombardi et al. 2009



Pannonian Basin

- different Neo-Tethys sub-basin fragments accumulate since Cretaceous



Hetényi et al. 2009

- 40 km thickening = 300 K temperature anomaly throughout the MTZ
- normal '410' + deflected '660': 800 K ???
- how to explain the strong lateral variations ???

Other transition zone thickness variations

Tauzin et al. 2008 [RF]:

- overall MTZ thickness variations: 35-40 km
- no clear evidence for thinning at hotspots: '410' deepened, '660' less T-sensitive

Huerta et al. 2009 [RF]:

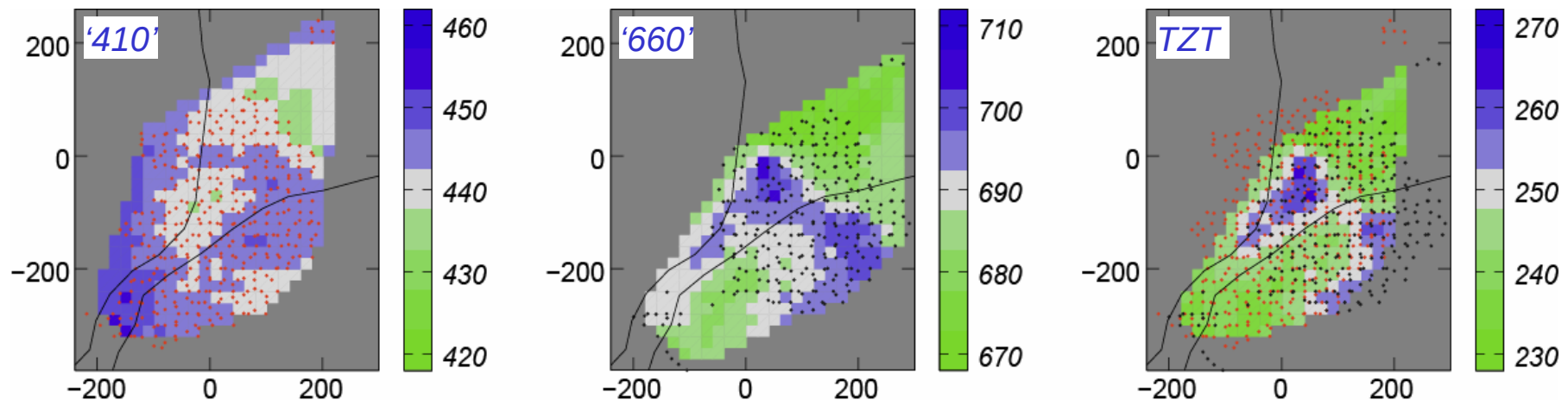
- East-African rift: '410' depressed (350°C hot anomaly), '660' also depressed

Day & Deuss 2010 AGU [PP and P'P' precursors]:

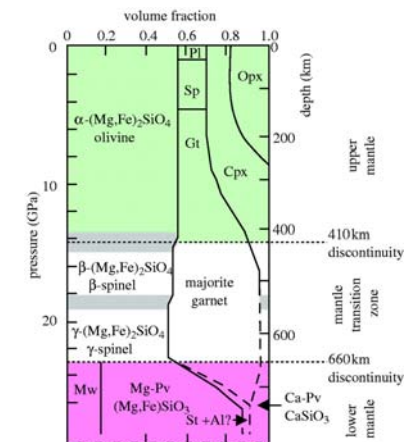
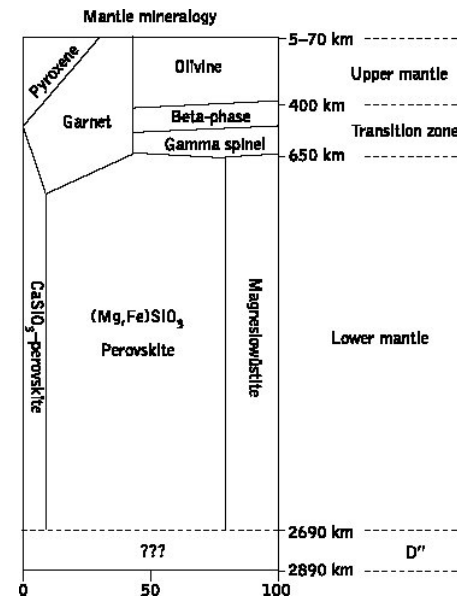
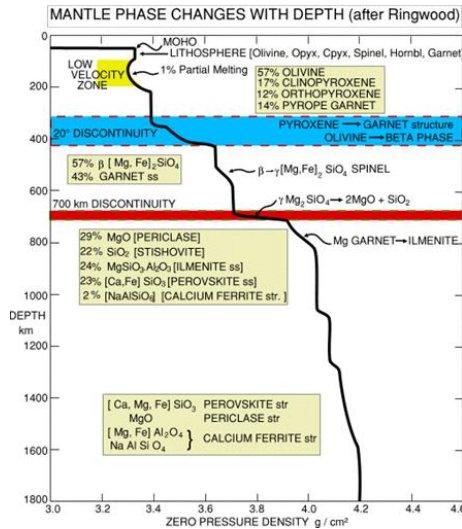
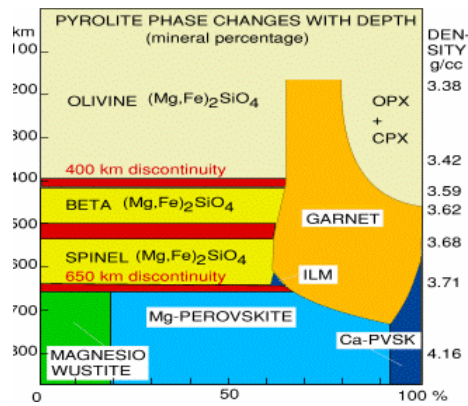
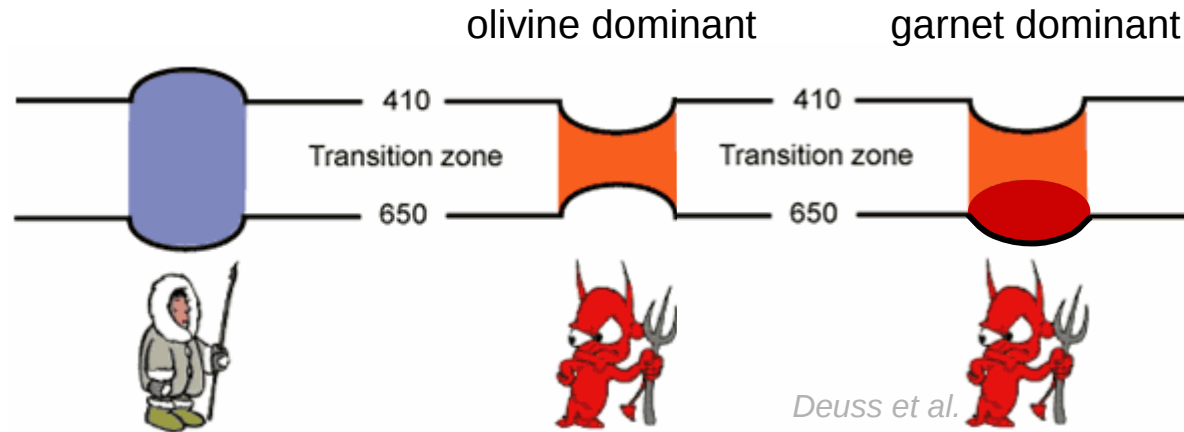
- '660': up to 35 km depth variation within few hundred lateral km
- double peak: 720 km depth: majorite garnet → Ca-perovskite phase change

Cornwell et al. (2010 AGU; paper in prep.) [RF]:

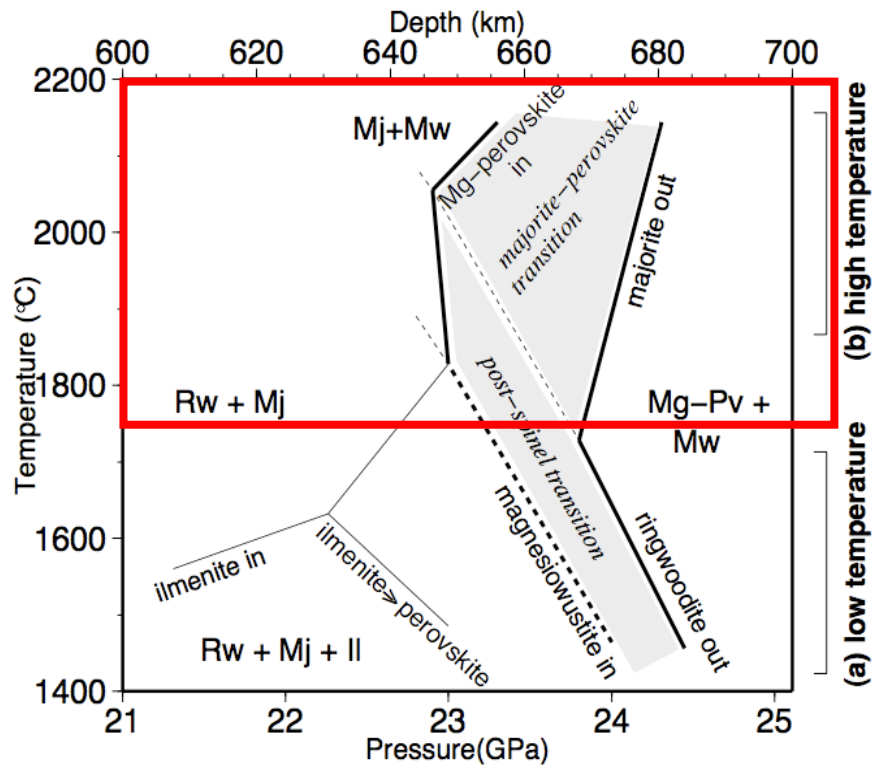
- Ethiopia: depressed '410', '660' also deepened and shows lateral variations



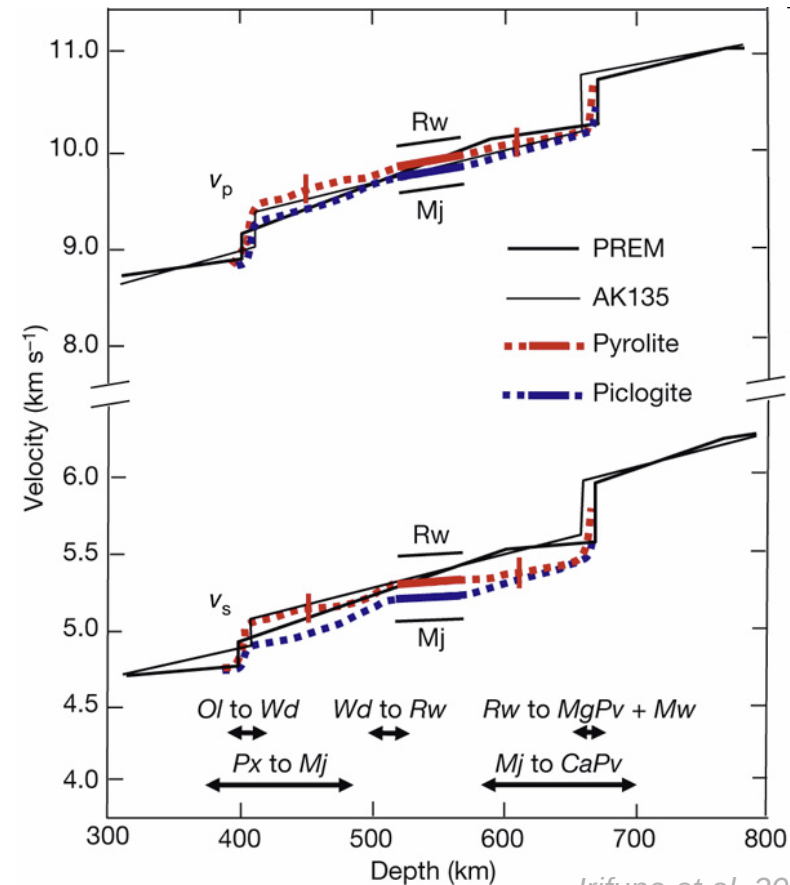
Phase change s control the transition zone



Phase change s control the transition zone



Hirose 2002

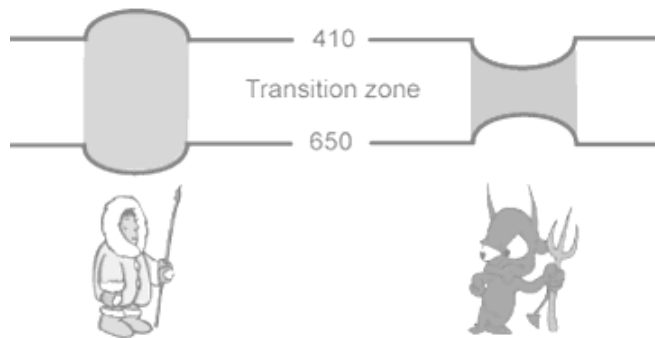


Irifune et al. 2008

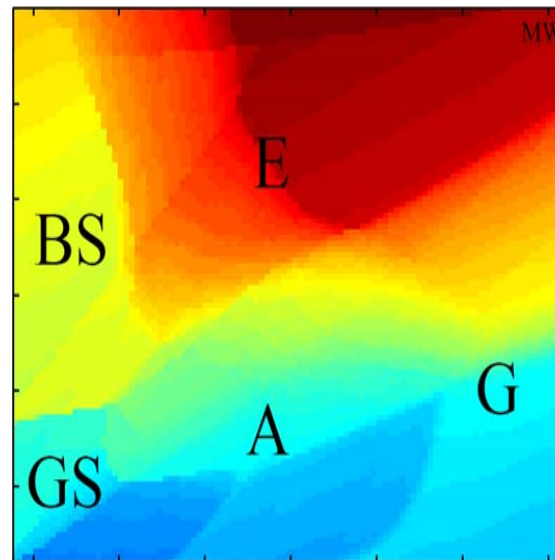
- 1) Phase changes of olivine and garnet shape the depth of discontinuities
- 2) Garnet Clapeyron-slope is positive → '660' more complex
- 3) Lateral variations likely to be explained by compositional heterogeneities

