

Inversion of travel times to estimate Moho depth in Shillong Plateau and Kinematic implications through stress analysis of Northeastern India

by

Saurabh Baruah

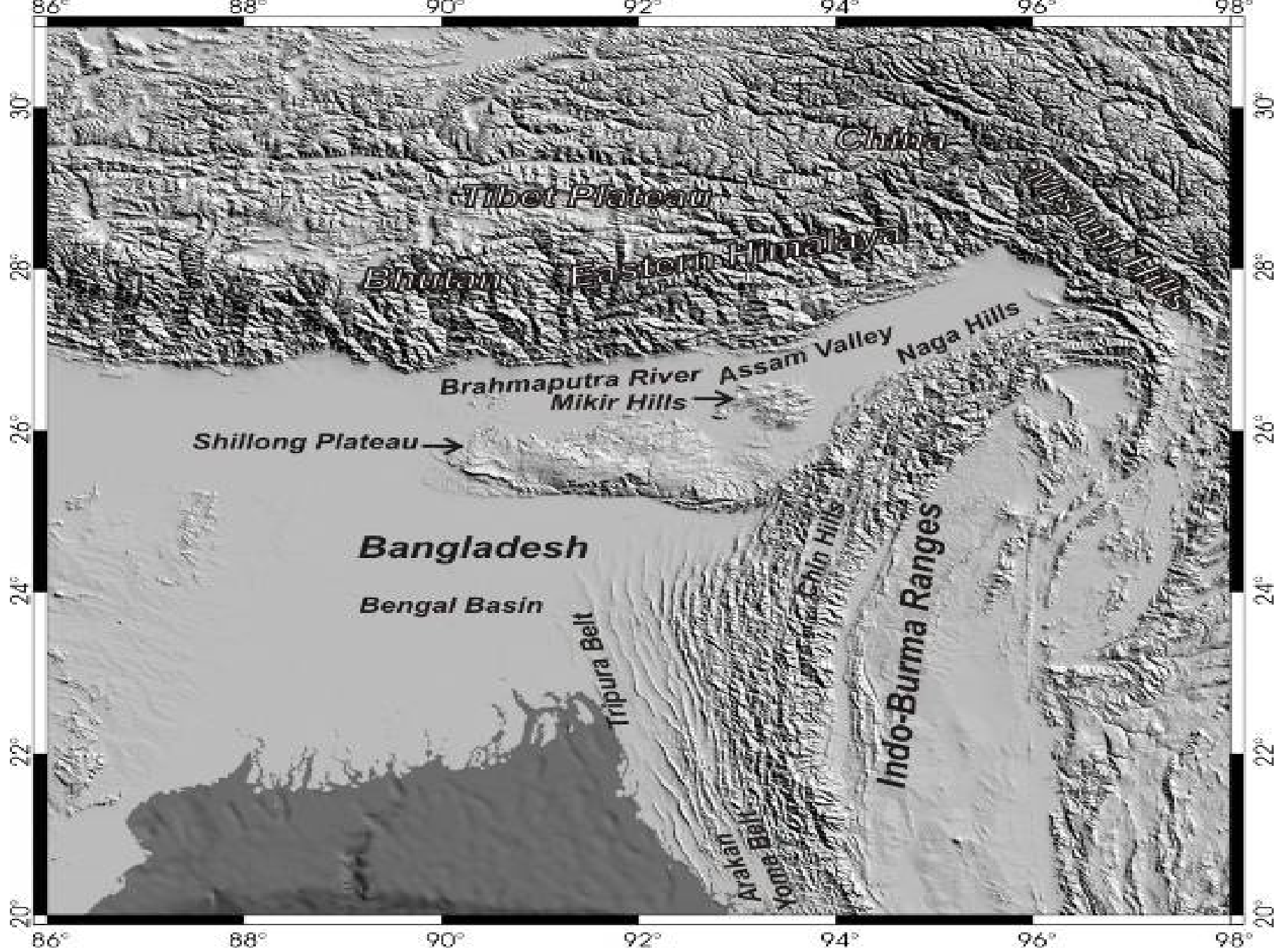
**Geoscience Division
North-East Institute of Science and Technology
(Council of Scientific & Industrial Research)
Jorhat-785 006, Assam, India
saurabhb_23@yahoo.com**

Plan

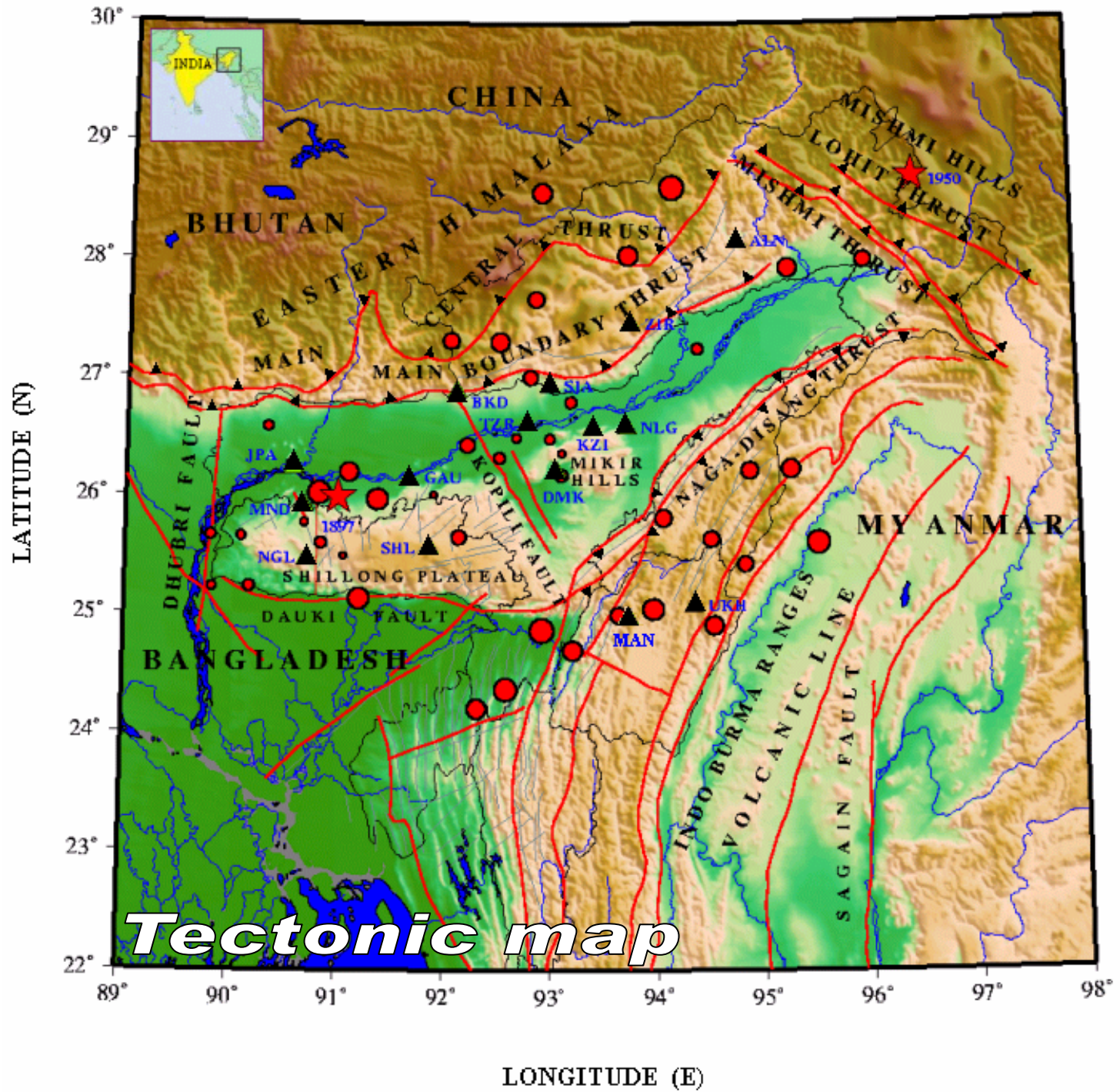
a. Tectonics and Seismicity

b. Inversion of travel times to estimate Moho

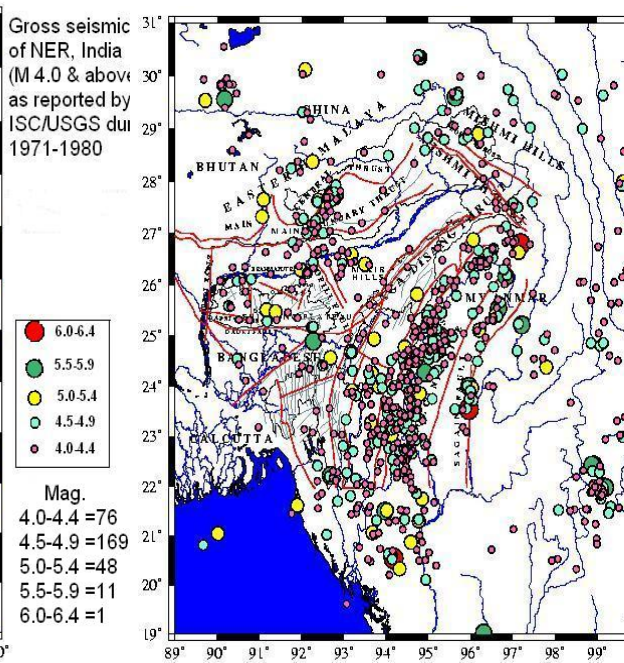
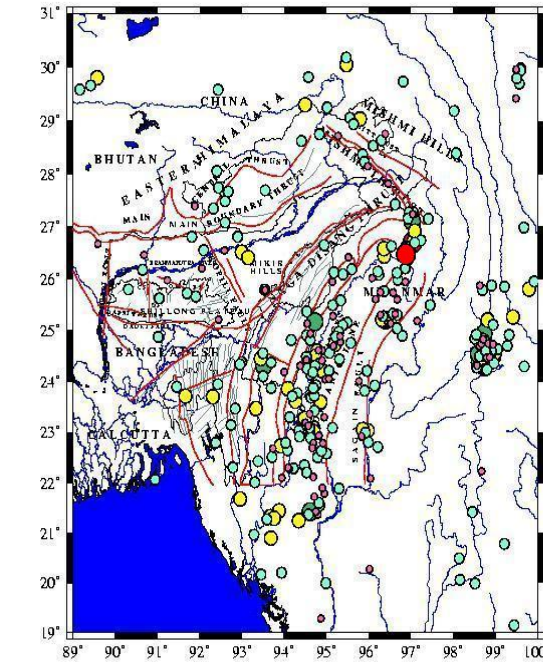
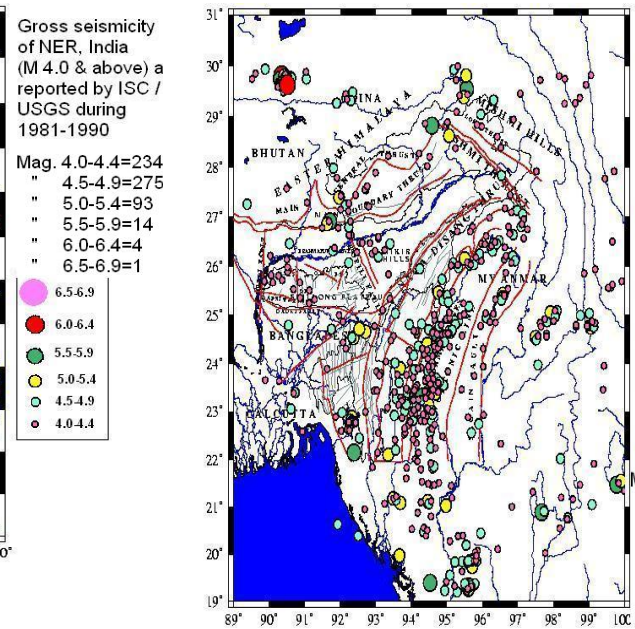
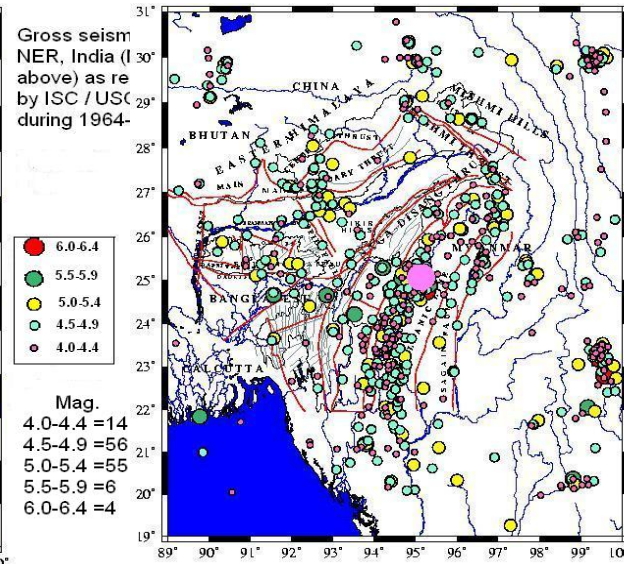
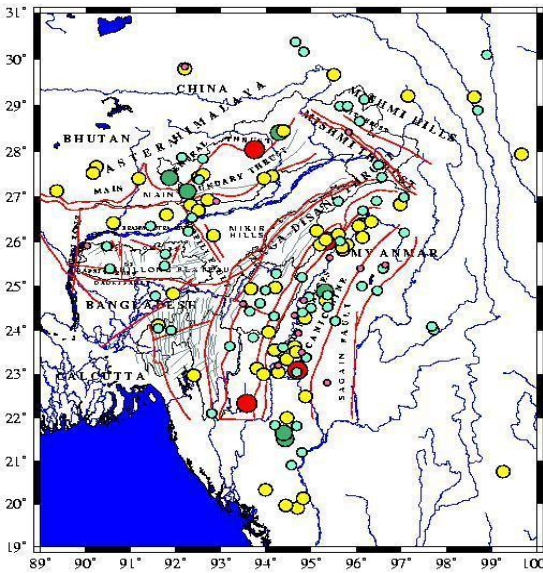
c. Stress analysis