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modélisation géodynamique



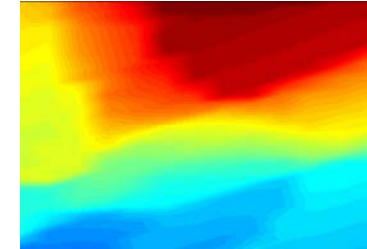
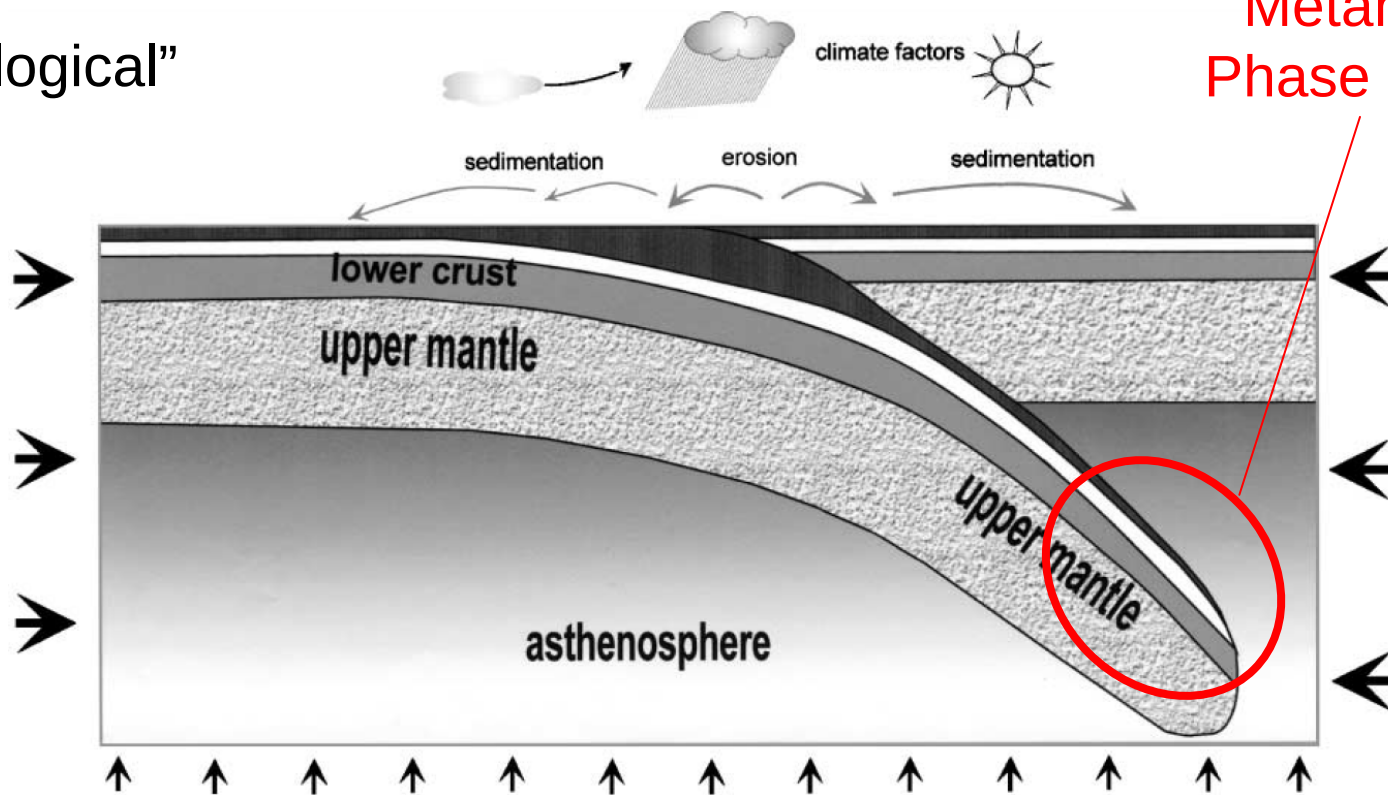
prismation des orgues de lave



Consistent implementation of phase changes into geodynamic models

Physical processes and laws

- mechanical
- thermal
- “geological”



Metamorphic
Phase Changes

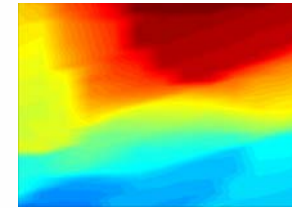
Burov et al. 2001

Collaboration: Vincent Godard (CEREGE), Rodolphe Cattin (Géosciences Montpellier), James A. D. Connolly (ETH)

Metamorphism and deformation

$t_0(P_0, T_0)$

initial
state

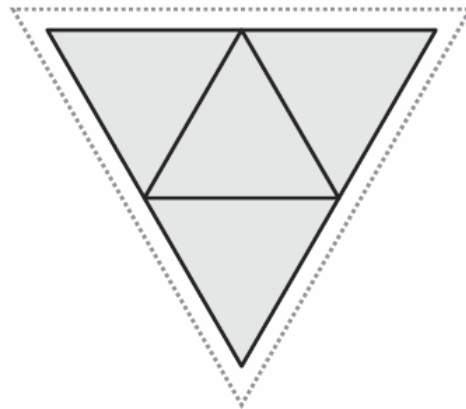


metamorphism: 0
deformation: δV

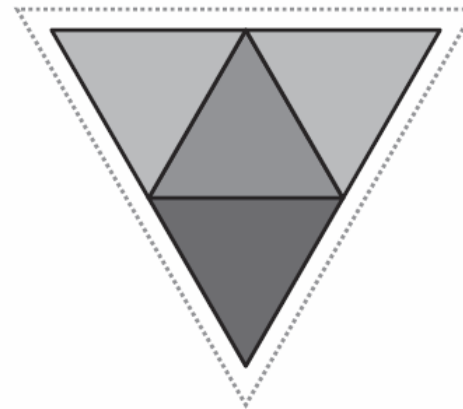
metamorphism: $\delta \rho$
deformation: δV

metamorphism } $\delta \rho$
deformation } δV

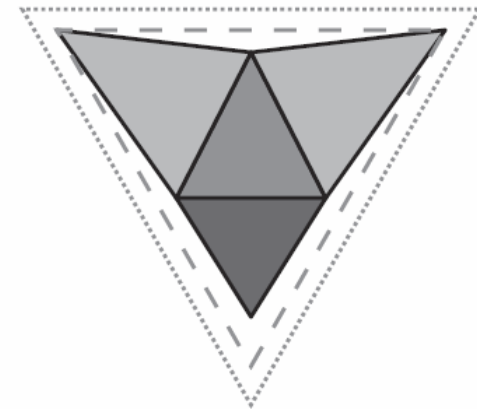
$t_1(P_1, T_1)$



a) *Simple
Boussinesq*



b) *Boussinesq
& metamorphism*



c) *Mass conserved
& metamorphism*

Equations to solve

- equilibrium: conservation of momentum

$$\nabla \cdot \sigma_{ij} + \mathbf{f} = m\mathbf{a}$$



- heat: conservation of energy

$$\frac{\partial T}{\partial t} - \kappa \cdot \Delta T - A + \mathbf{v} \cdot \nabla T = 0$$



- continuity: conservation of mass

$$\frac{d\rho}{dt} + \nabla \cdot (\rho \mathbf{v}) = 0$$



Simplification

» Cette indépendance subsiste même quand, supposant le fluide incompressible (ce qui n'est nullement obligé, même pour un liquide), on s'impose de ne choisir que des déformations compatibles avec la conservation parfaite des volumes aux divers instants. En effet, celle-ci revient, comme on sait, à établir, entre les vitesses effectives de dilatation dans les sens des axes, la relation linéaire

$$(4) \quad \frac{du}{dx} + \frac{dv}{dy} + \frac{dw}{dz} = 0;$$

Boussinesq 1897

$$\cancel{\frac{dp}{dt}} + \nabla \cdot (\cancel{p}\mathbf{v}) = 0$$

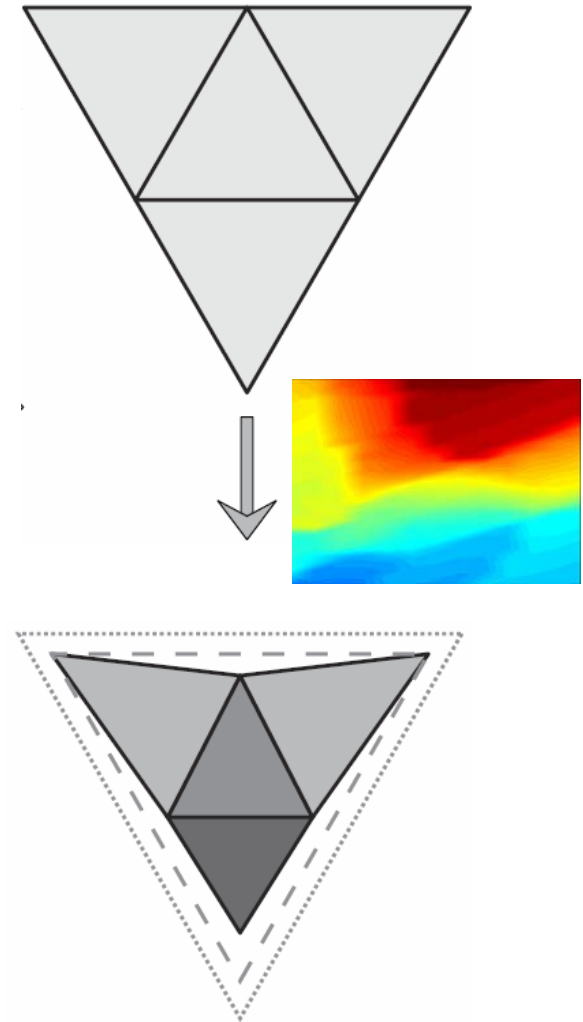
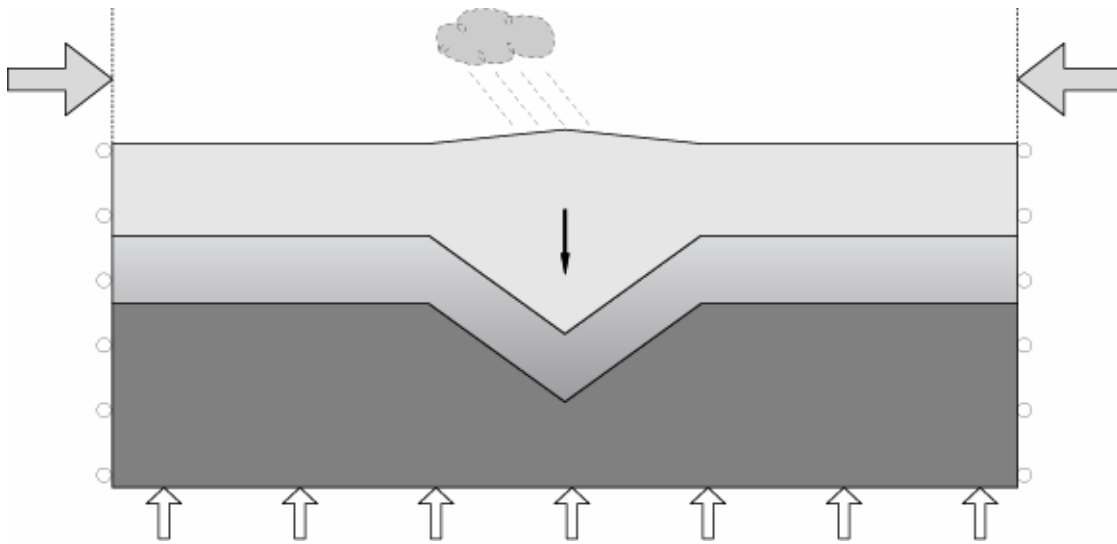
incompressibility

Goal

- implement

$$\frac{d\rho}{dt} + \nabla \cdot (\rho \mathbf{v}) = 0 \quad \checkmark$$

- explore the consequences



Modelling tool: Castem

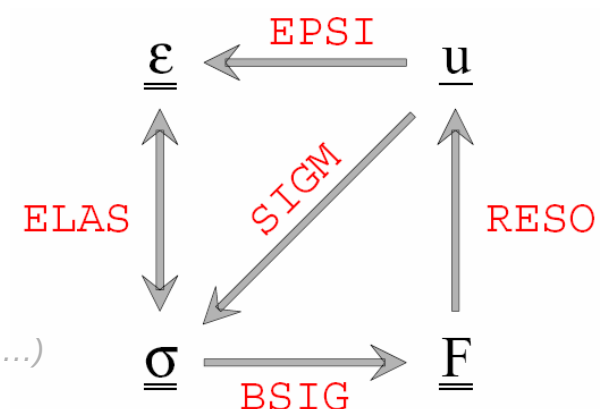
- finite element modelling tool (2D, 3D)
- mechanical and thermal problems
- pre-defined operators: flexibility ☺, efficiency ☹



<http://www-cast3m.cea.fr>

Geodynamics toolbox (Godard, 2006)

- rheologies (elastic, viscous, plastic)
- erosion (diffusive, fluvial incision)
- benchmarked against ADELI (J. Chéry, R. Hassani, ...)
- re-meshing



Implementation of mass conservation

Mass to conserve:

$$m_c = \rho_1 V_1 = \rho_0 V_0$$

$$\Delta V / V = -\Delta \rho / \rho$$

$$\Delta V / V = tr \underline{\underline{\epsilon}} = \epsilon_{11} + \epsilon_{22} + \epsilon_{33}$$

$$\epsilon_{ii} = \Delta V / 3V = -\Delta \rho / 3\rho$$

Stress modification:

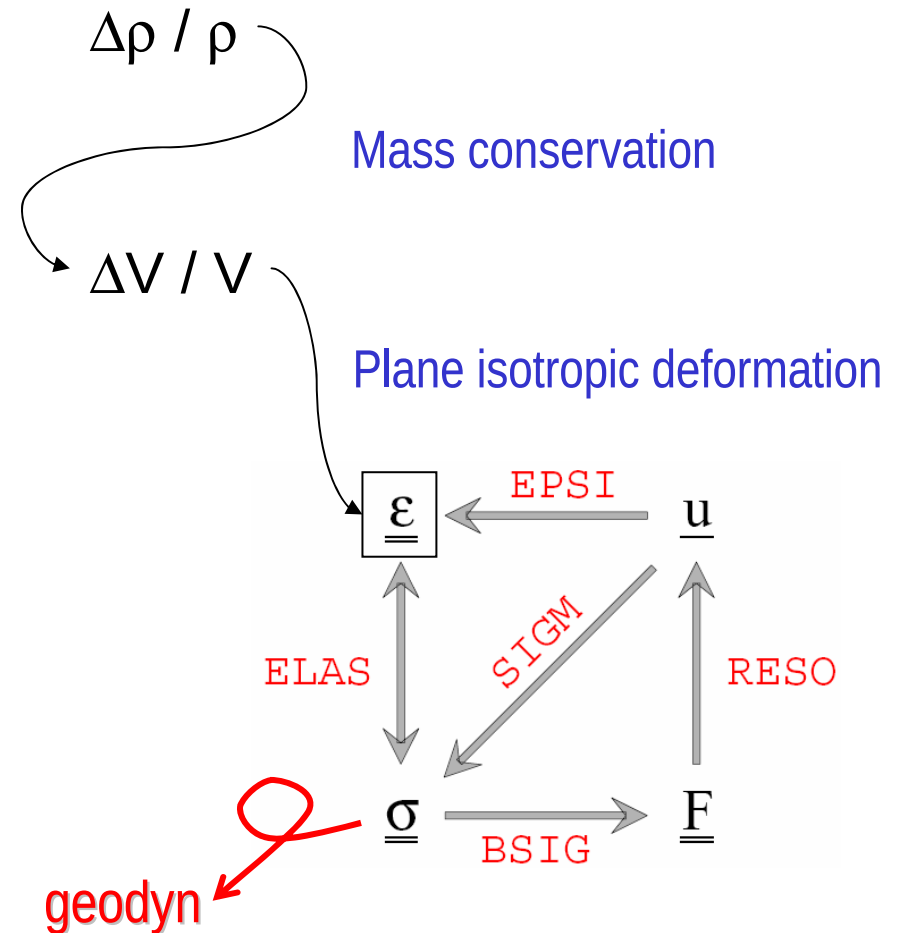
- subtract from internal stress field

$$\underline{\underline{\sigma}}_{int} = \underline{\underline{\sigma}}_{int} - \underline{\underline{\sigma}}_{MPC}$$

- “numerical” damping

$$\underline{\underline{\sigma}}_{int} = \underline{\underline{\sigma}}_{int} - \underline{\underline{\sigma}}_{MPC} / \mathbf{F}$$

- stability tests

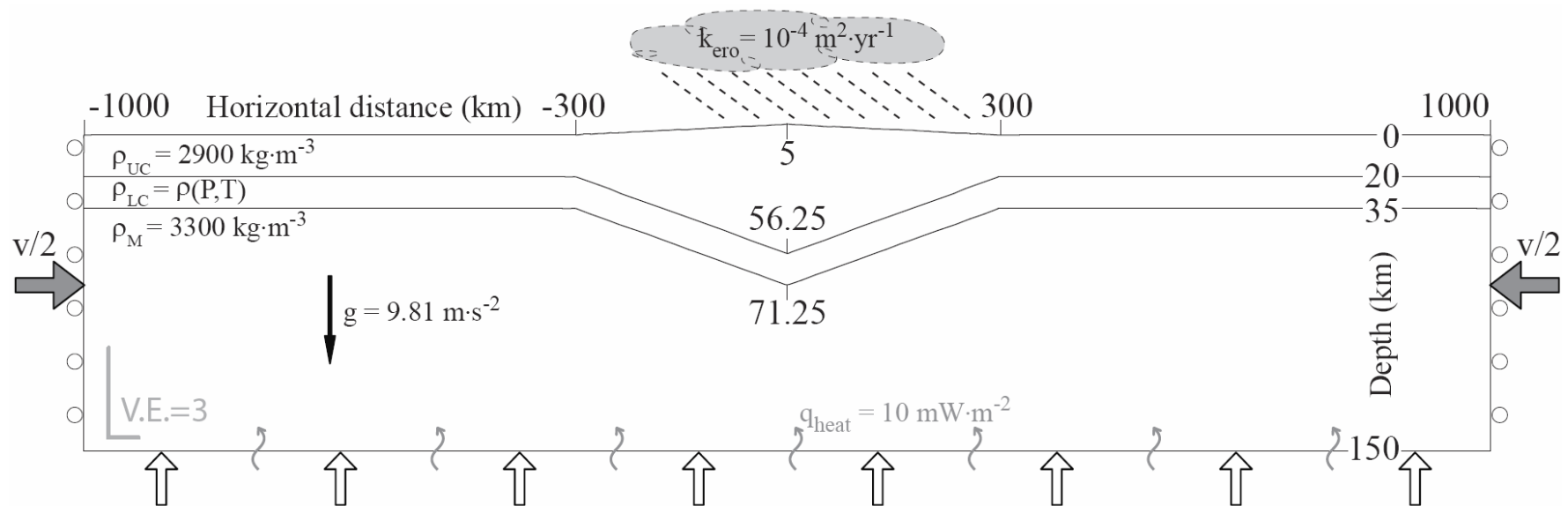
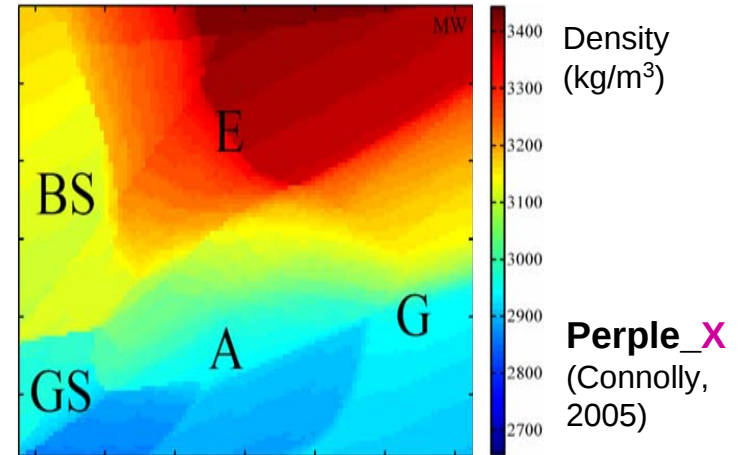


Benchmark setup

2D continental lithosphere with topography & isostatic crustal root

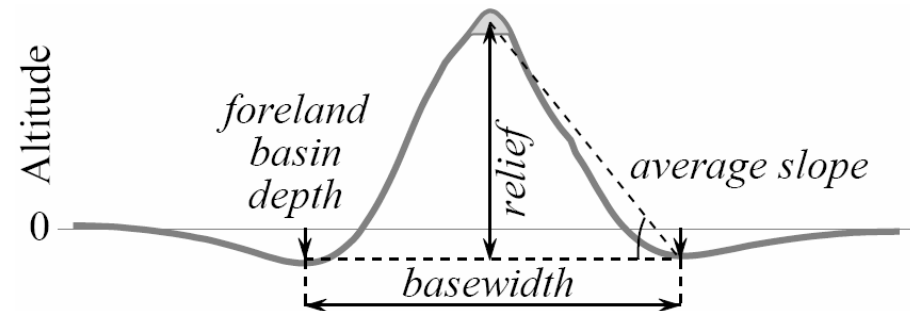
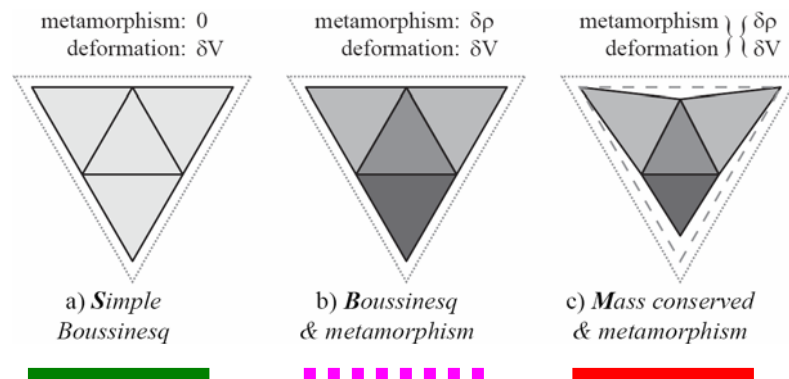
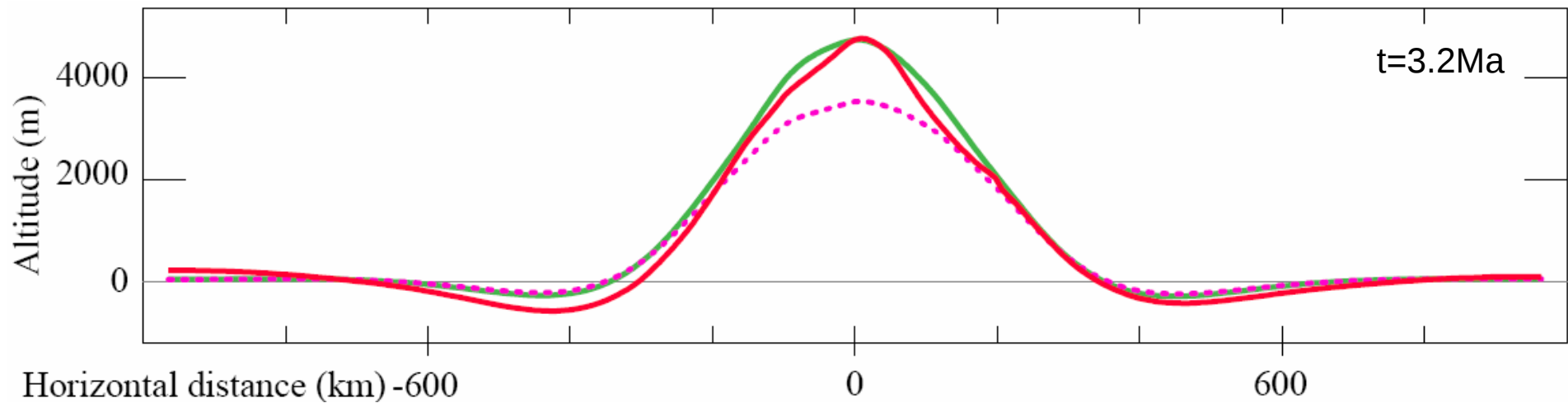
- visco-elastic rheology
- horizontal convergence: 20 mm/yr
- lower crustal metamorphism
- linear diffusive erosion

$$\frac{\partial h}{\partial t} = -k_{ero} \cdot \frac{\partial^2 h}{\partial x^2}$$