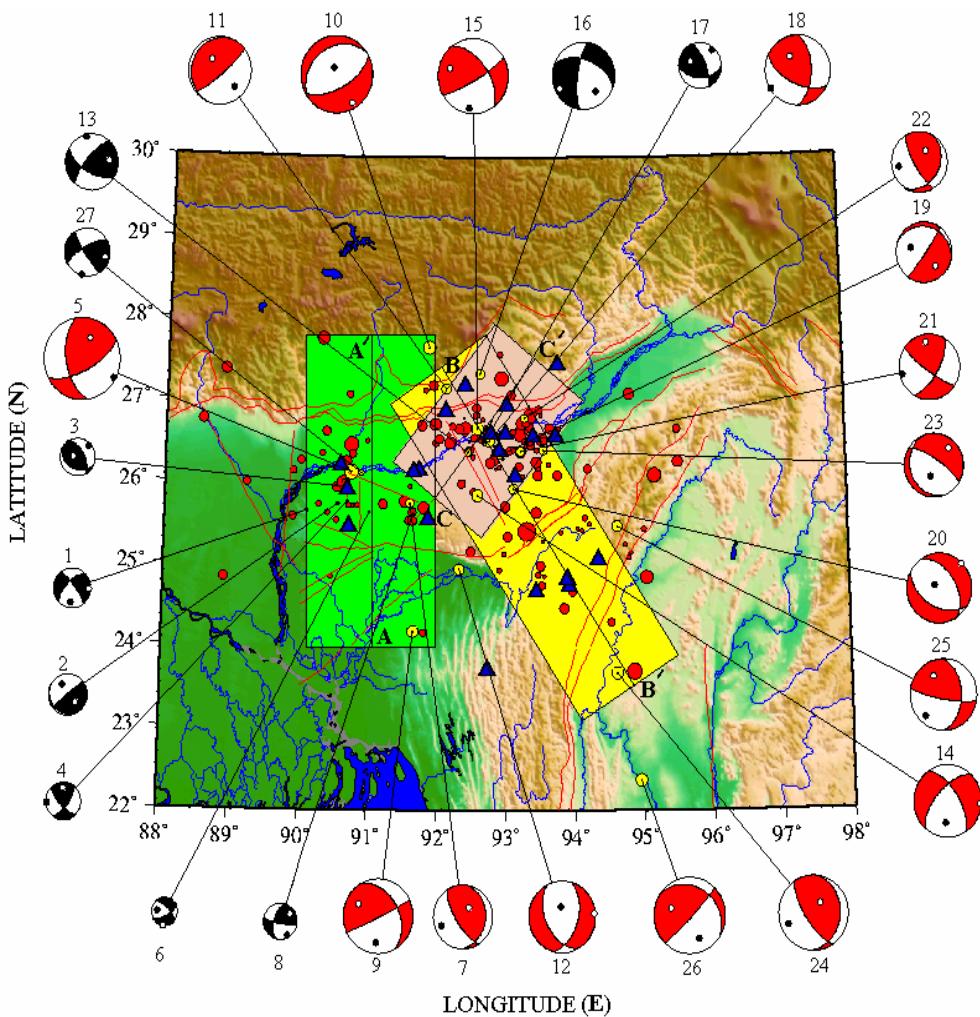
RELIEF MAP of NER, India



Epicentral map of earthquake of magnitude 4.0 and above occurred in NE India

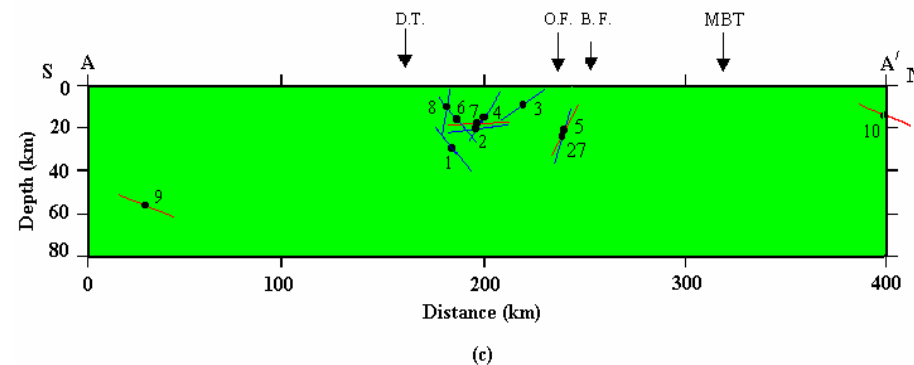
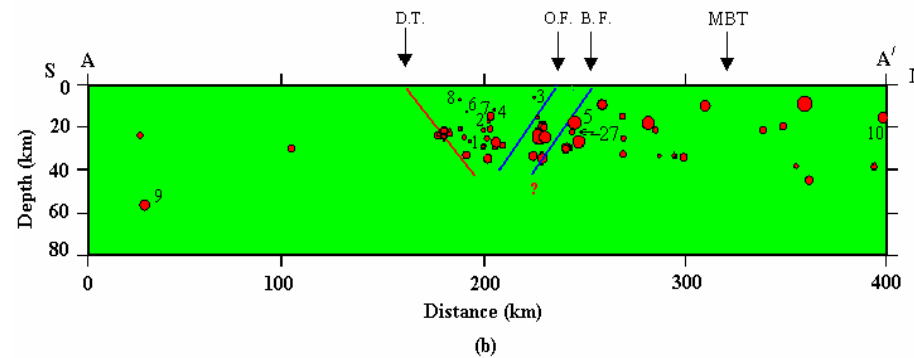
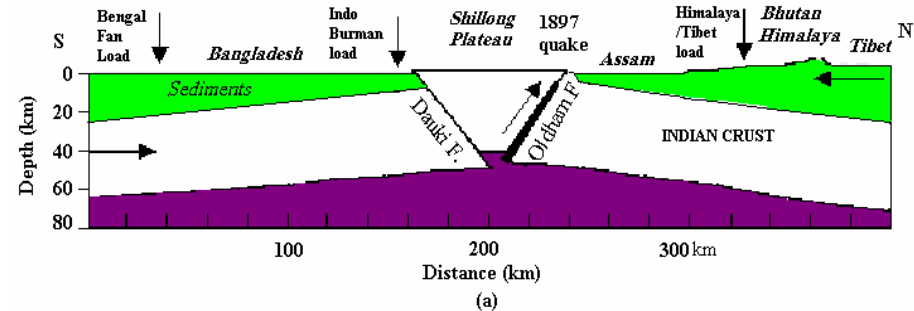


Since 1964-2008

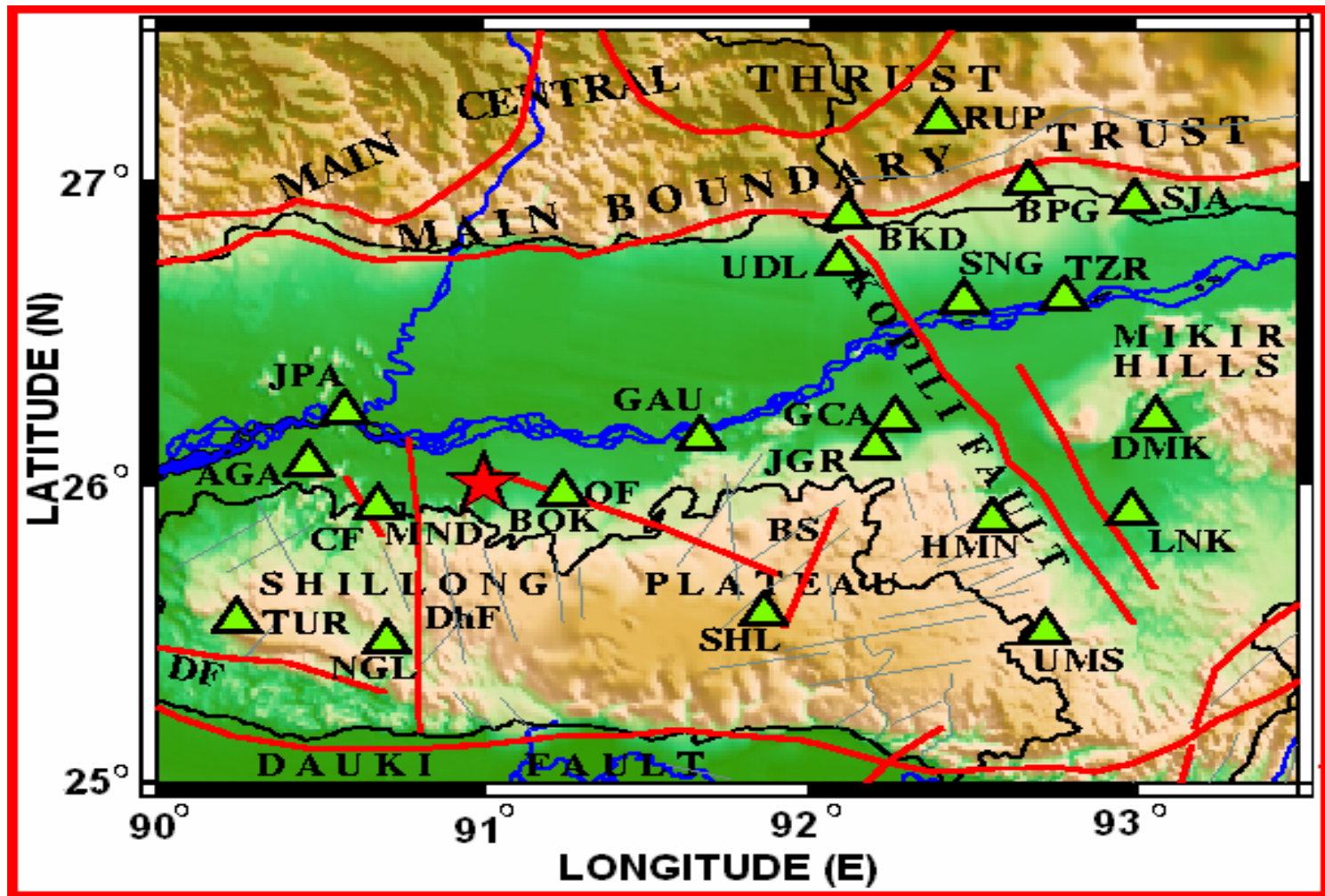


The map showing epicenters of 184 well located earthquakes; smaller circles indicate lower magnitude ($M < 3.0$) and bigger circles higher magnitude ($M > 3.0$) earthquakes. 26 fault-plane solutions shown, are obtained by wave form inversion(2001-2004)

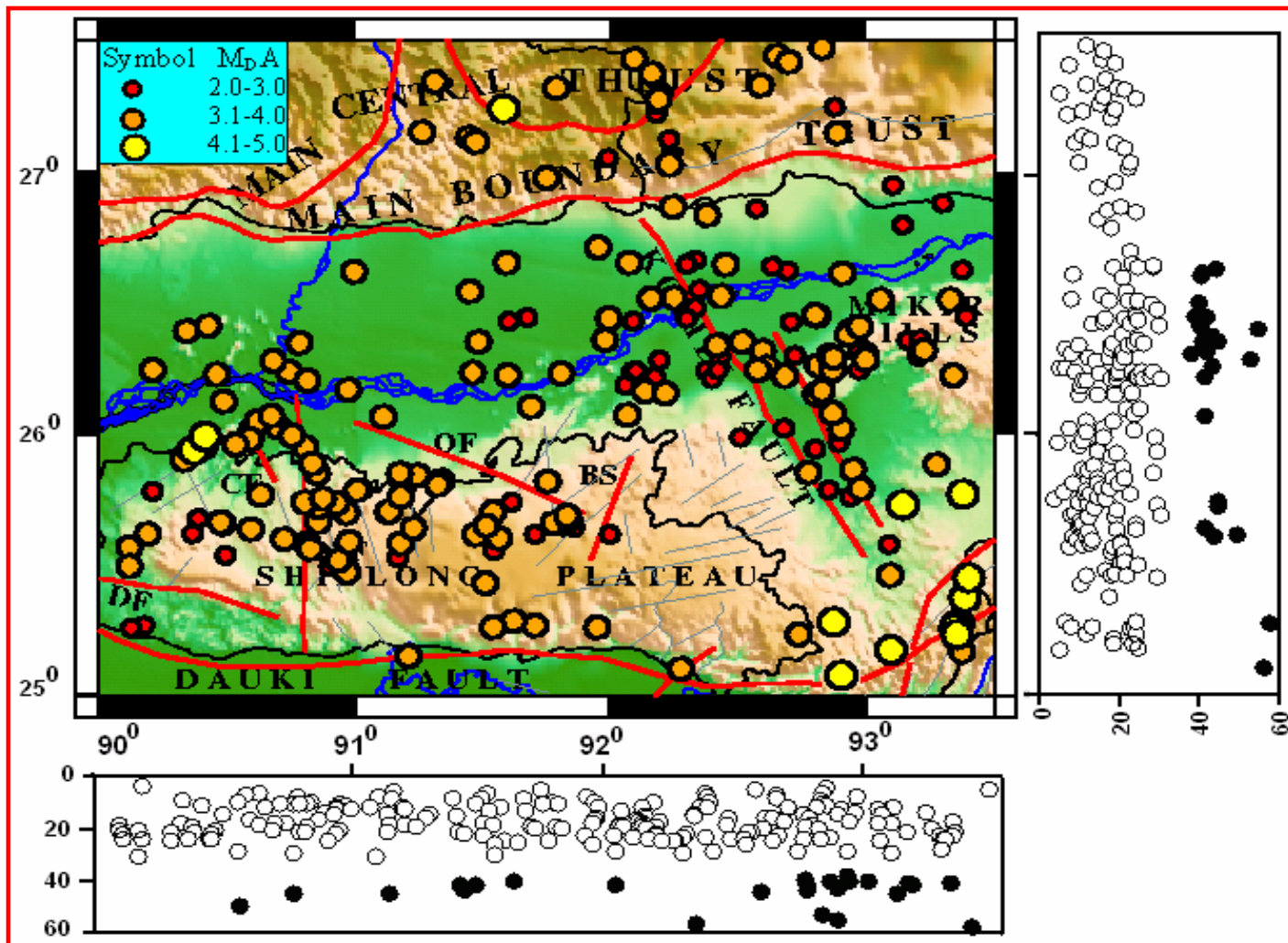
(a) The pop-up tectonic model of the Shillong Plateau [after *Bilham and England, 2001*], (b) cross section of the events across the Dapsi/Oldham/Brahmaputra fault zone, the considered events are shown by the shaded zone A-A' (Fig.3), and (c) cross section of the inferred fault planes of the selected events.