Effect of metamorphism

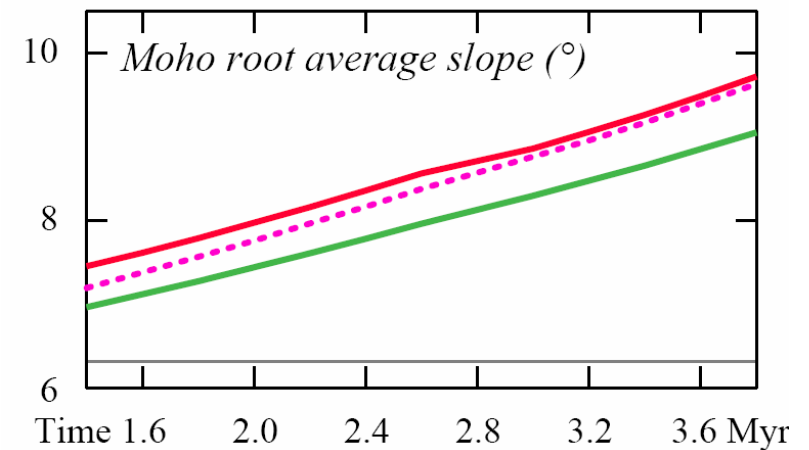
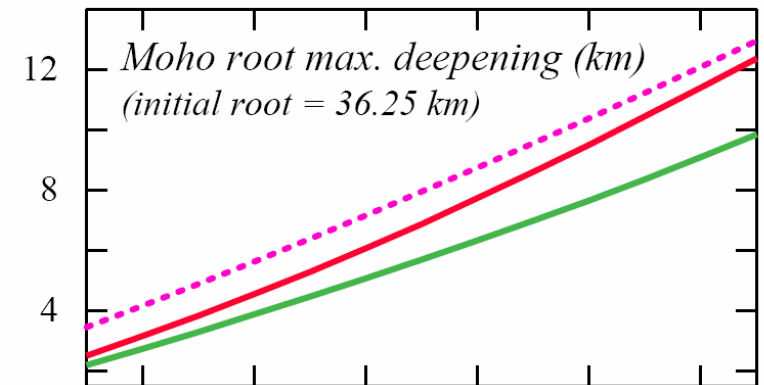
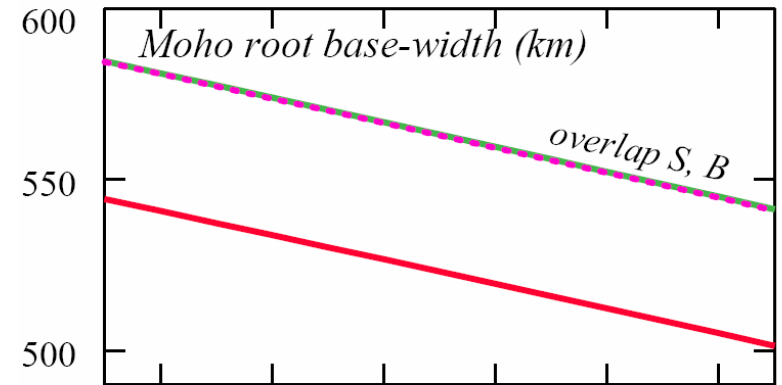
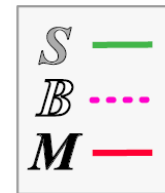
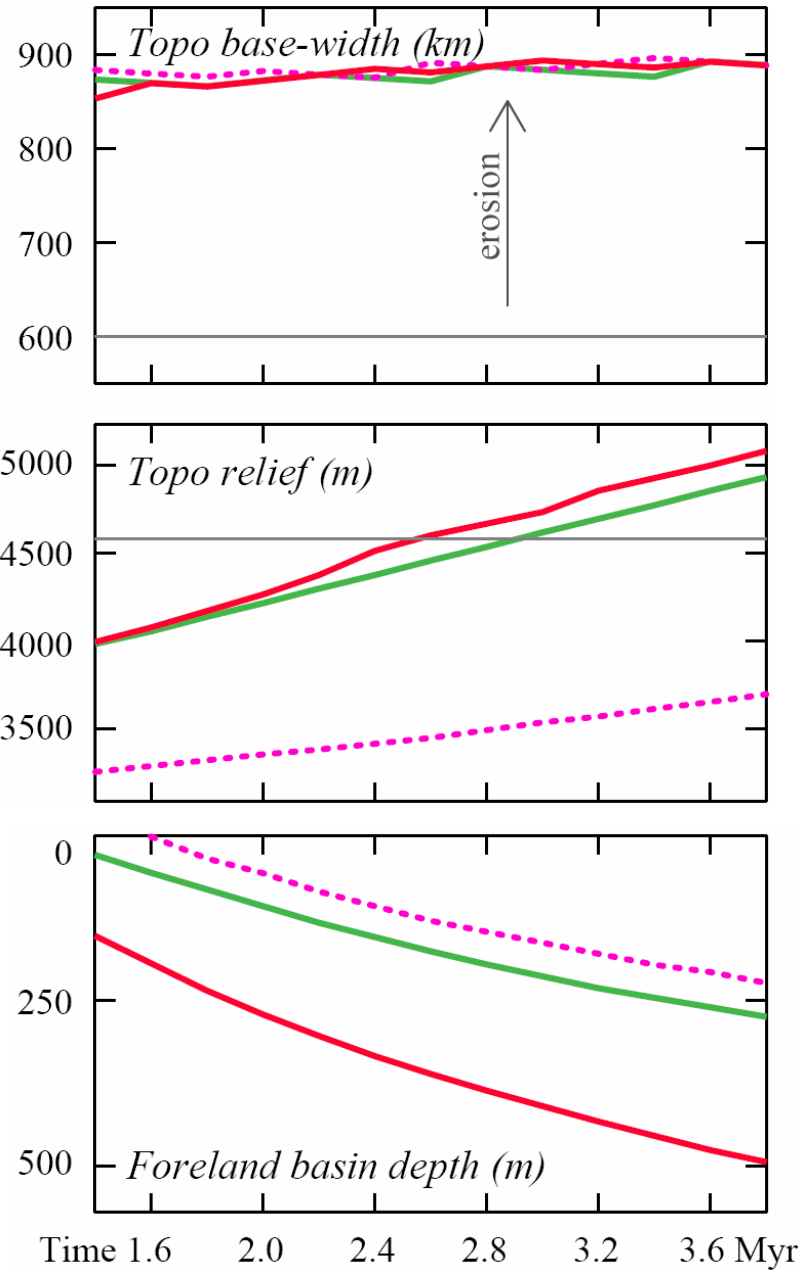
with visco-elastic rheology, erosion, convergence



Topography

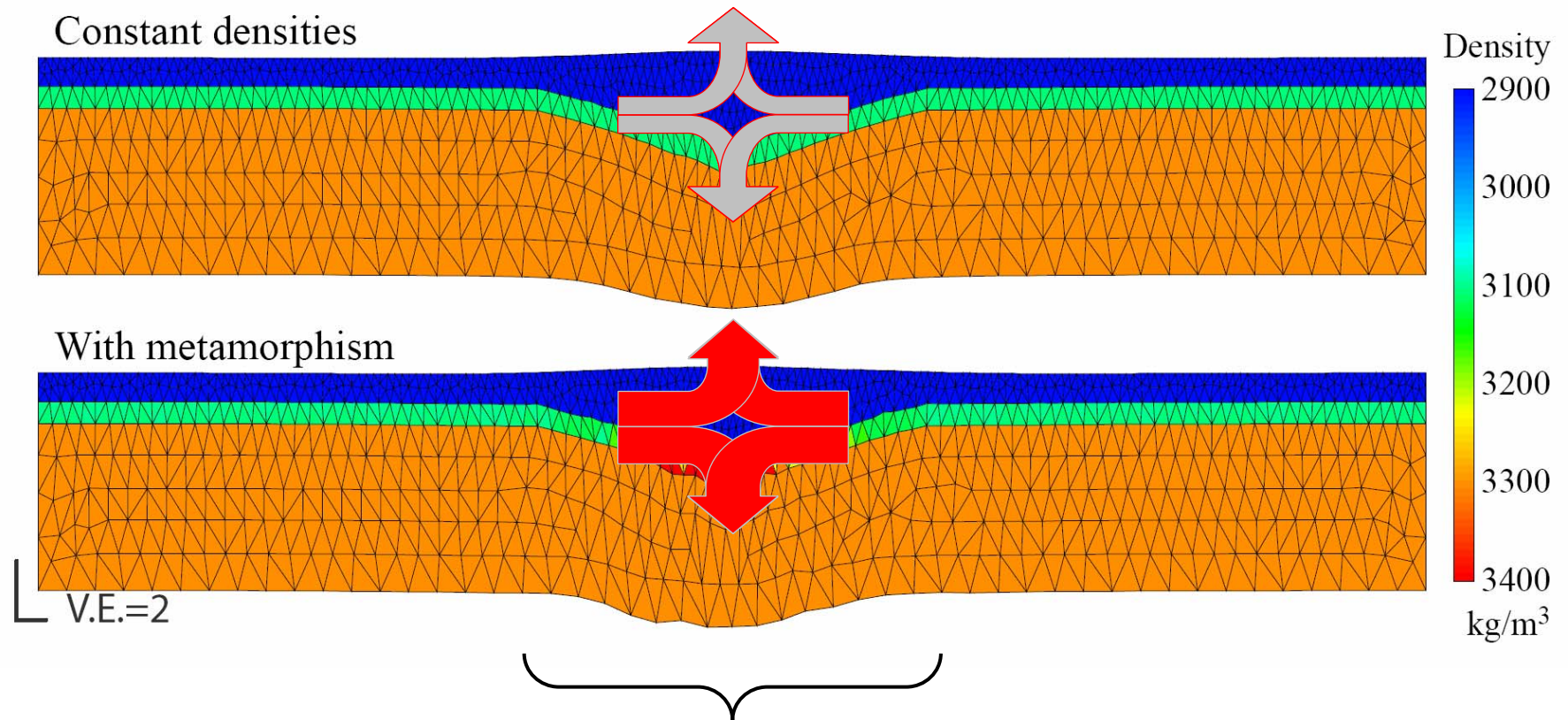
Effect of metamorphism

Crustal root



Effect of metamorphism

Density profiles



Conclusions and implications

Metamorphism in orogeny:

- deformation localization
- amplitudes similar to erosion

May need revision:

- erosion rates
- basin evolution
- flexural studies

Respect continuity !!!

Earthquake triggering ($\sigma_{MPC} \approx$ few bars)

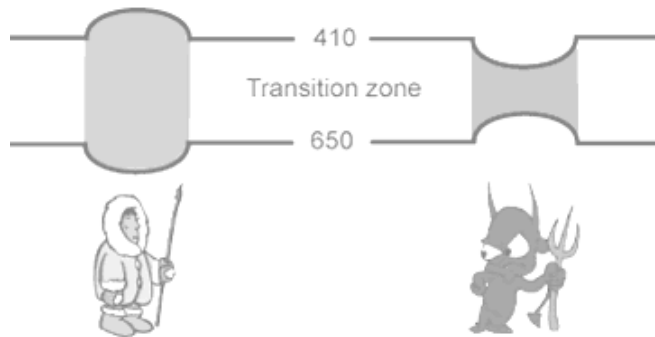
[e.g., eclogitization beneath Tibet]

Delamination / break-off

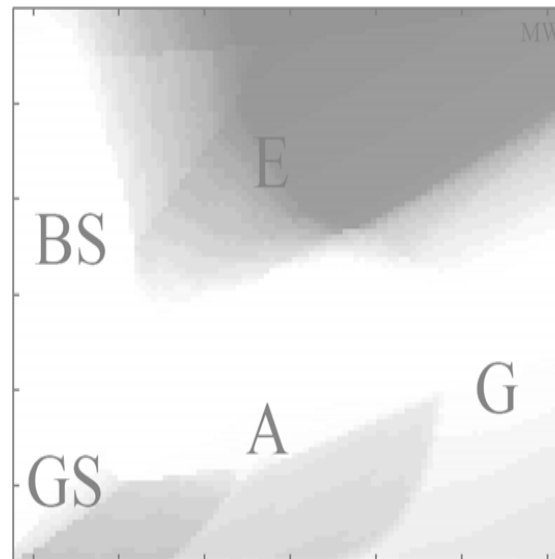
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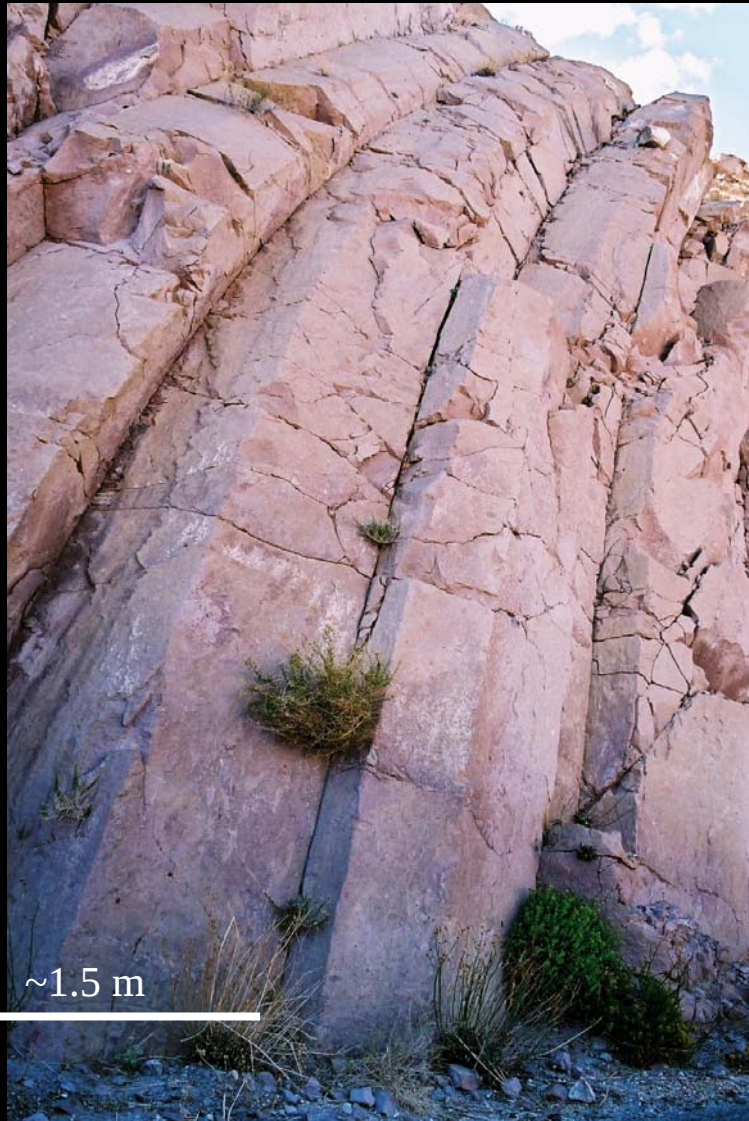
modélisation géodynamique



prismation des orgues de lave



Where this study has started...



Bishop Tuff (CA, USA) – rhyolite

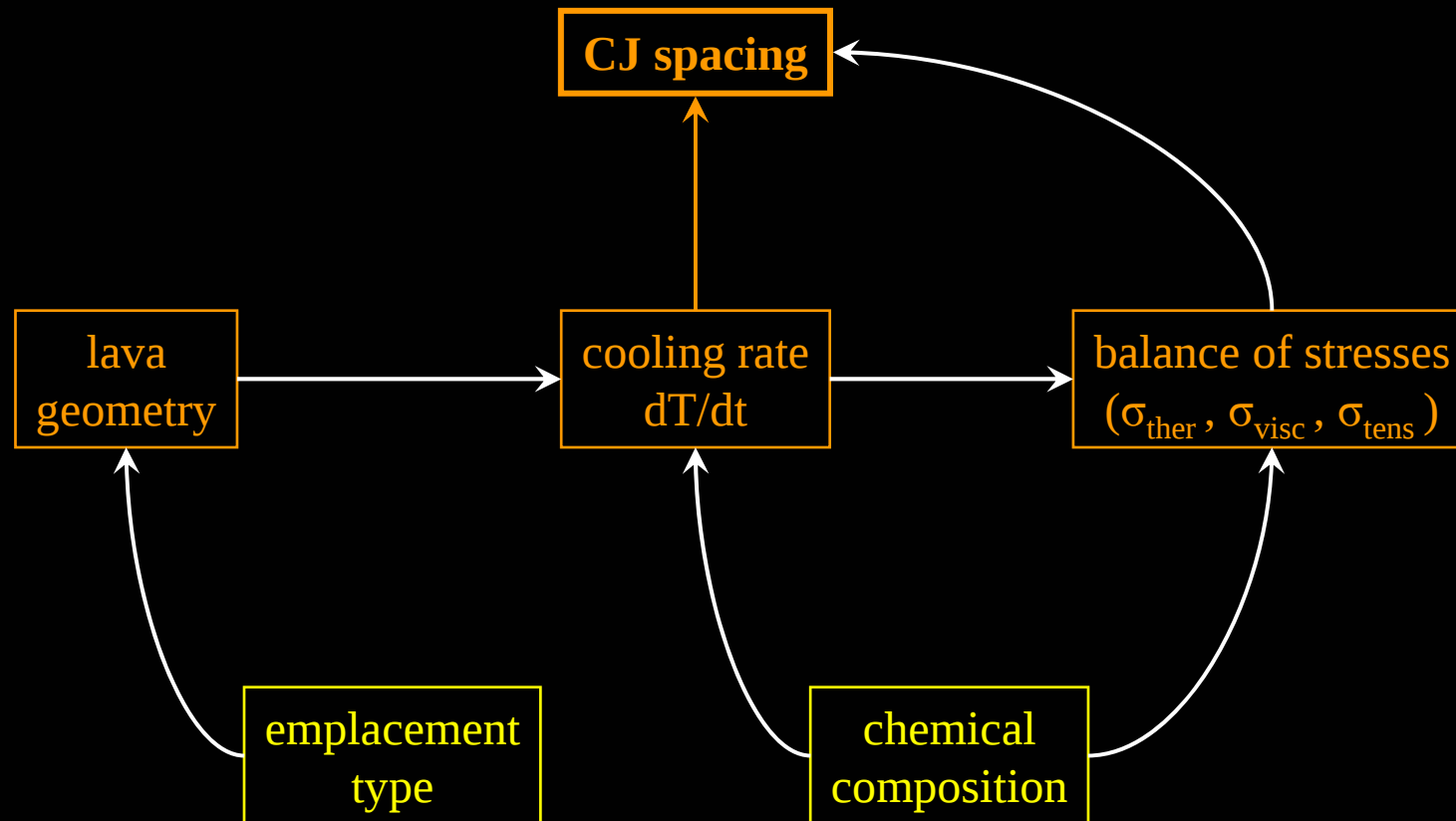


Somoskő [Šiatorská Bukovinka] (H-SK) – basalt



Szanda (H) – andesite

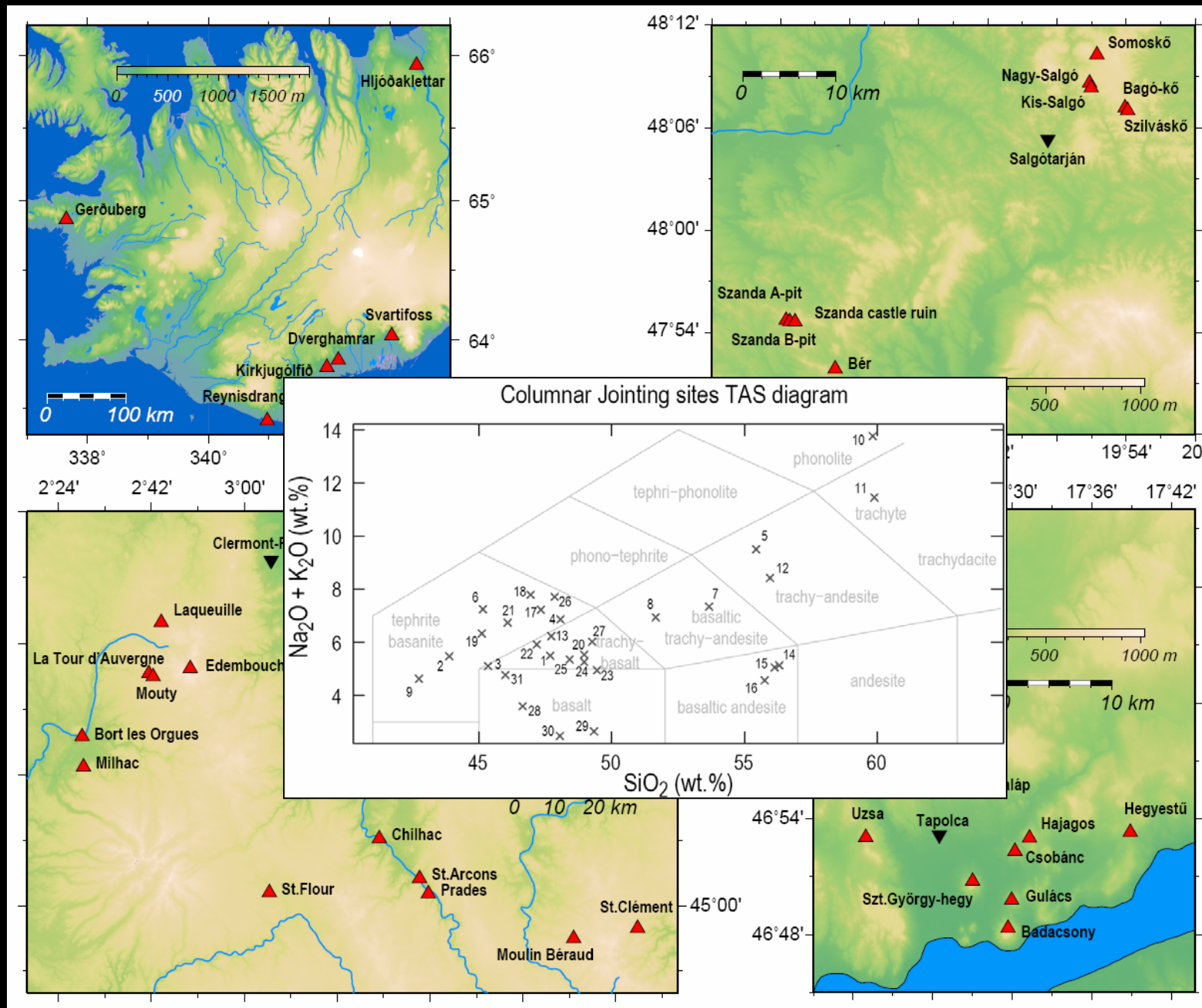
What controls the size of columns?



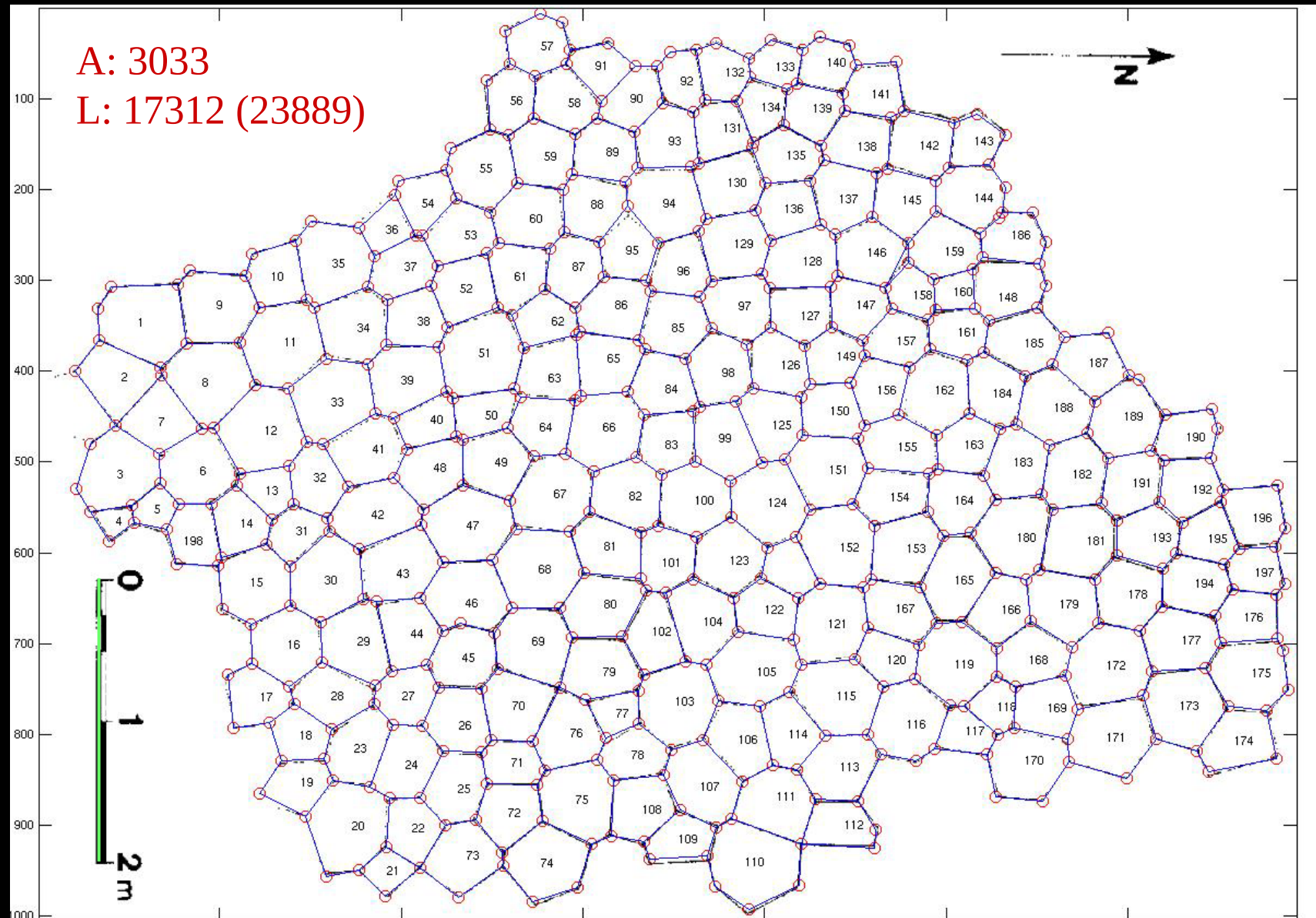
Field studies in search for: characteristic size, emplacement type, composition...