INVERSION OF TRAVEL TIMES TO ESTIMATE THE MOHO DEPTH



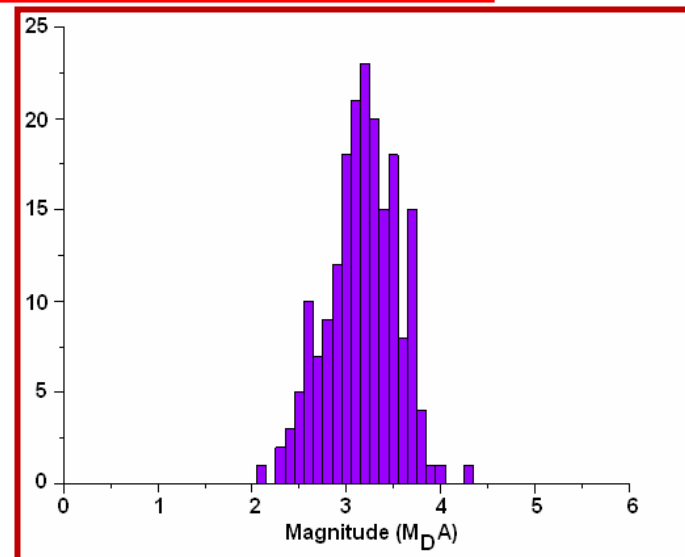
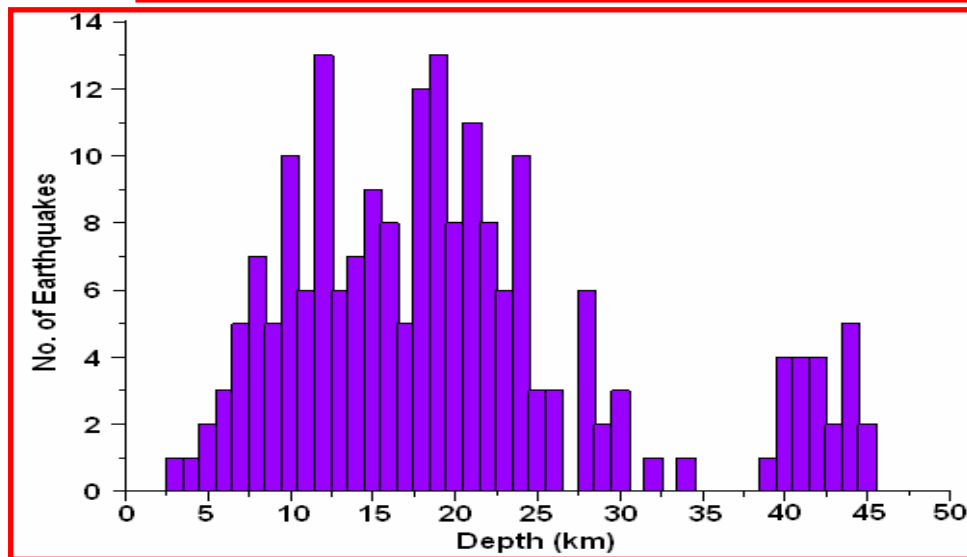
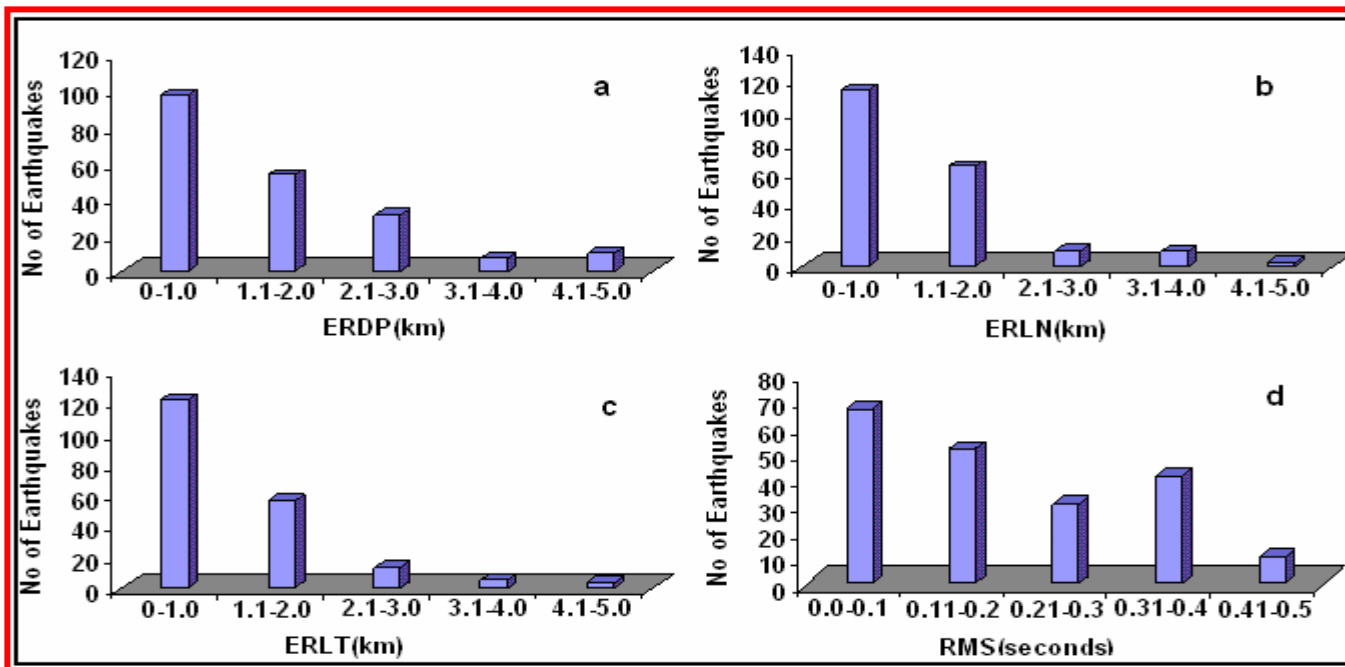
Map showing the major tectonic features of the study region. The epicenter of Great earthquake of 12 June, 1897(M=8.7) is shown by a star. Green triangles represent the digital broadband seismic station.



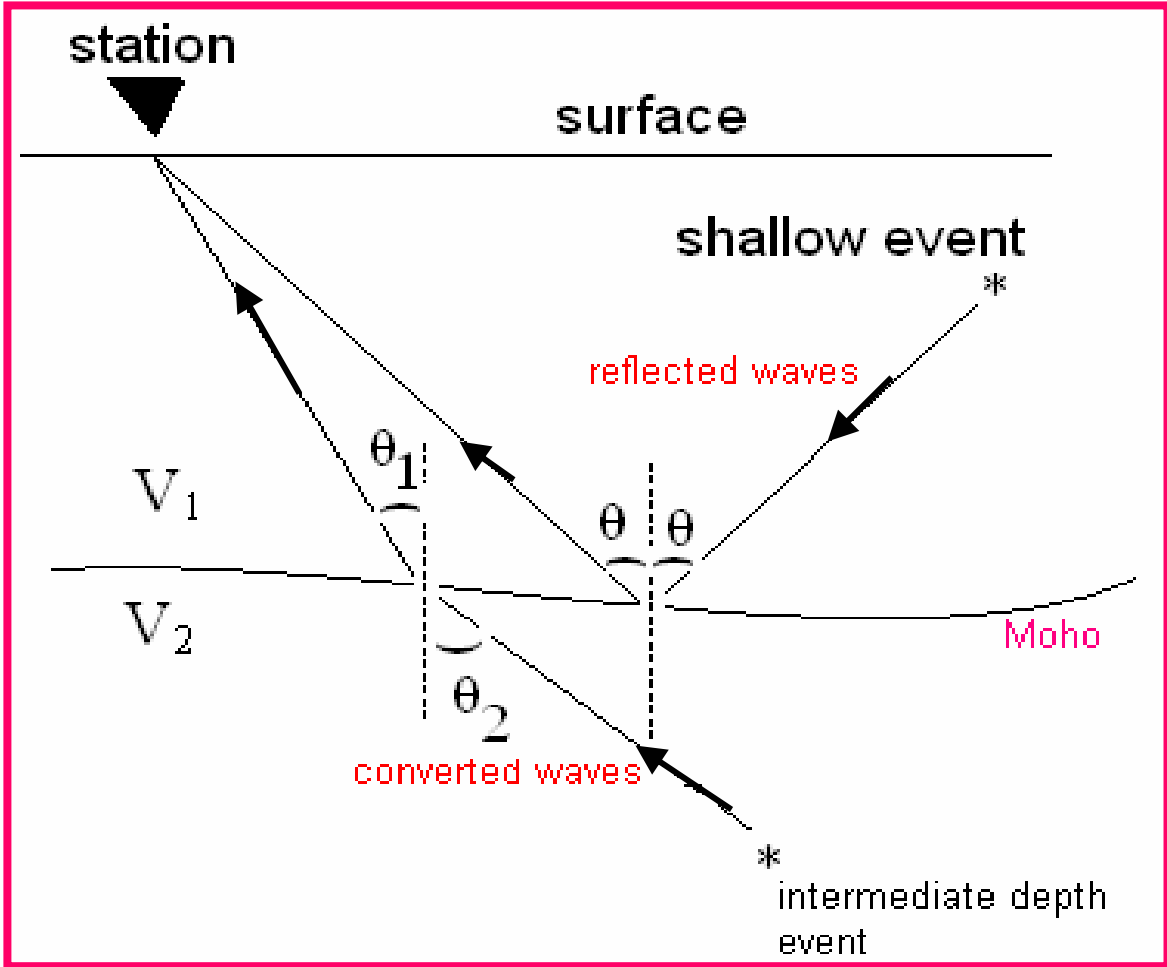
Hypocentral distributions of earthquakes along with depth sections. (a) Open and solid circles denote shallow and intermediate-depth earthquakes, respectively illustrated by different magnitude range. (b) Depth section plot along Longitude. (c) Depth section plot along Latitude.

Uncertainties involved in the estimates of epicenters:

203 earthquakes used ,966 reflected(PmP & SmS) and 70(PS & SP) converted phases

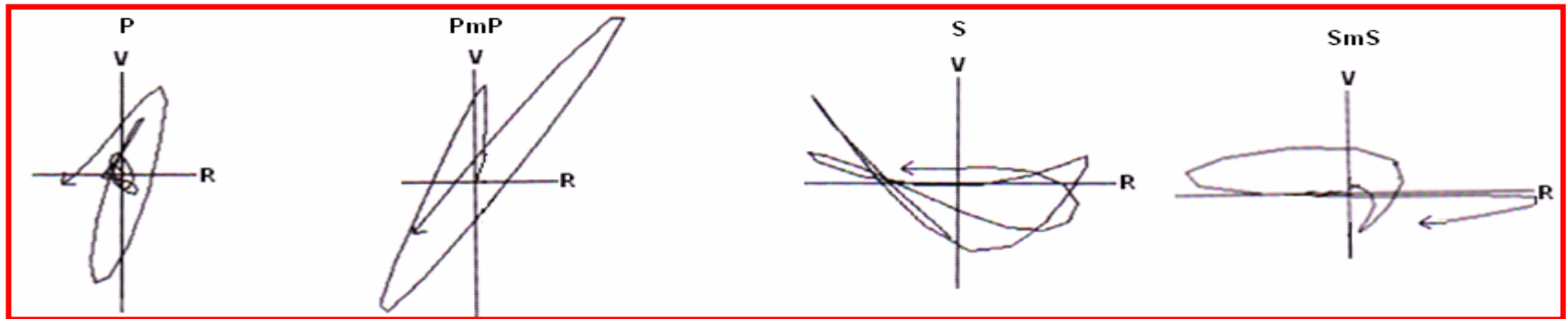
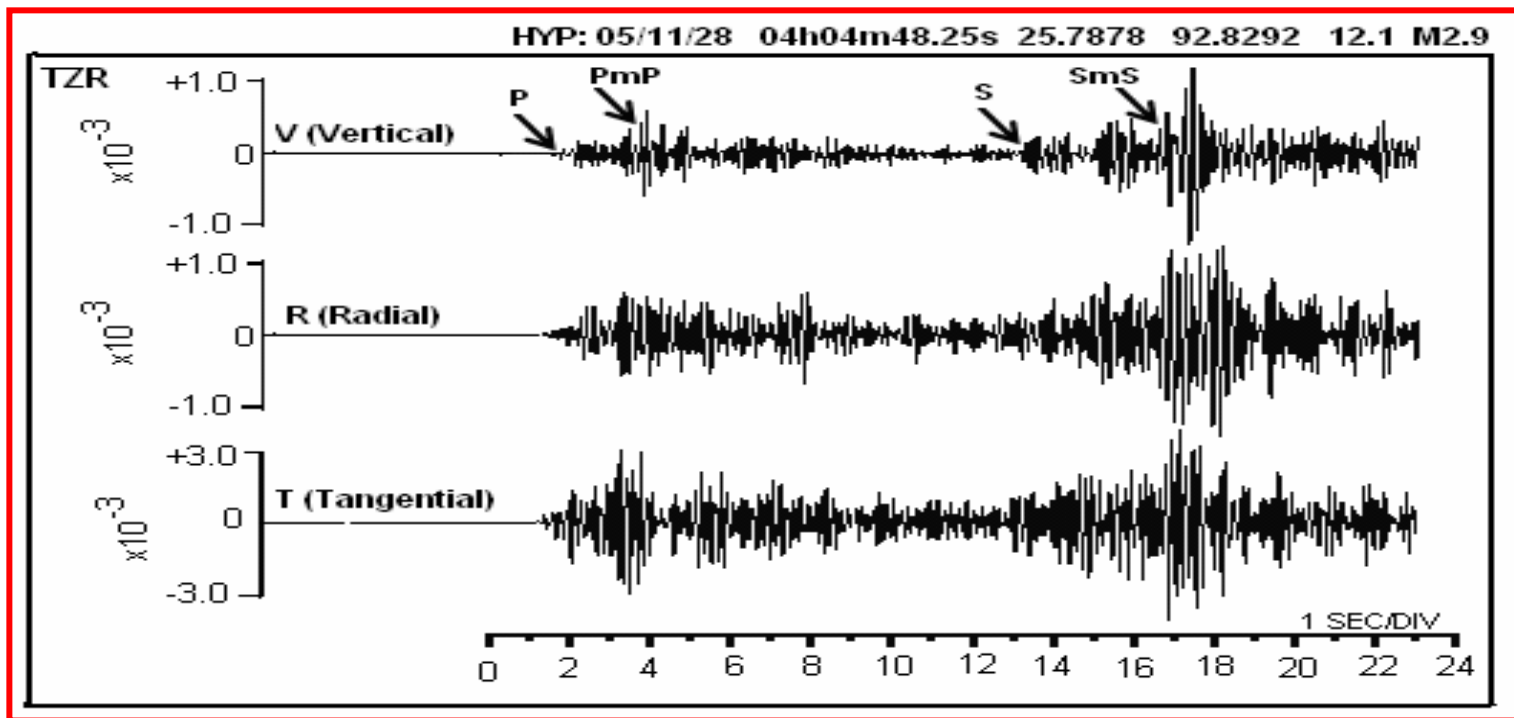