Collaboration: Etienne Médard (Clermont-Ferrand), Fanny Garel and Benoît Taisne (IPGP)

Columnar jointing sites and compositions

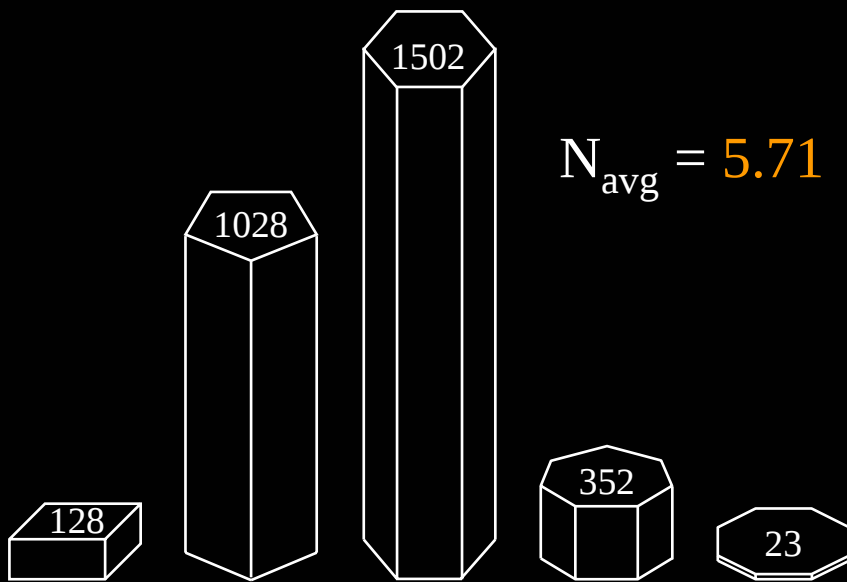
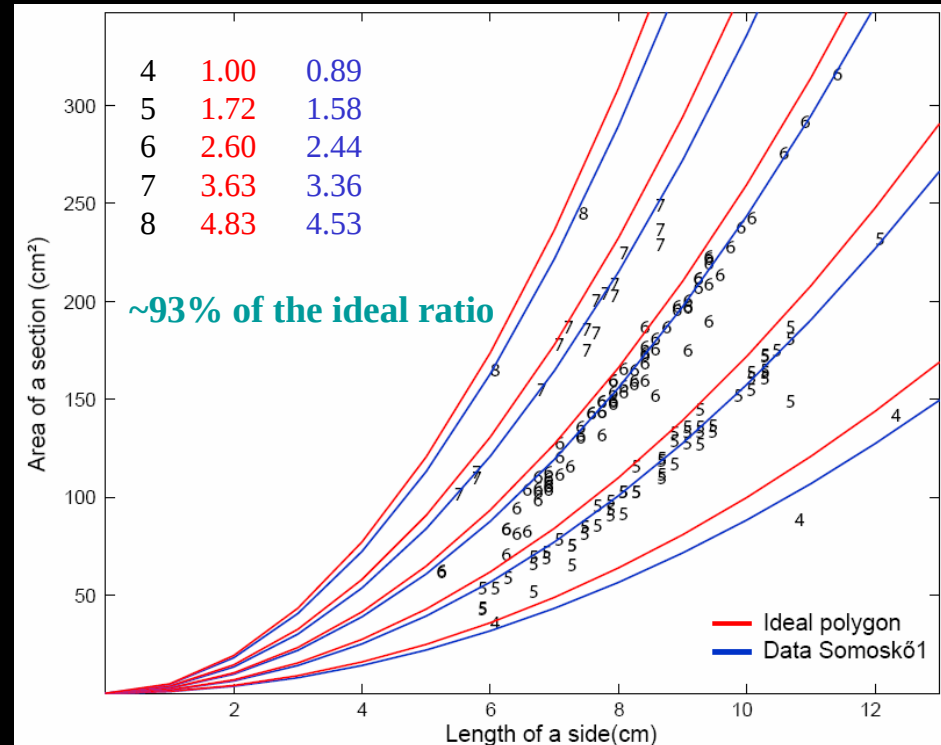
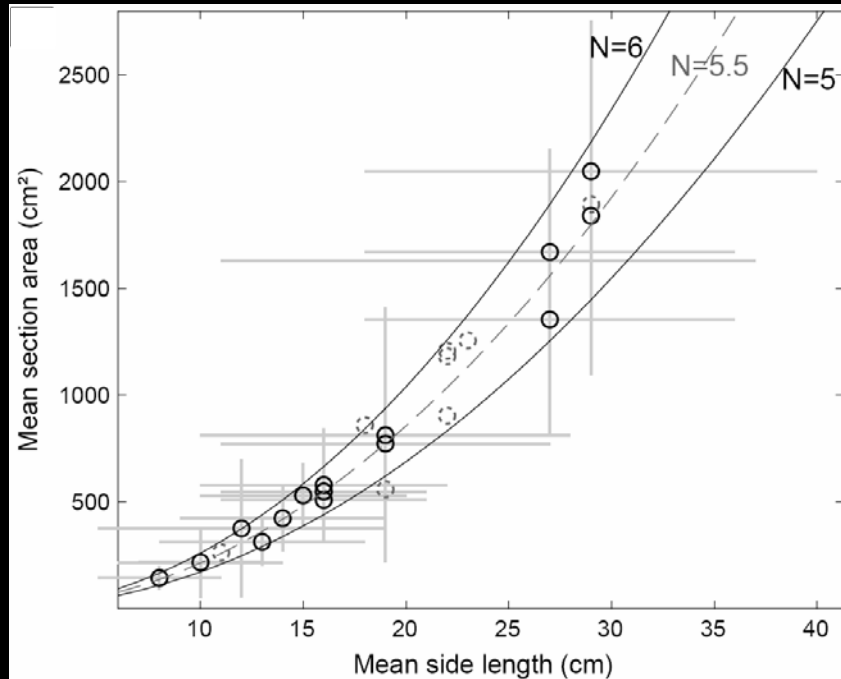


Field measurements: photos $\rightarrow$ software $\rightarrow$ statistics on L, A, N



Background drawing from O'Reilly, 1879 (Trans. R. Irish Acad., 26: 641-728; cited in Budkewitsch & Robin, 1994)

Some results on geometry



- areas reduced compared to ideal polygons (“distorted” columns)
- $N_{avg} < 6$

→ energy spent on creating joints is not minimized

→ dynamic process

Field measurements: simple flow



Saint-Flour (F)

Field measurements: ponded flow



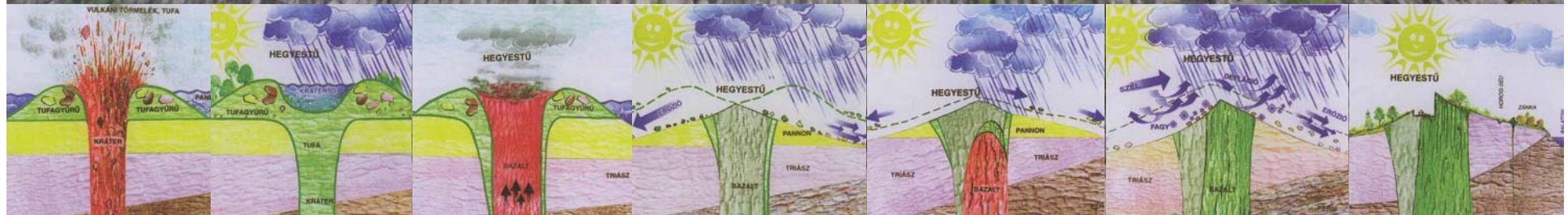
Prades (F)

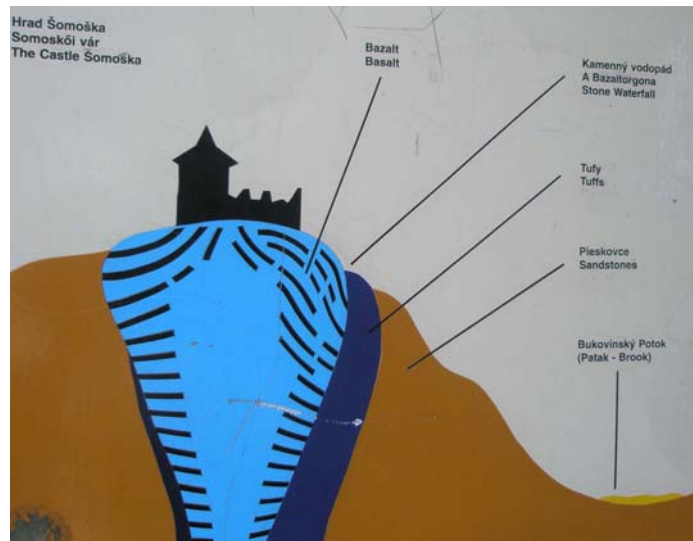
Field measurements:
dyke



Moulin Béraud (F)

Field measurements: neck/plug

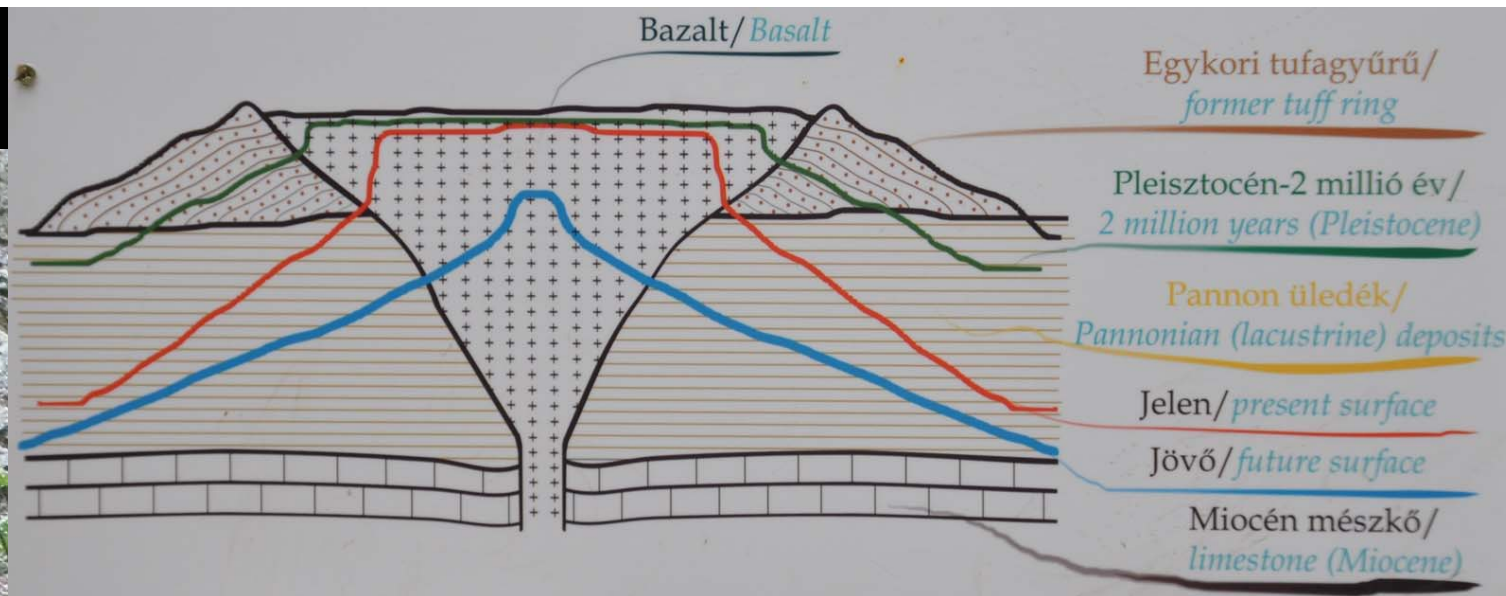




Field measurements: neck/plug

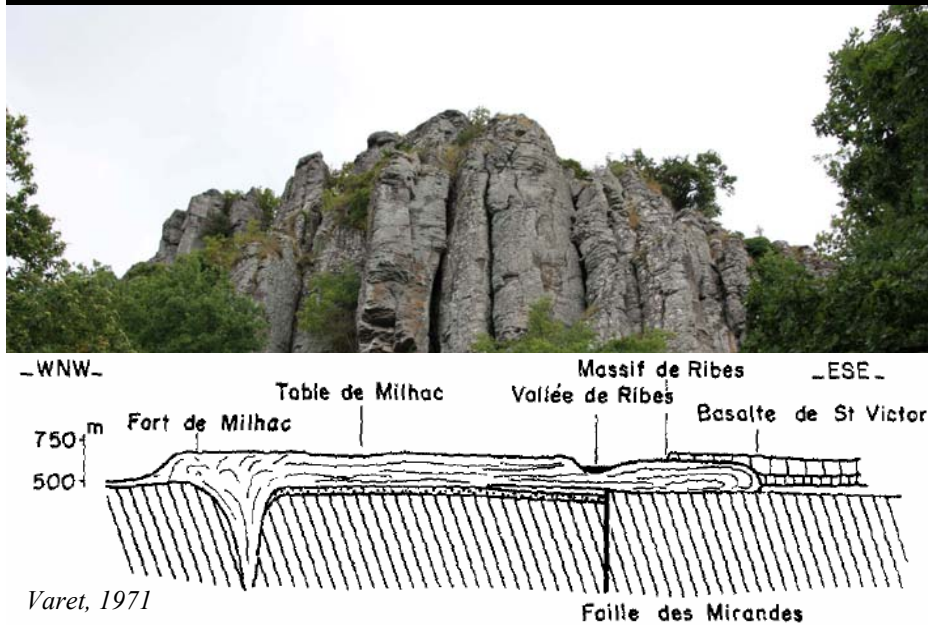


Somoskő [Šiatorská Bukovinka] (H-SK)

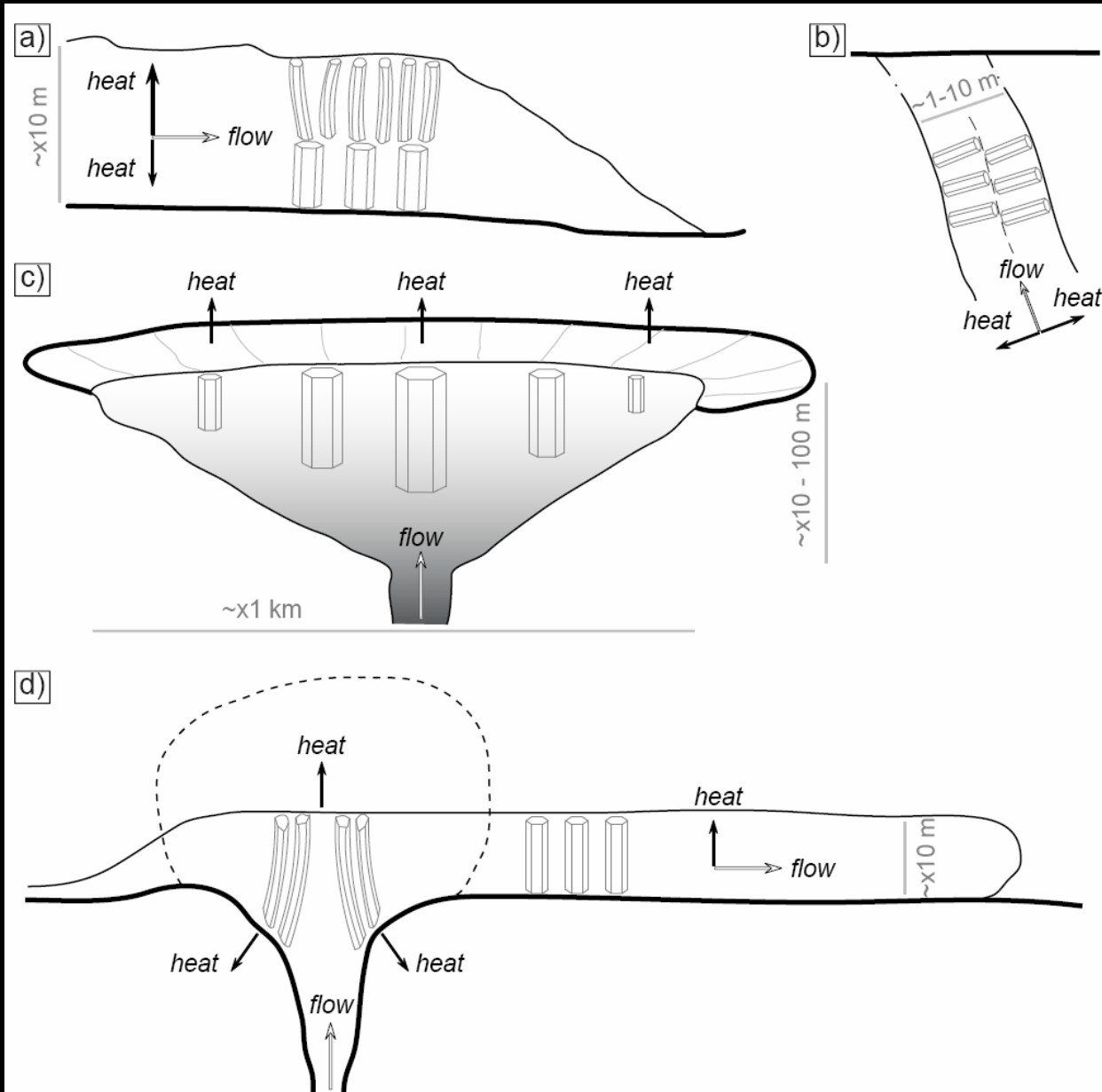


Szent-György-hegy (H)

Field measurements: viscous flow and lava lake



Solidification geometry affects the shape of columns



flow
(ponded)

dyke
(sub-surface)

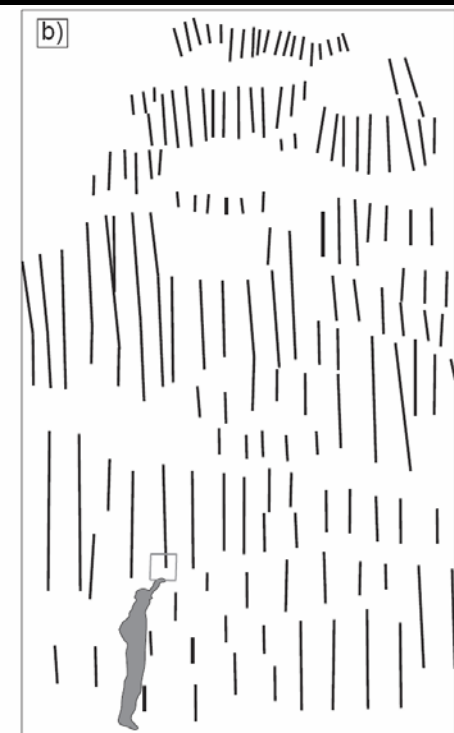
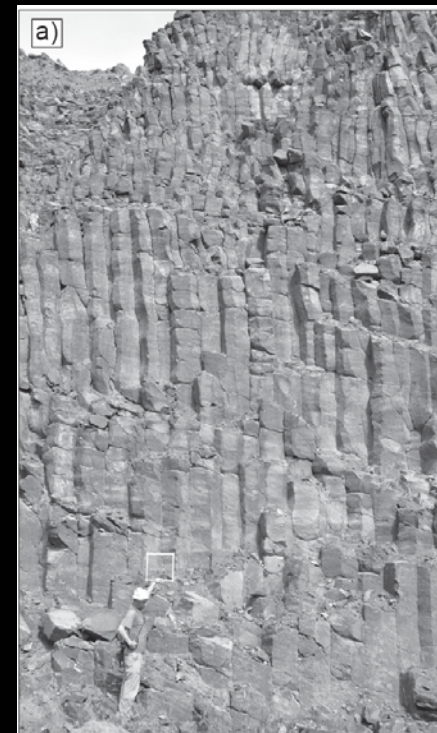
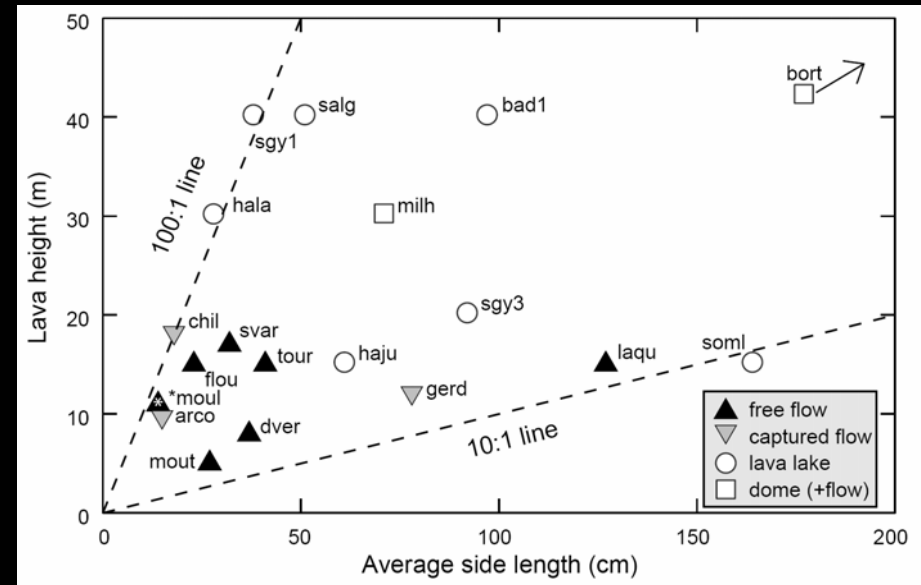
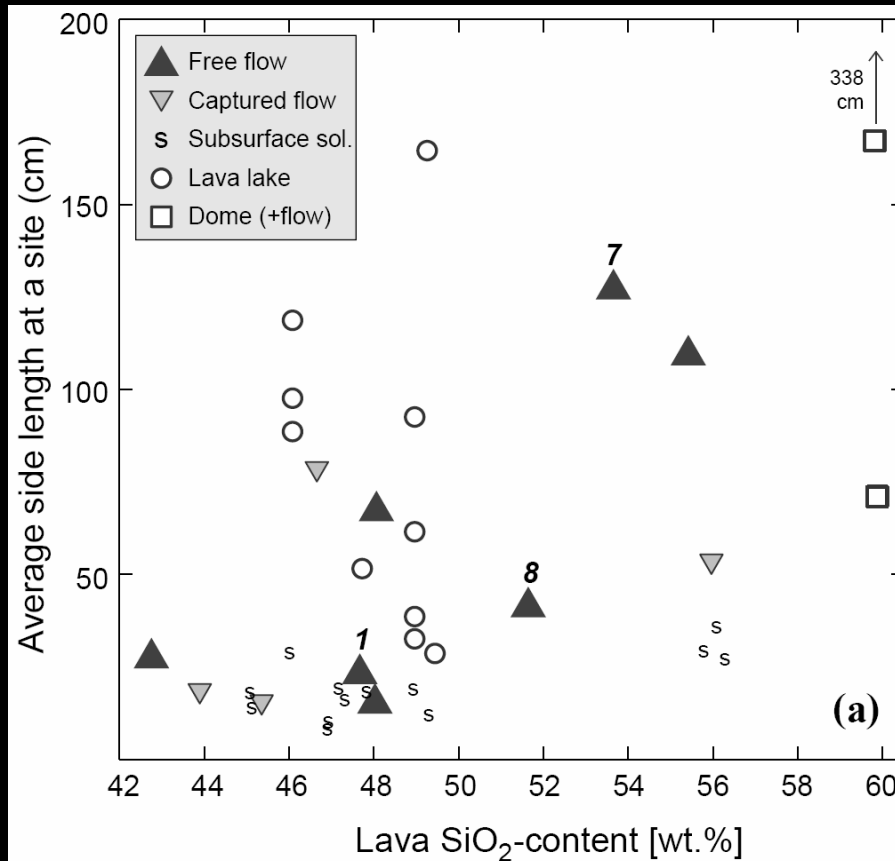
lava lake

dome (+flow)

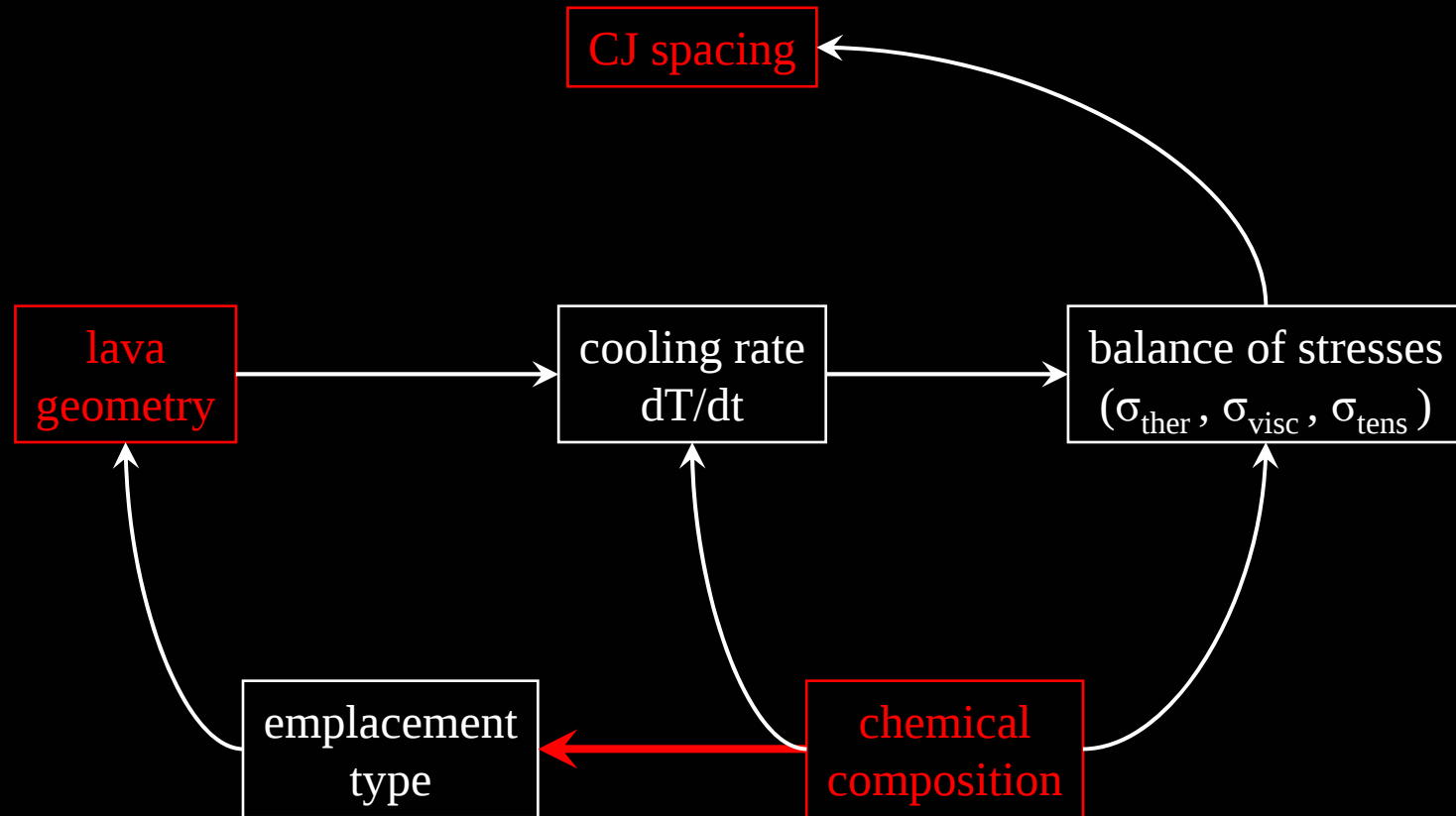
Simple, free flows

Column diameter correlates with:

- SiO_2 -content
- lava flow thickness



CONCLUSIONS



What happens inside a column?