A schematic illustration of ray path of reflected and converted waves

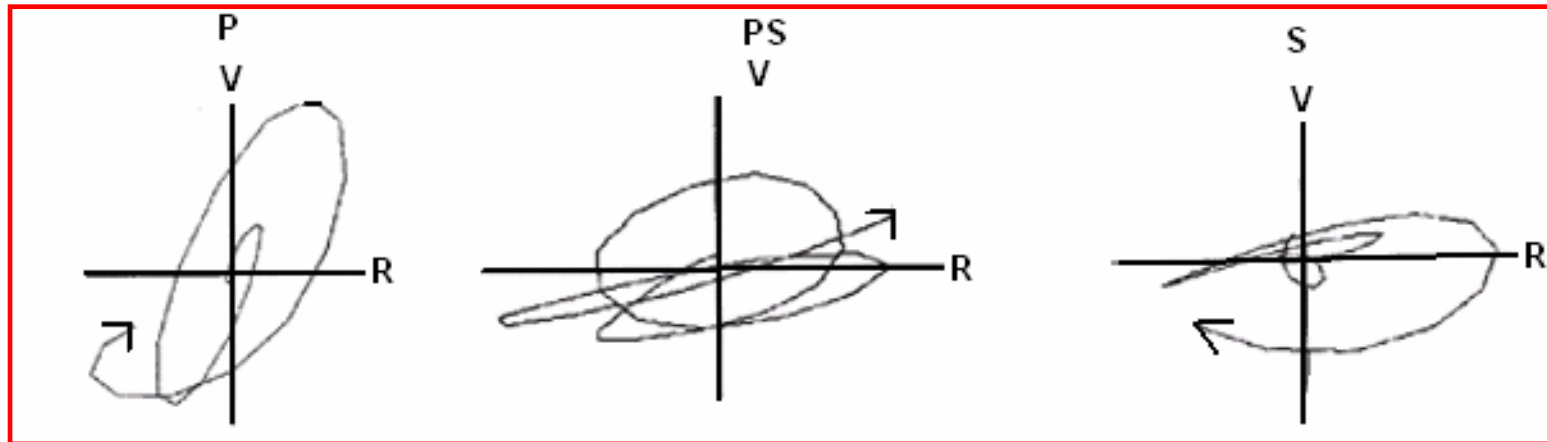
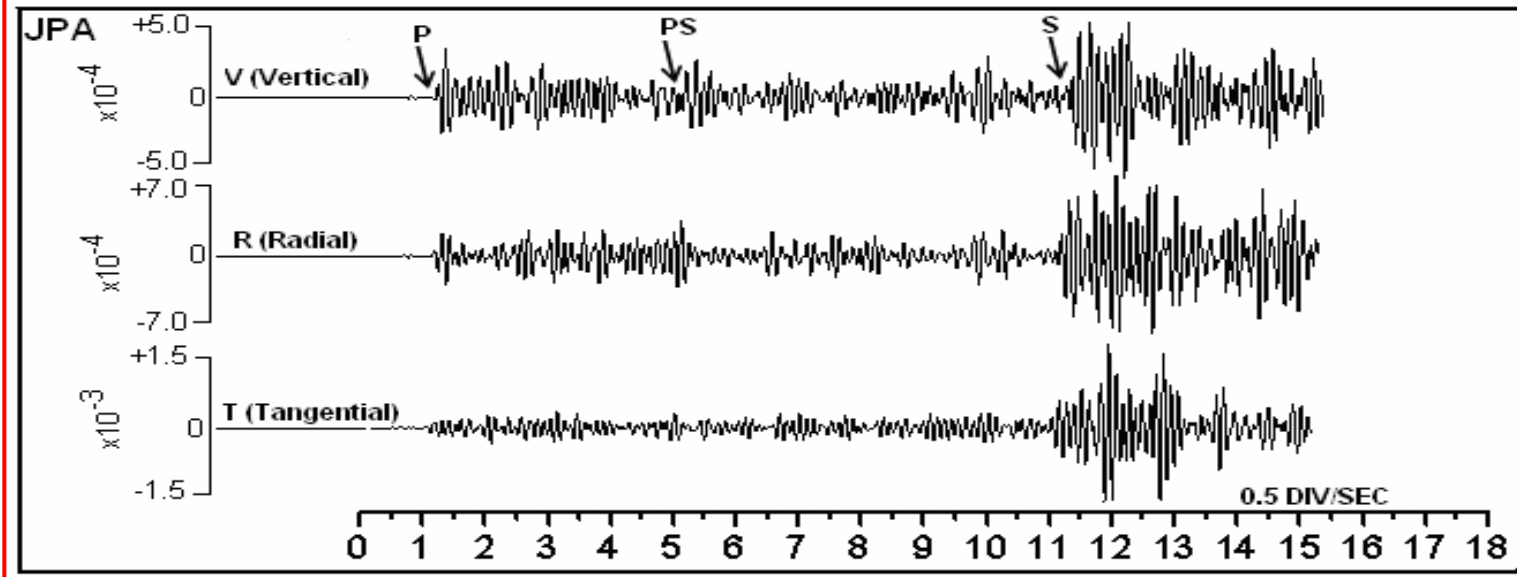


We consider two cases of ray propagation corresponding to the reflected wave and the converted wave at the Moho. One is the ray propagates in medium 1(velocity V_1) and is reflected at the interface with medium 2. The other is that the ray propagates from the medium 2(velocity V_2) into the medium 1 with a phase conversion at the interface.

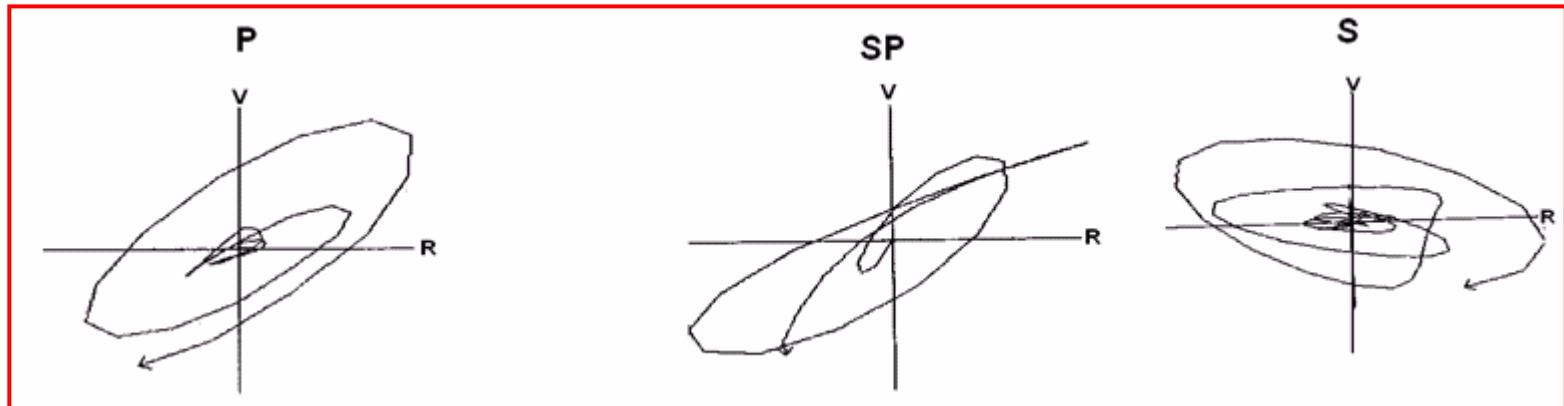
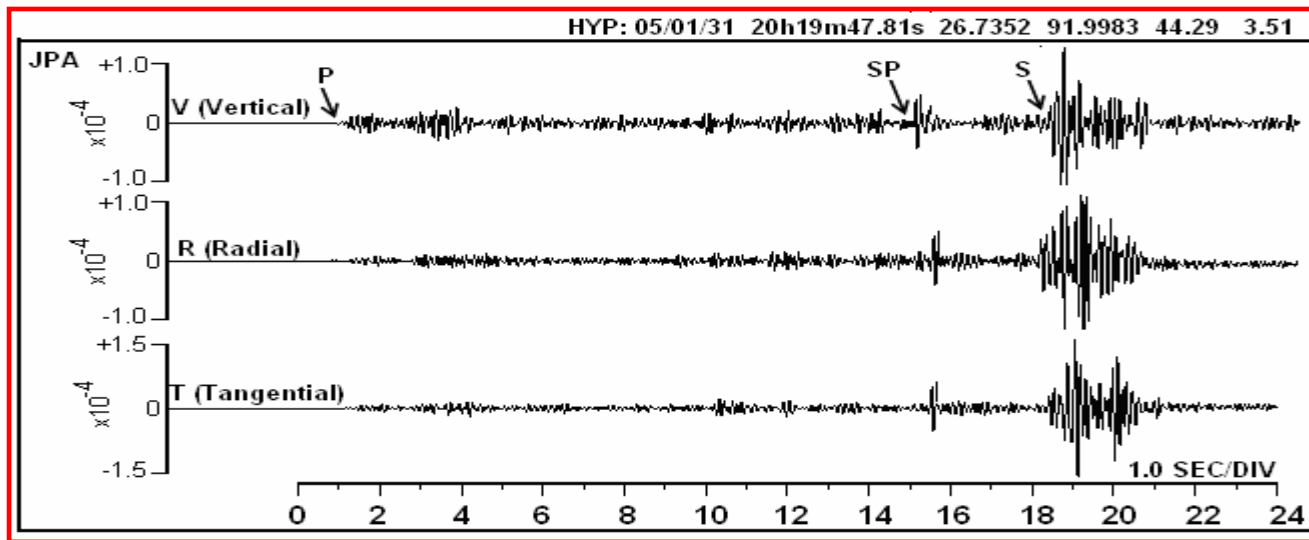
V_1 and V_2 denote the velocity in each medium. θ shows incident angle of reflected waves. θ_1 and θ_2 represent emergent and incident angles of converted waves, respectively.



An example of rotated three-component seismogram of a shallow earthquake recorded at station TEZPUR (TZR). Prominent later phases (PmP and SmS) can be seen after the first P- and S-wave arrivals, respectively. Particle motions of P, PmP, S and SmS phases are also shown.



An example of rotated three-component seismogram of an intermediate depth earthquake recorded at station JPA .A prominent later phases (PS) can be seen about 3.8 s after the P-wave arrival. Particle motions of P, PS and S phases are also shown.



An example of rotated three-component seismogram of an intermediate depth earthquake recorded at (JPA). A prominent later phase (SP) can be seen about 3.1 s before the S- wave arrival. Particle motions of P, SP and S phases are also shown.

Theoretical consideration

The reflected, refracted phases and converted phases (PS and SP) at Moho are observed in seismograms for local earthquakes. Travel times of these phases are inverted to estimate depth of the Moho discontinuity.

Depth distribution of the Moho are expressed as a function of latitude and longitude. The Moho depth (H_m) at a location (Φ', λ') is expressed as :

$$H_m(\Phi', \lambda') = C_0 + C_1 \Phi' + C_2 \lambda' + C_3 \Phi'^2 + C_4 \Phi' \lambda' + \text{-----} + C_{14} \lambda'^4 \quad \text{-----}(1)$$

where Φ' and λ' are the latitude and longitude respectively. C_k 's are unknown parameters which may be determined by inversion of the observed travel time data.

Travel time residuals can be written as :

$$T_{p2-p1}^{obs} - T_{p2-p1}^{cal} = \sum_k \left(\frac{\partial T_{p2}}{\partial C_k} - \frac{\partial T_{p1}}{\partial C_k} \right) \Delta C_k + e \quad \text{----- (2)}$$

Where P_1 and P_2 denote the first and the later phases, respectively. For example, P_1 and P_2 are P and PmP for the PmP-P data. ΔC_k is a correction term for unknown parameters and e is an error.

Partial derivatives of T are expressed as:

$$\frac{\partial T}{\partial C_k} = \frac{\partial T}{\partial H_m} \frac{\partial H_m}{\partial C_k} \quad \text{----- (3)}$$

Partial derivatives of H_m against C_k is calculated from equation (1)

Change in Travel times for reflected and converted

The travel time change for the reflected wave due to change in depth of the interface as expressed by

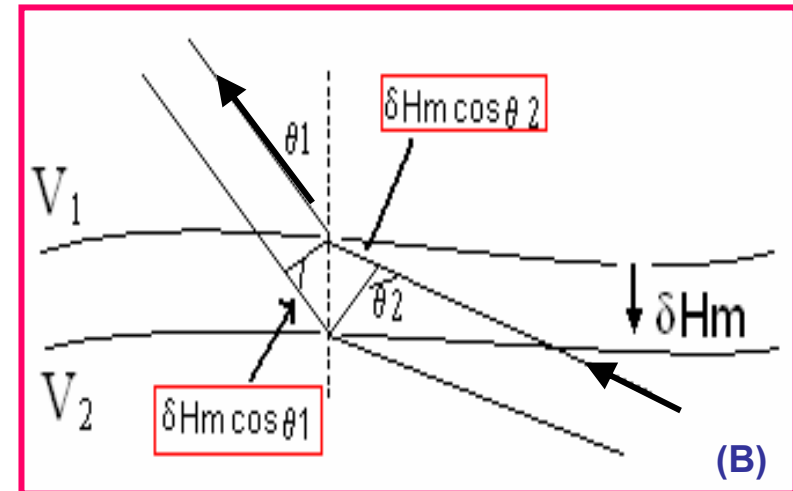
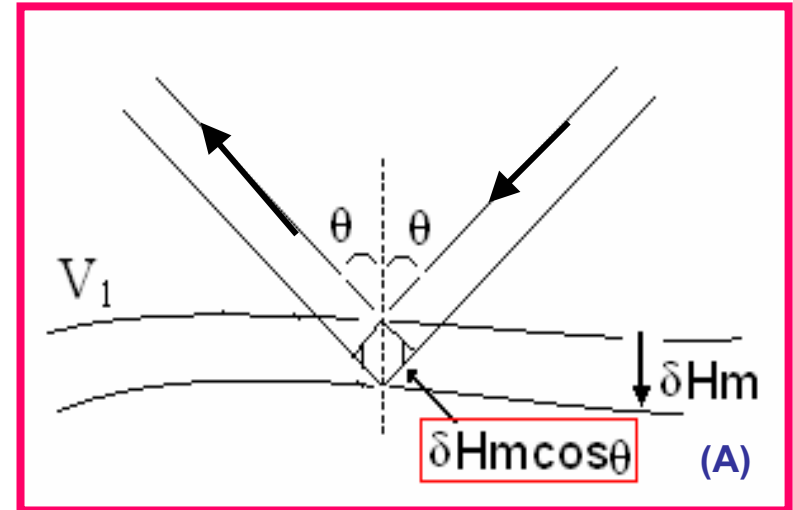
$$\partial T = \frac{2\partial H_m \cos\theta}{V_1} \quad \text{----- (4)}$$

and we obtain

$$\frac{\partial T}{\partial H_m} = \frac{2\cos\theta}{V_1} \quad \text{----- (5)}$$

Similarly, we obtain the partial derivative $\partial T/\partial H_m$ for the converted wave as

$$\frac{\partial T}{\partial H_m} = \frac{\cos\theta_1}{V_1} - \frac{\cos\theta_2}{V_2} \quad \text{----- (6)}$$



(A) Travel time change of reflected waves due to the depth change of the Moho (B) Travel time change of converted waves due to the depth of the Moho.

Eq. (2) can be expressed by matrix form as

$$d = G m + e$$

----- (7)

where

d=the column vectors of residuals between observed and theoretical travel time difference

G=matrix of partial derivatives

m= column vector of correction term of unknown parameters ΔC_K

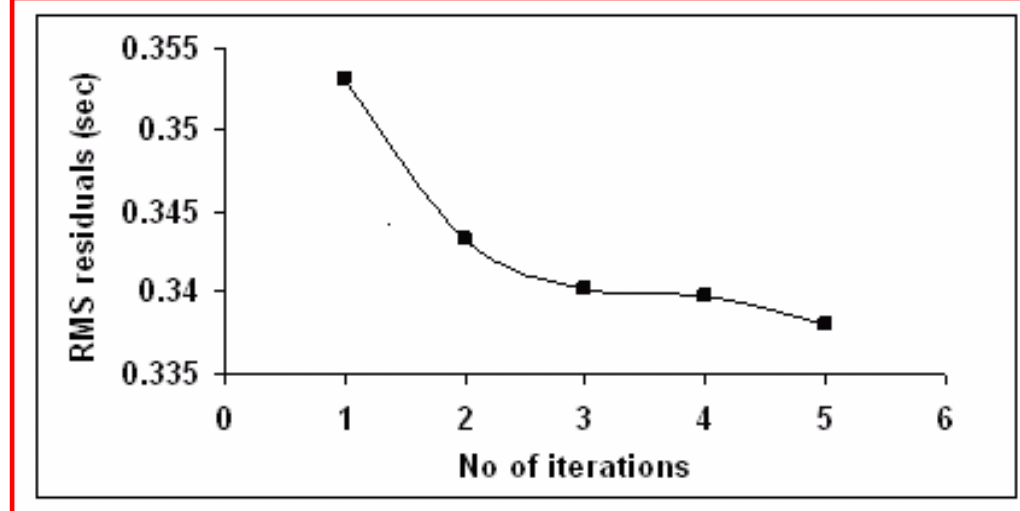
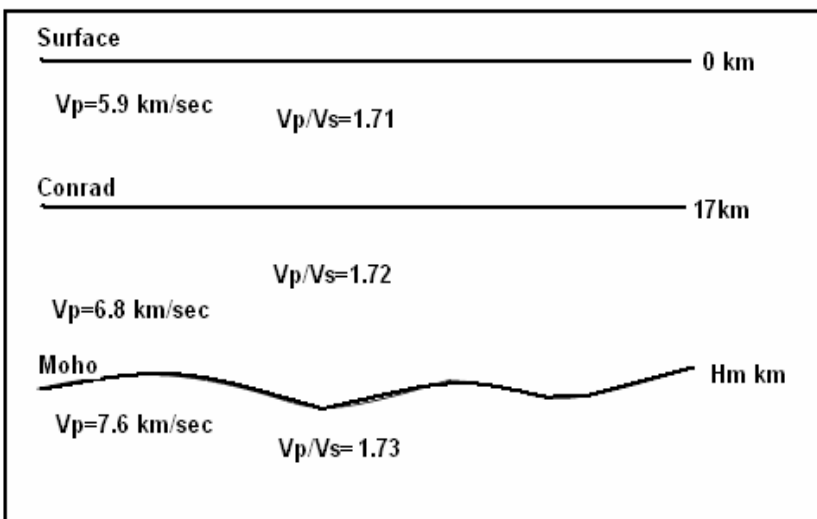
e=error vector

This equation can be written as follows by solving with least squares method:

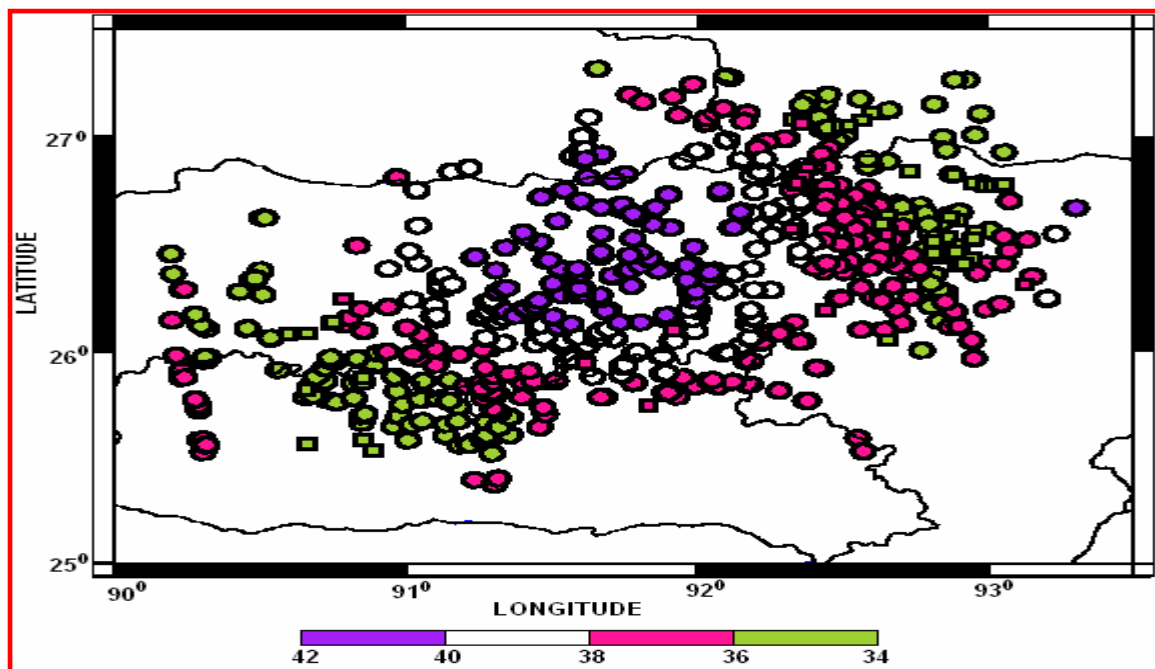
$$m = (G^T G)^{-1} G^T d$$

----- (8)

This calculation is carried out iteratively until ΔC_K becomes sufficiently small.



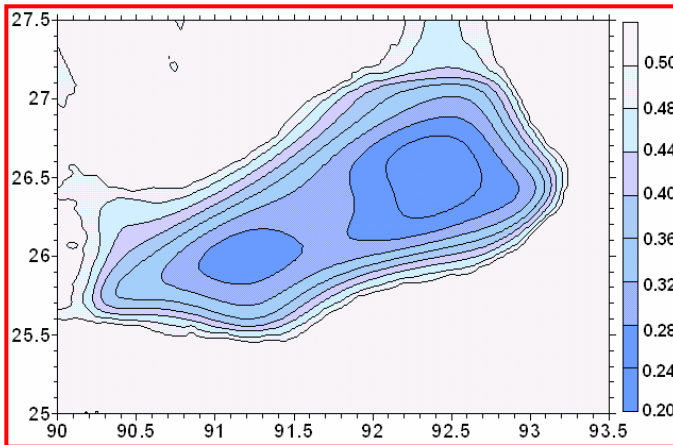
Root Mean Square (RMS) residuals versus number of iterations



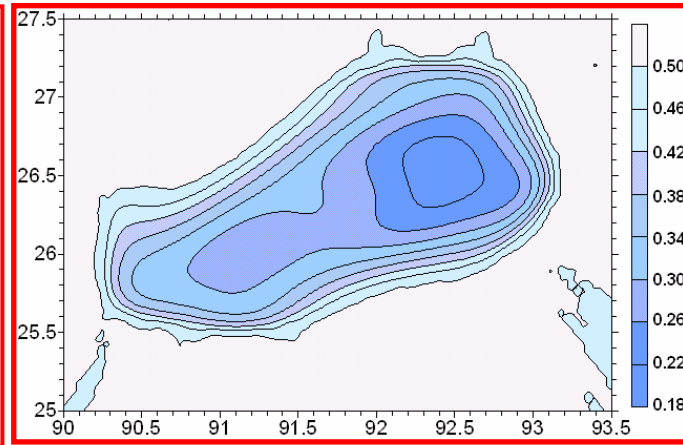
Moho depth obtained for each data point. Circles and squares denote reflection and conversion points, respectively.