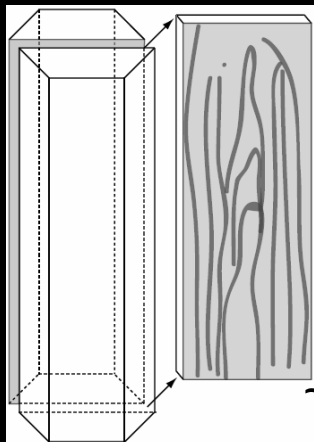
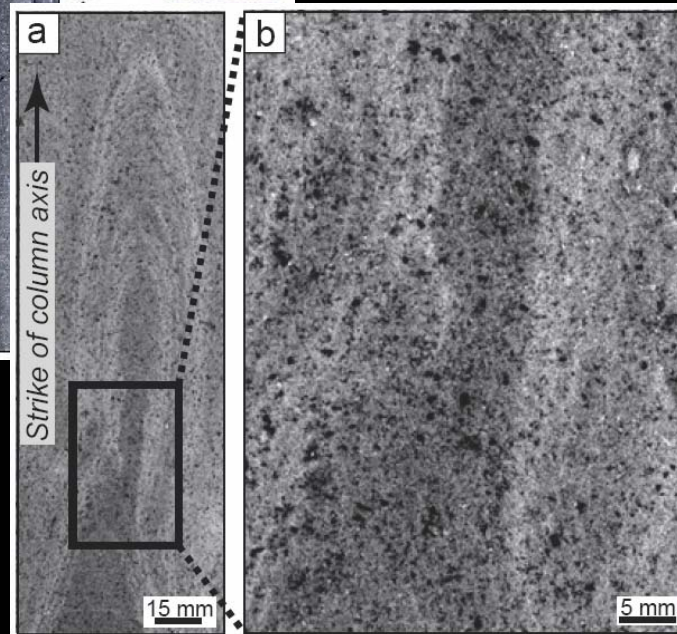
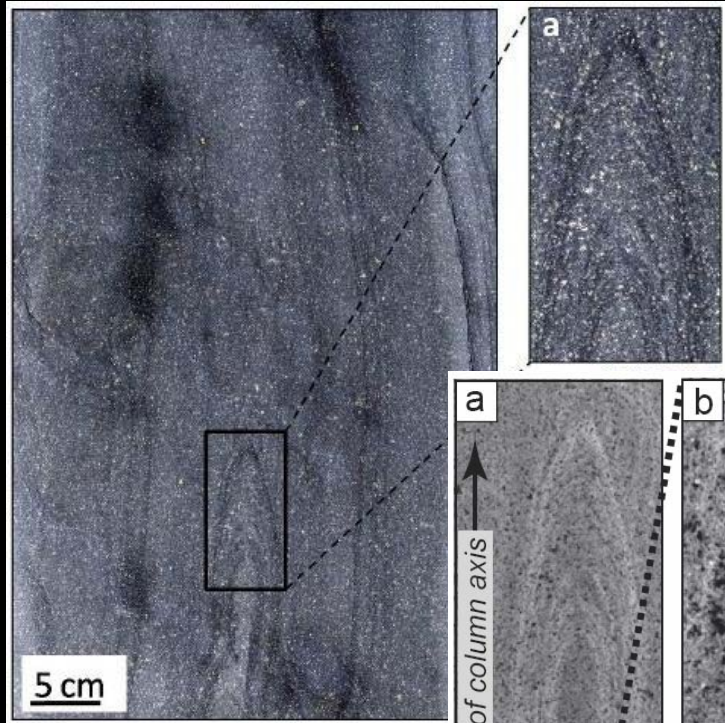


Collaboration: Hannes B. Mattson, Sonja Bosshard, Bjarne Almqvist, Ann M. Hirt (ETH)

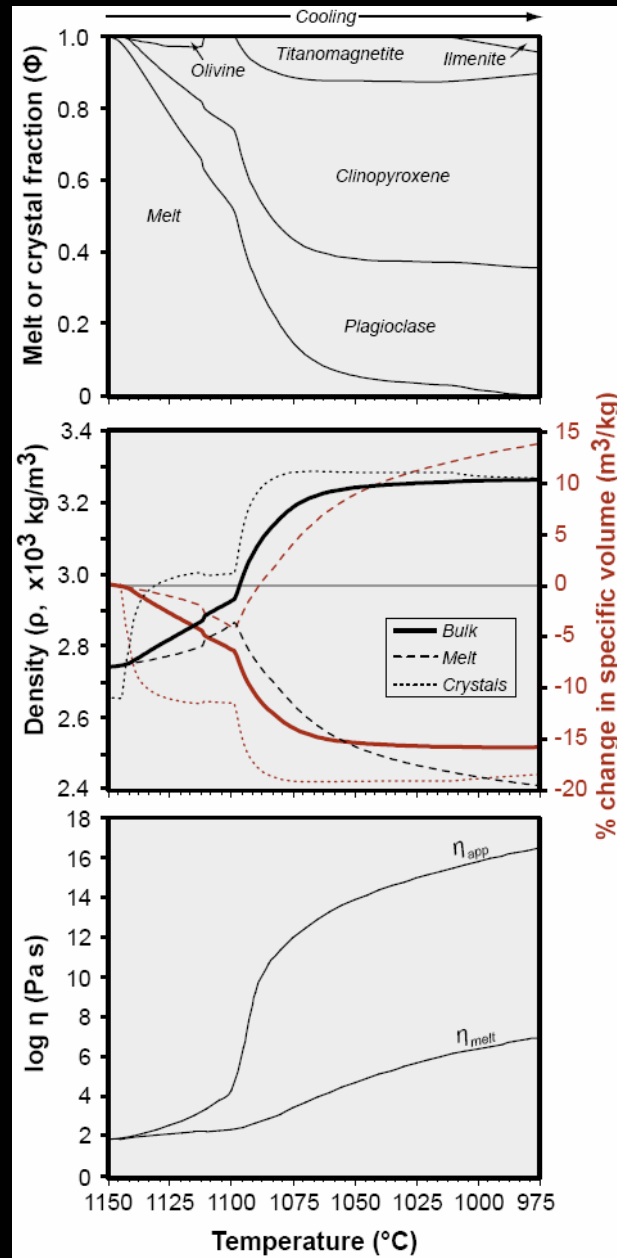
What happens inside a column?

Micro-scale observations across the column

- variation in composition
- variation of the crystal size distribution
- AMS sub-vertical, slight change in the susceptibility-ellipsoid orientation



Thermodynamic model

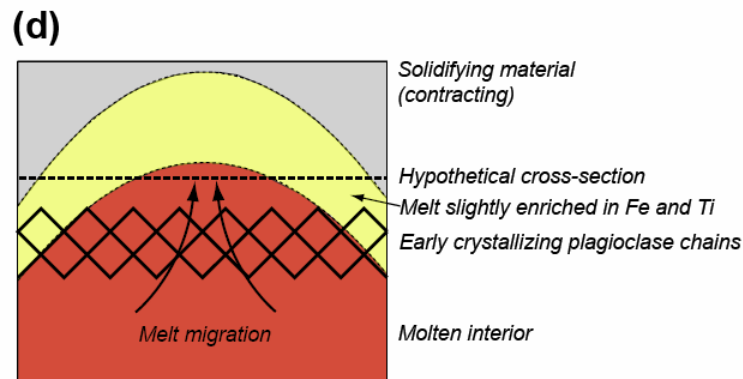
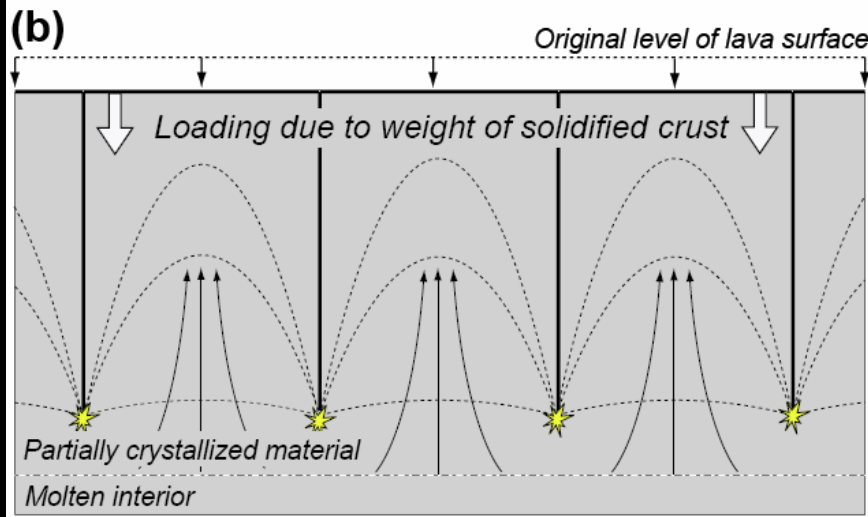
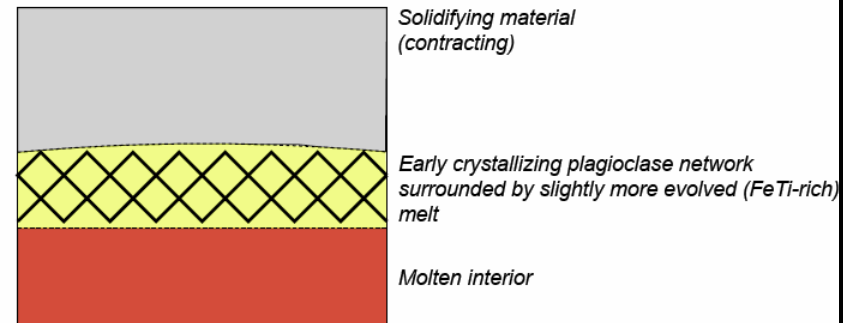
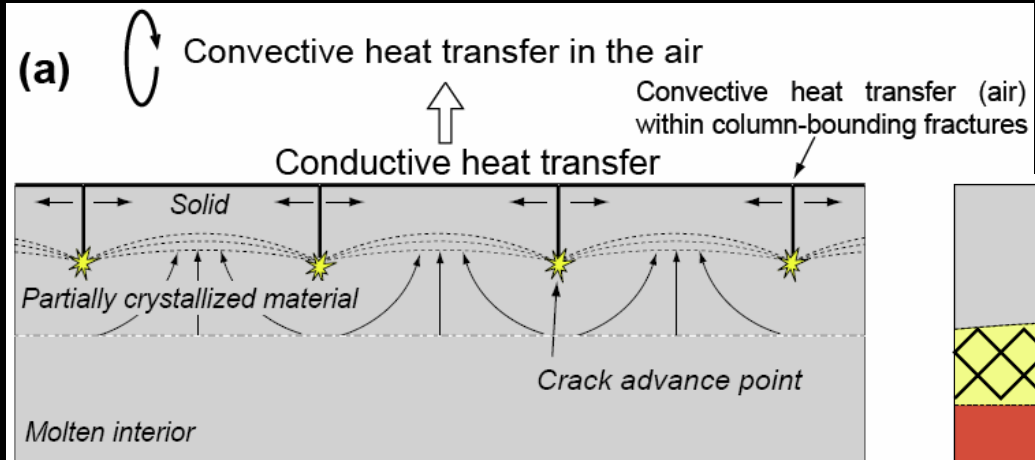


MELTS model *Ghiorso & Sack 1995, Asimow & Ghiorso 1998*

- crystallization order
- physical properties
 - volume decrease of ~15%
- joints can account for 0.5-1%
 - significant vertical component
- early crystallization of plagioclase
- late crystallizing titano-magnetite
- important change in physical properties below 1100°C

Model for melt migration within the column

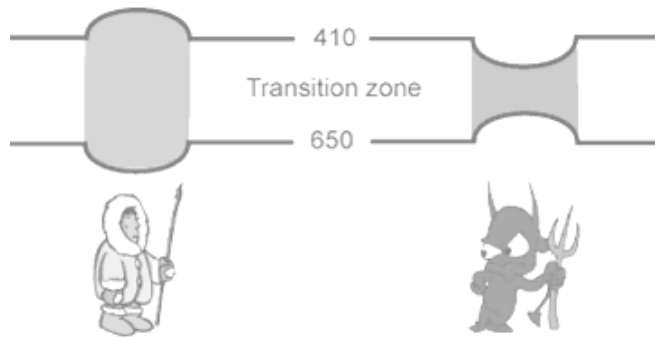
- early crystallization of plagioclase
- late crystallizing titano-magnetite
- important change in physical properties below 1100°C



« subsidence » by weight & passive upward melt migration into the column

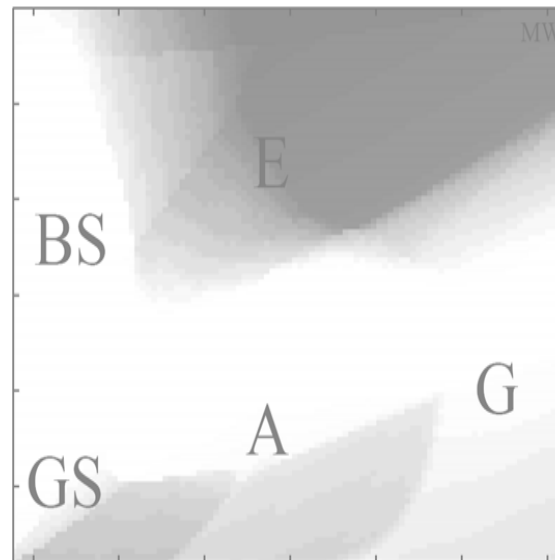
Triplet de changements de phases

zone de transition du manteau



COMPOSITION

modélisation géodynamique



VOLUME

prismation des orgues de lave



**COMPOSITION
& VOLUME**