Results

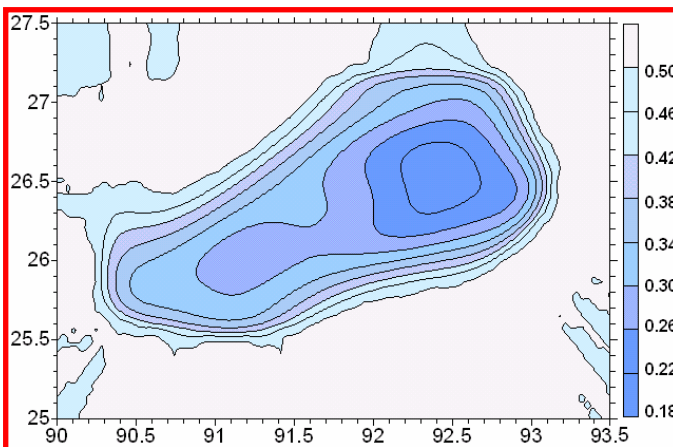
Iteration 1



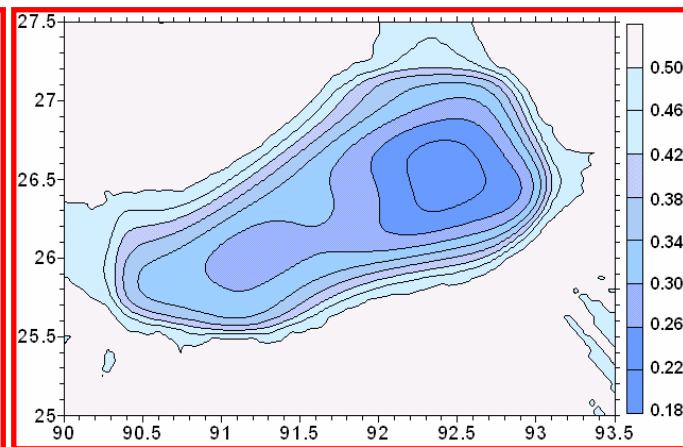
Iteration 2



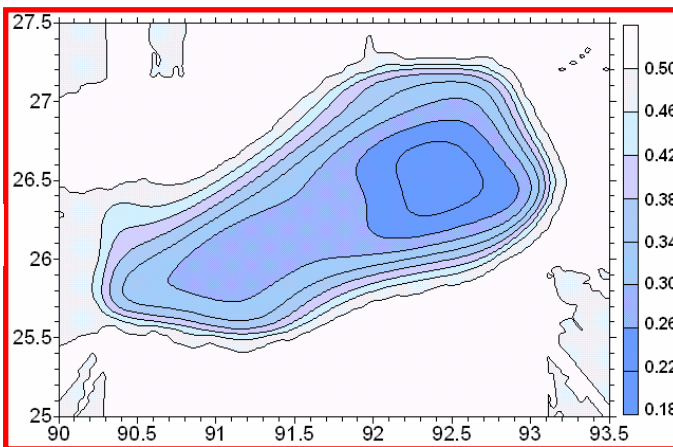
Iteration 3



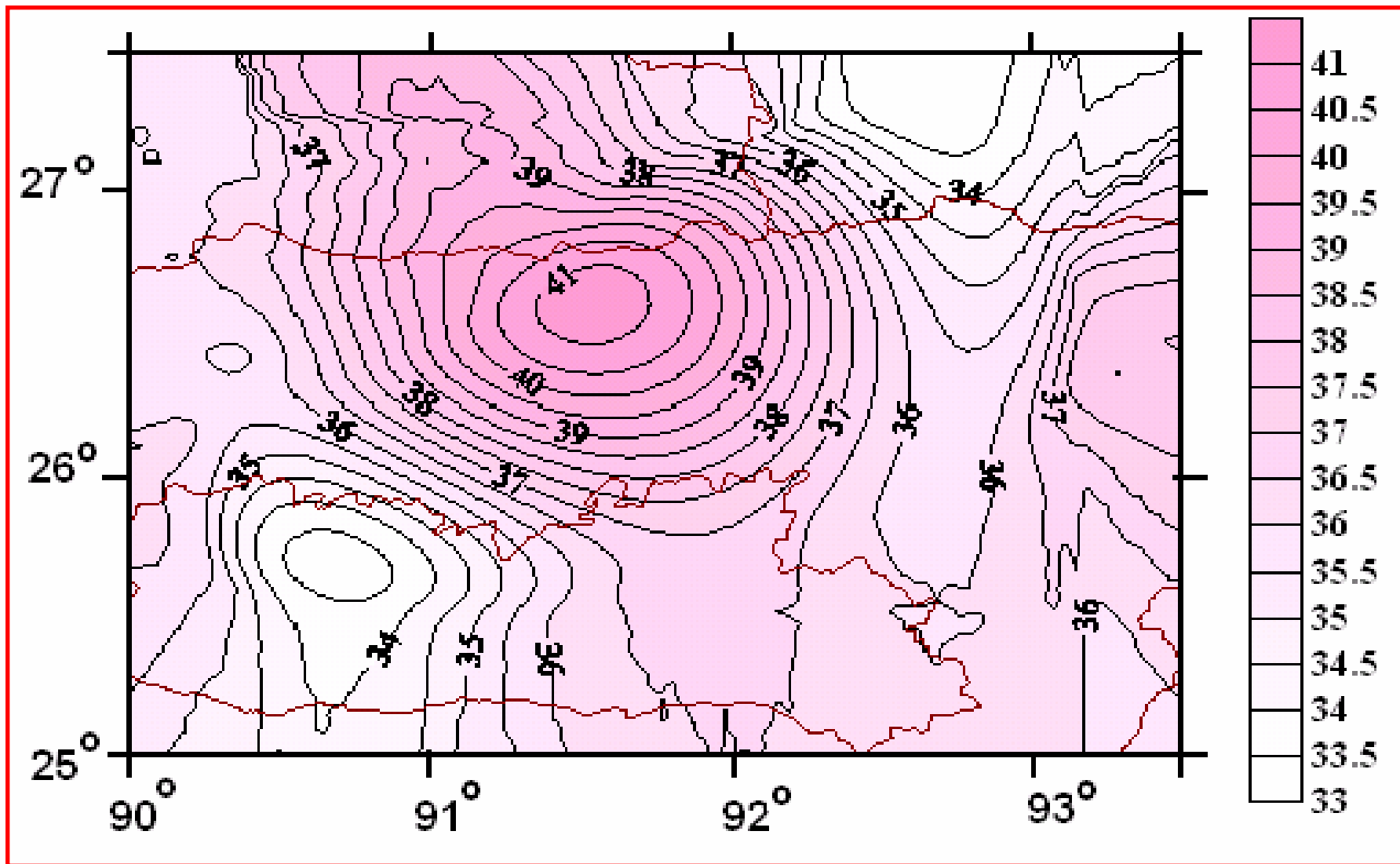
Iteration 4



Iteration 5



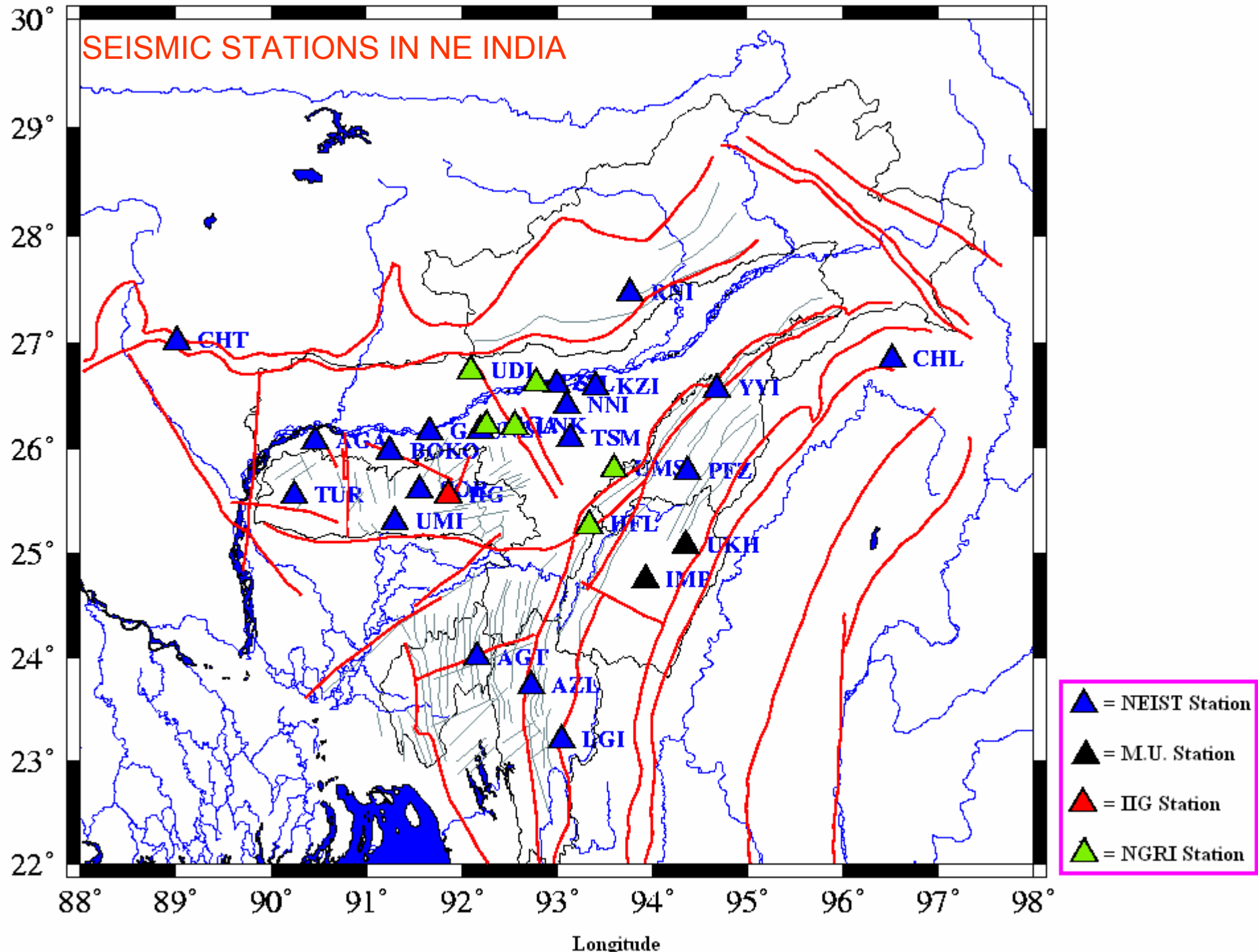
Contour diagram represents the RMS residuals of each of the iteration.

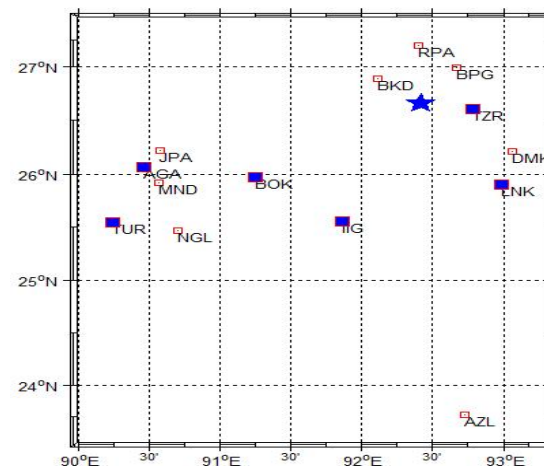
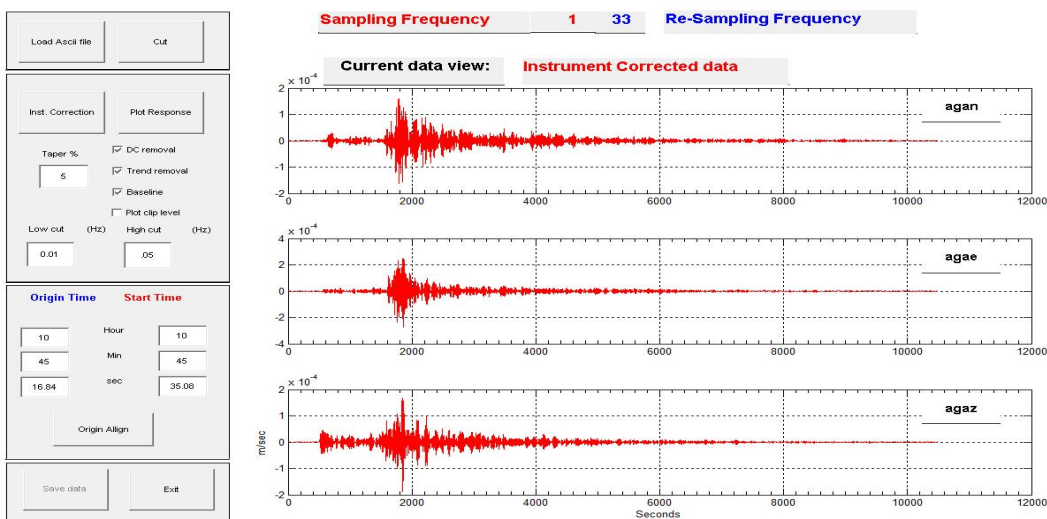


Depth distribution of Moho obtained at the reflection and conversion points

STRESS ANALYSIS IN NORTHEASTERN INDIA AND ITS KINEMATICS IMPLICATIONS

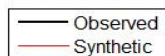
SEISMIC STATIONS IN NE INDIA





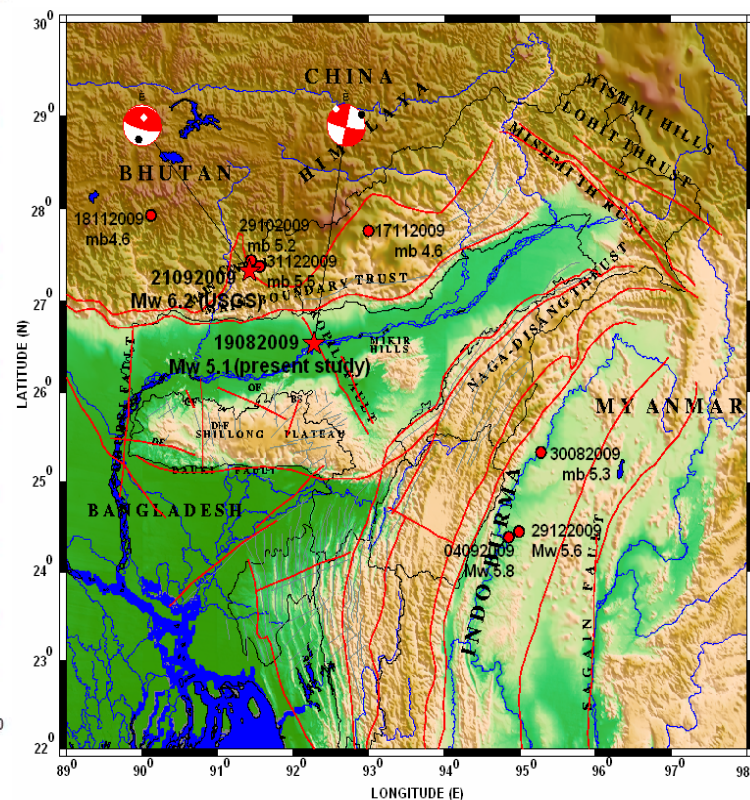
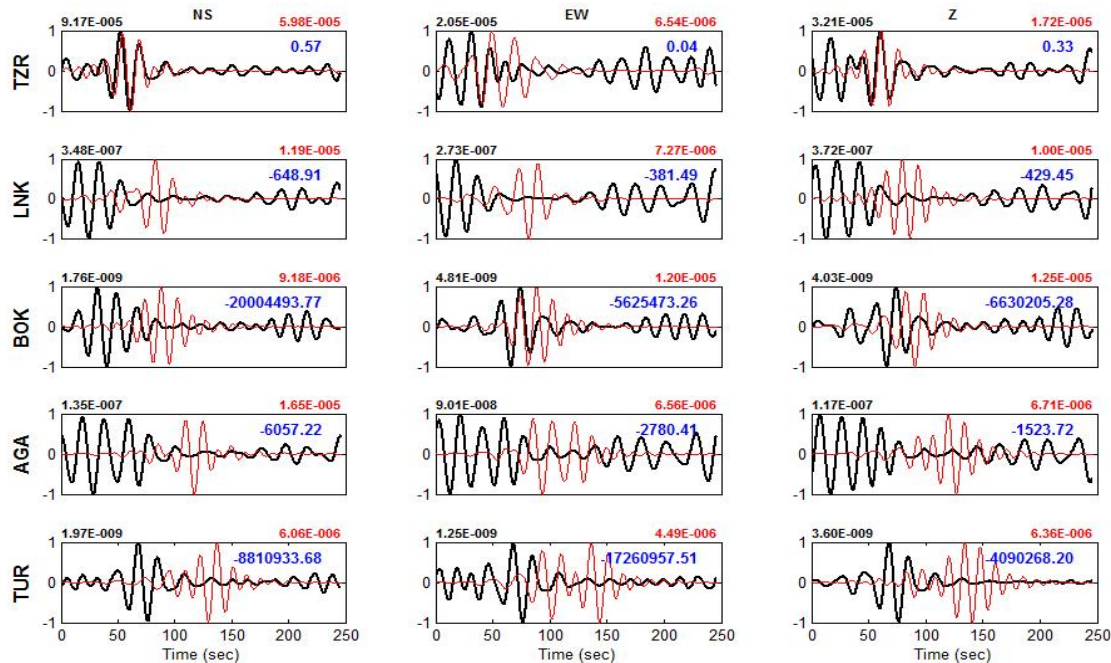
WAVEFORM INVERSION OF 19 AUGUST 2009 EVENT

Even date-time: 20090819 10:45:16 Displacement (m). Inversion band (Hz) 0.04 0.05 0.08 0.09

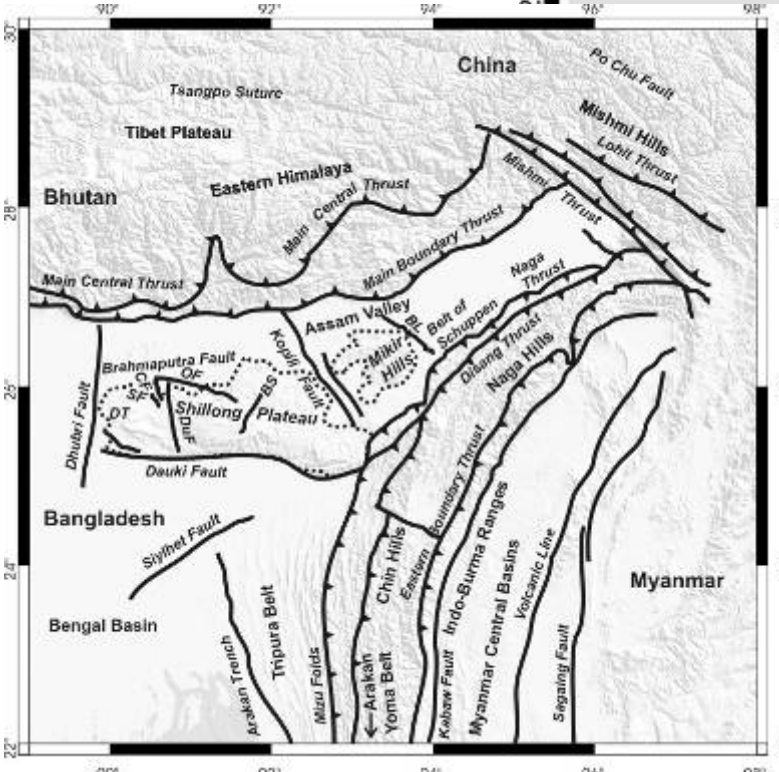
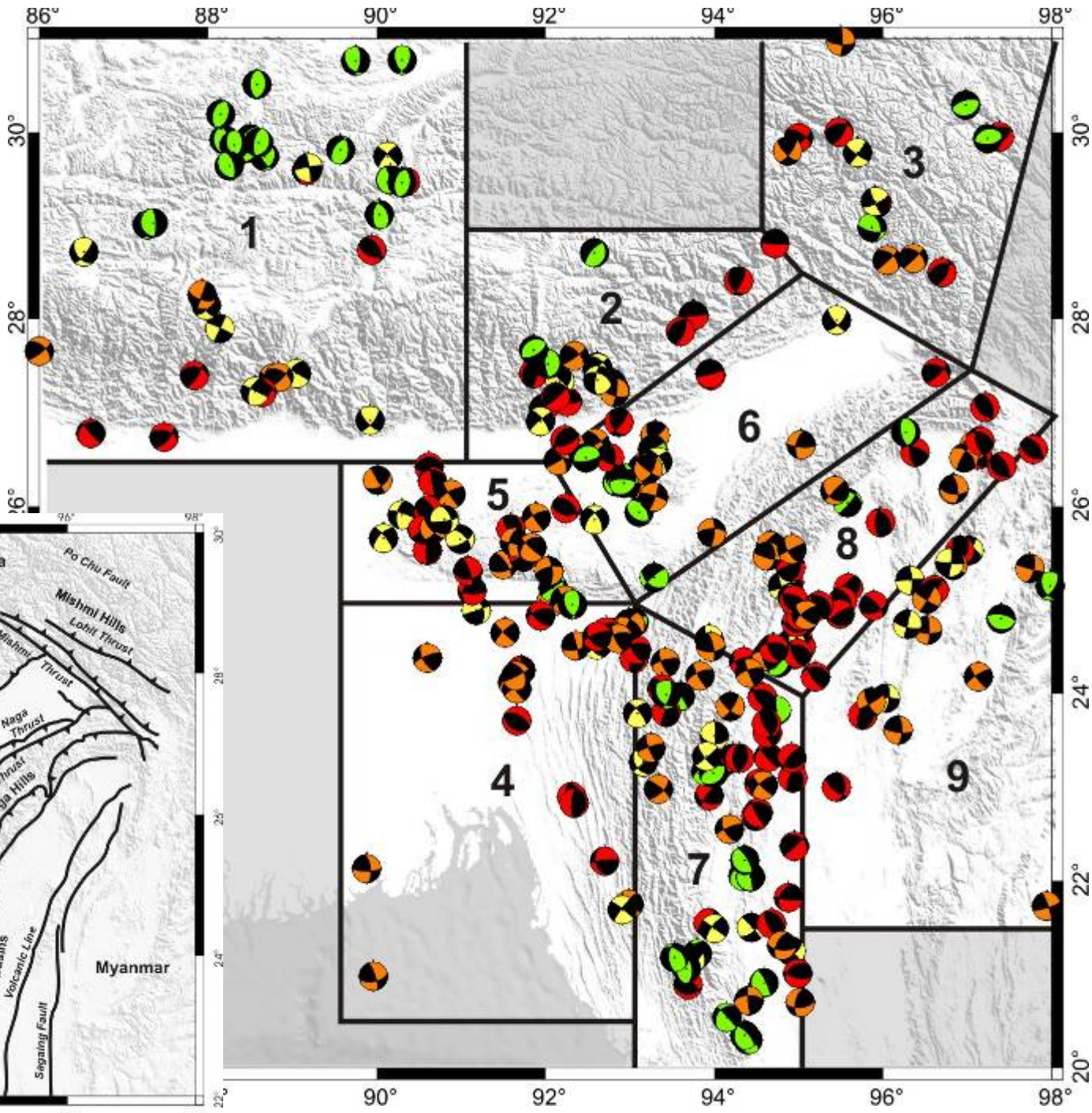


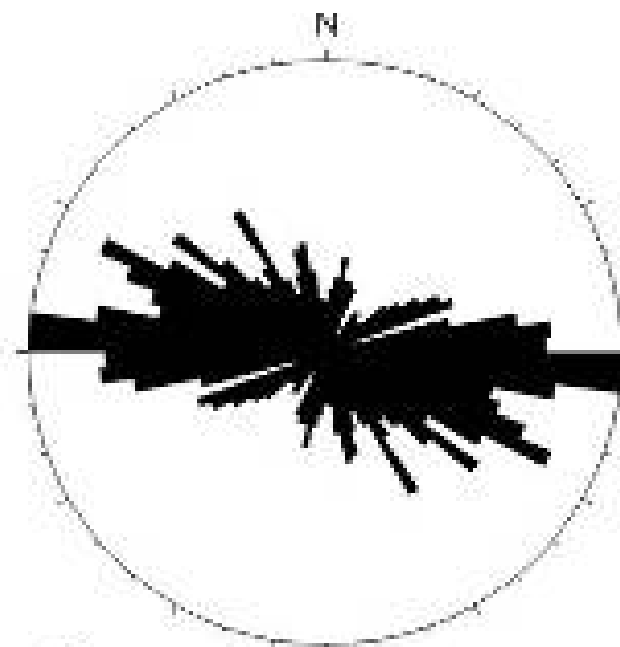
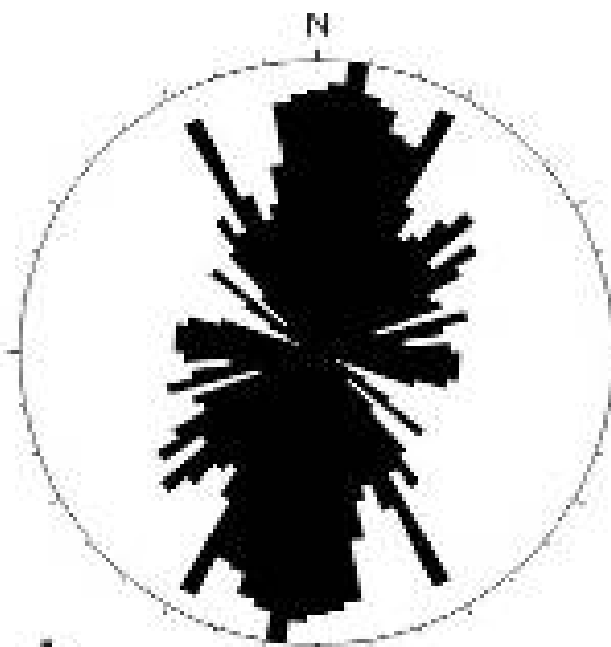
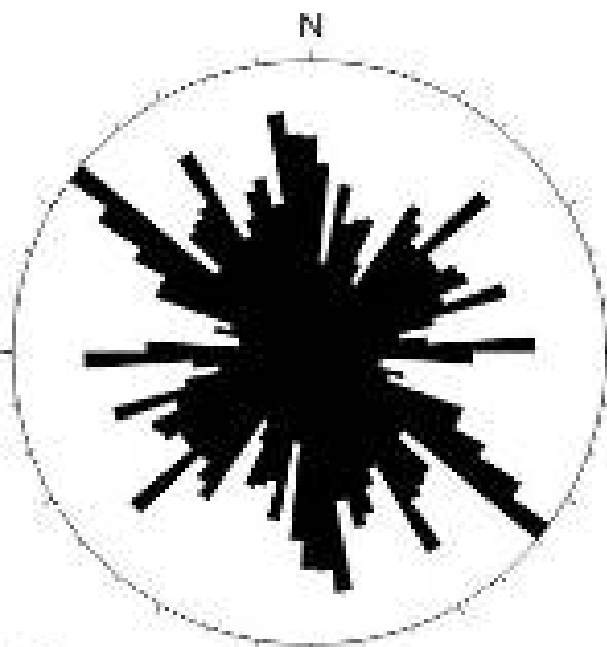
Gray waveforms weren't used in inversion.

Blue numbers are variance reduction



MAP DISTRIBUTION OF EPICENTERS OF 285 EARTHQUAKES CONSIDERED IN THIS STUDY

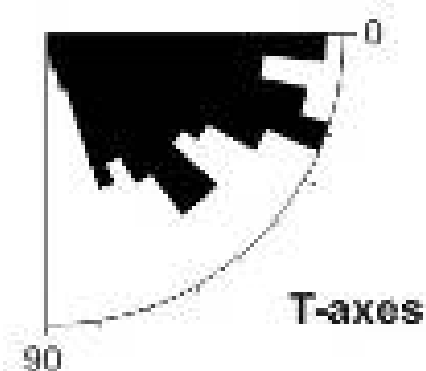
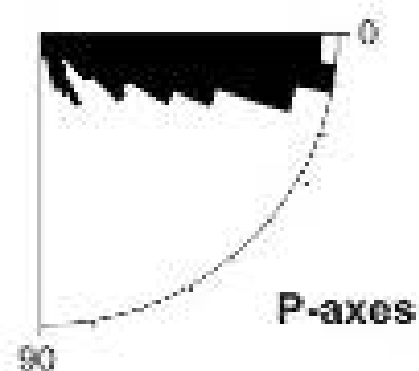
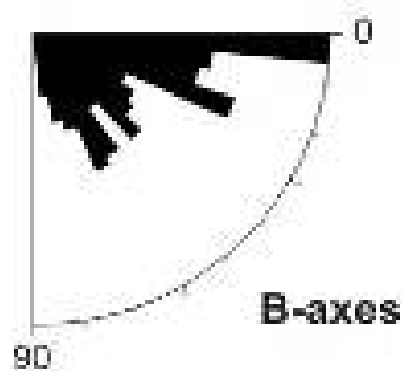




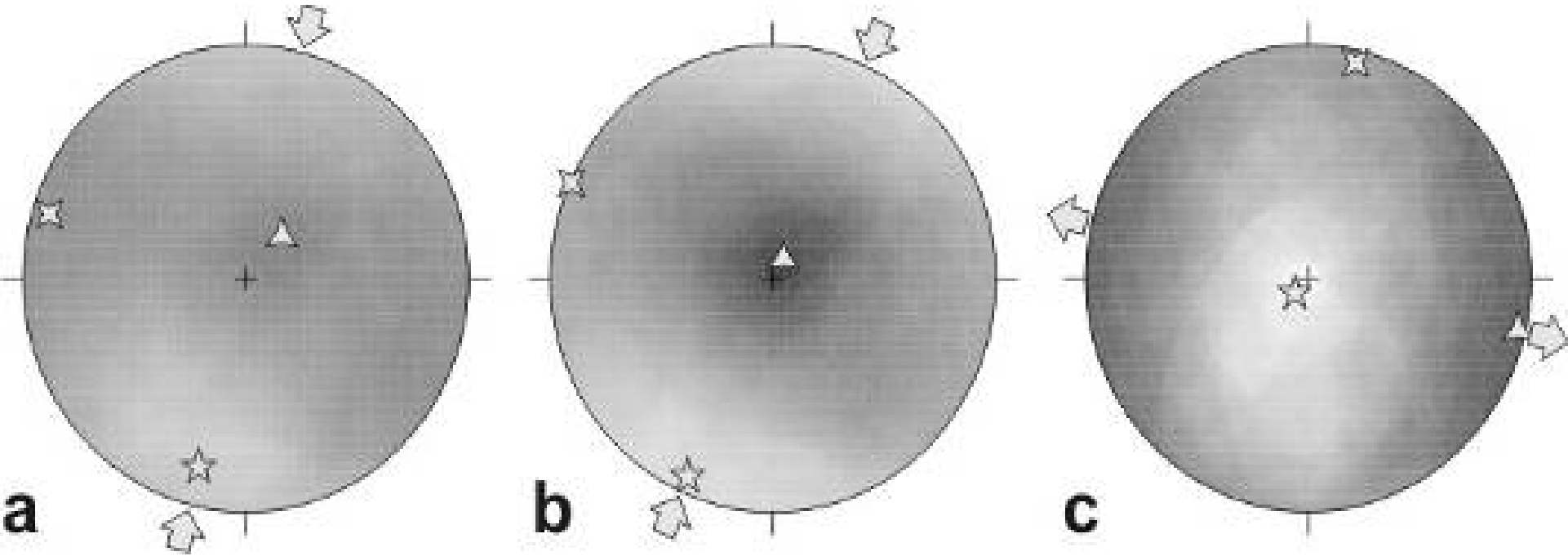
a

b

c



ROSE DIAGRAMS SHOWING THE ANGULAR DISTRIBUTION OF THE TREND (UPPER ROW) AND PLUNGES (LOWER ROW) OF THE INDIVIDUAL B-AXES, P-AXES AND T-AXES

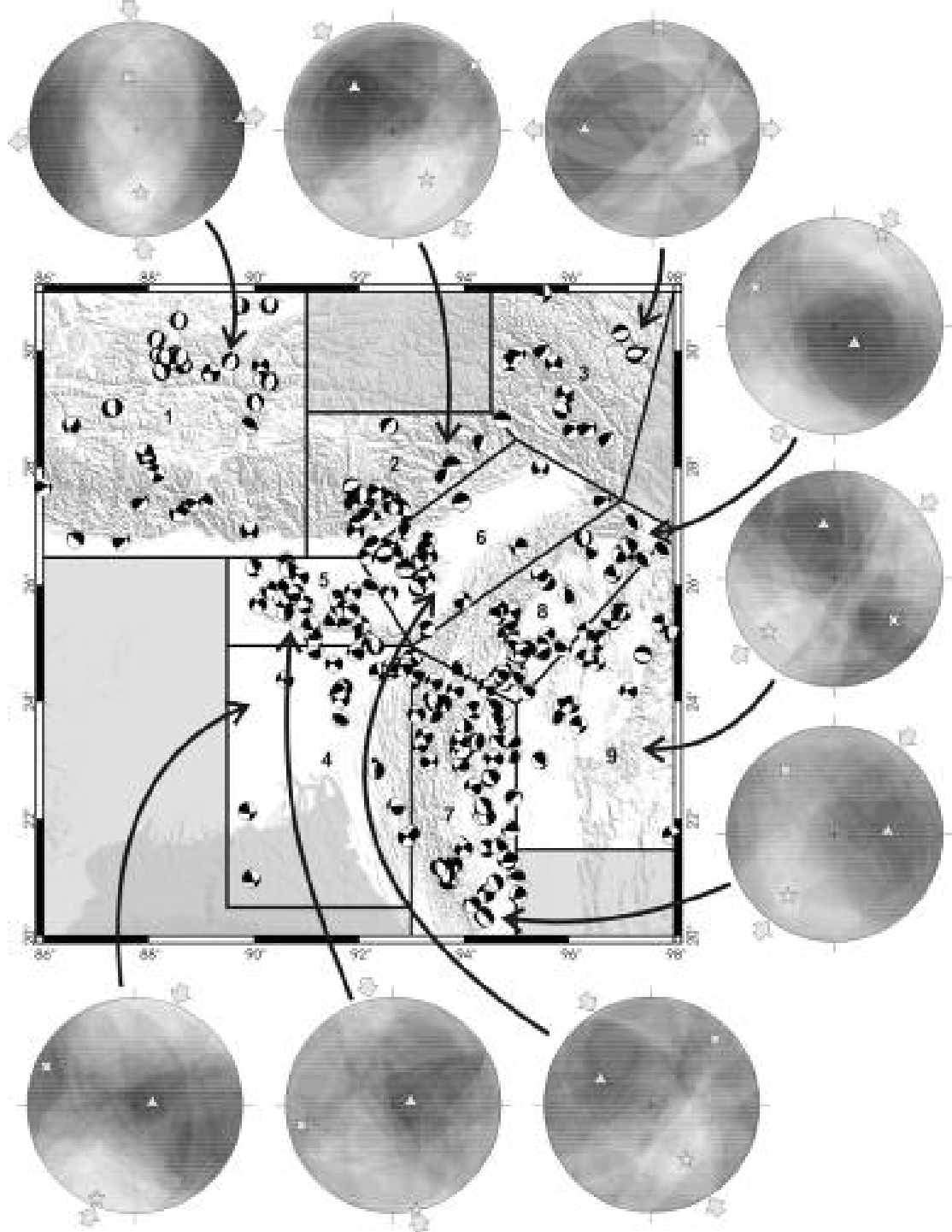


LARGE ARROWS SHOW INFERRED TREND OF COMPRESSION (CONVERGENT PAIRS OF ARROWS) AND EXTENSION (DIVERGENT ONES).

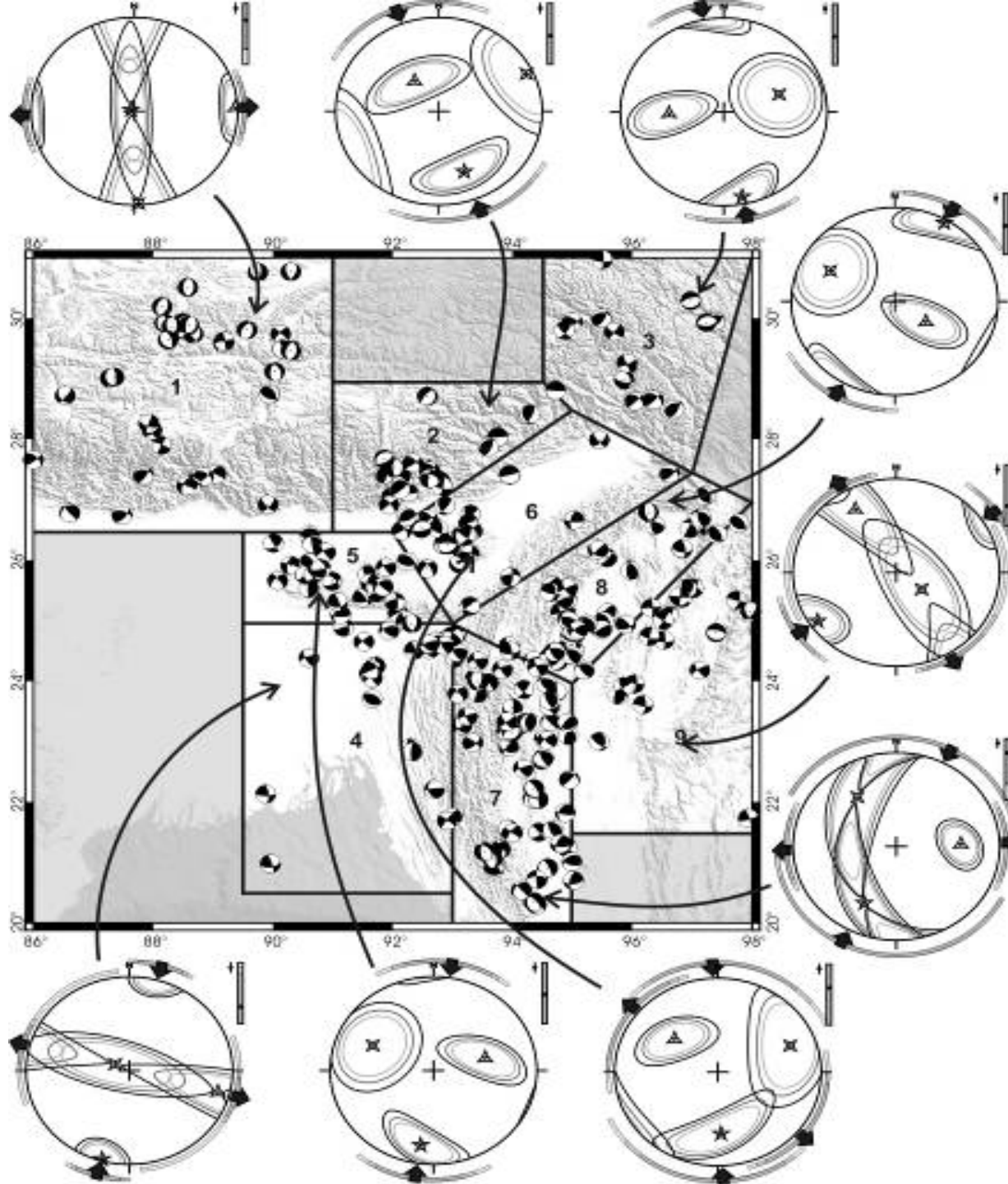
(A) WHOLE SET OF DATA, WITH 285 EARTHQUAKE FOCAL MECHANISMS.

(B) COMPRESSIVE SUBSET, INCLUDING 182 MECHANISMS.

(C) EXTENSIONAL SUBSET, INCLUDING 103 MECHANISMS



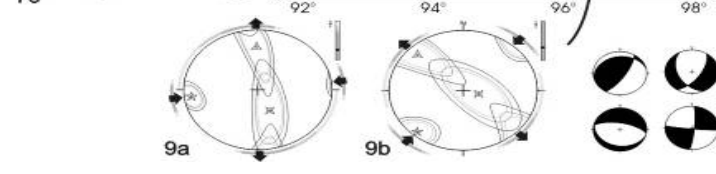
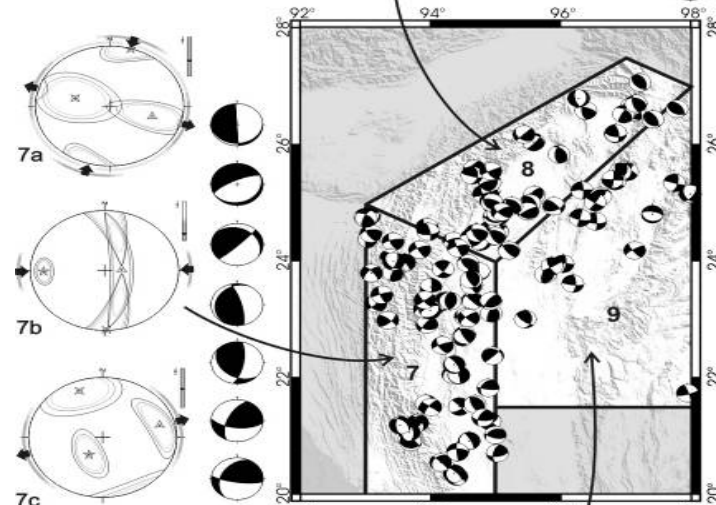
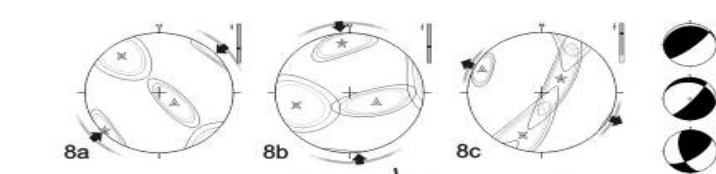
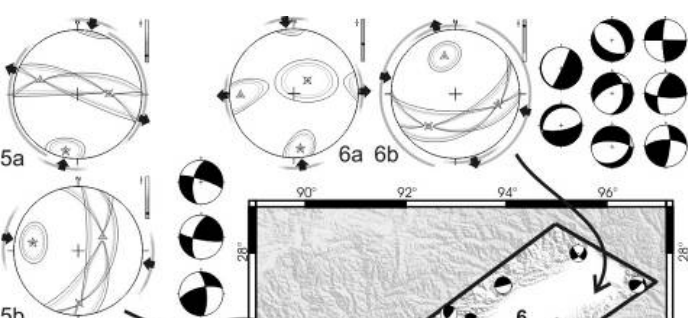
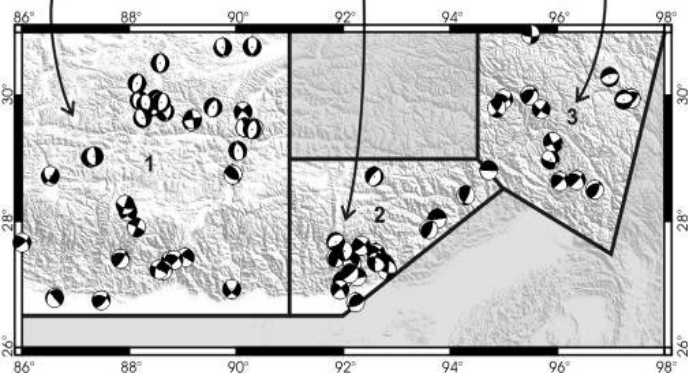
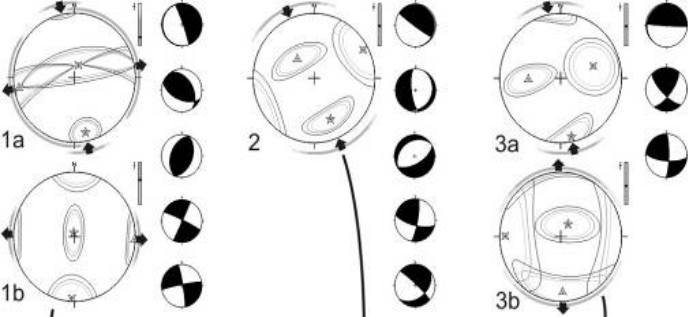
**STEREOPLOTS
SHOWING
THE
DENSITY
DISTRIBUTION OF
PRESSURE AND
TENSION
ON THE SPHERE OR
ALL
SPATIALLY
DEFINED SUBSETS**

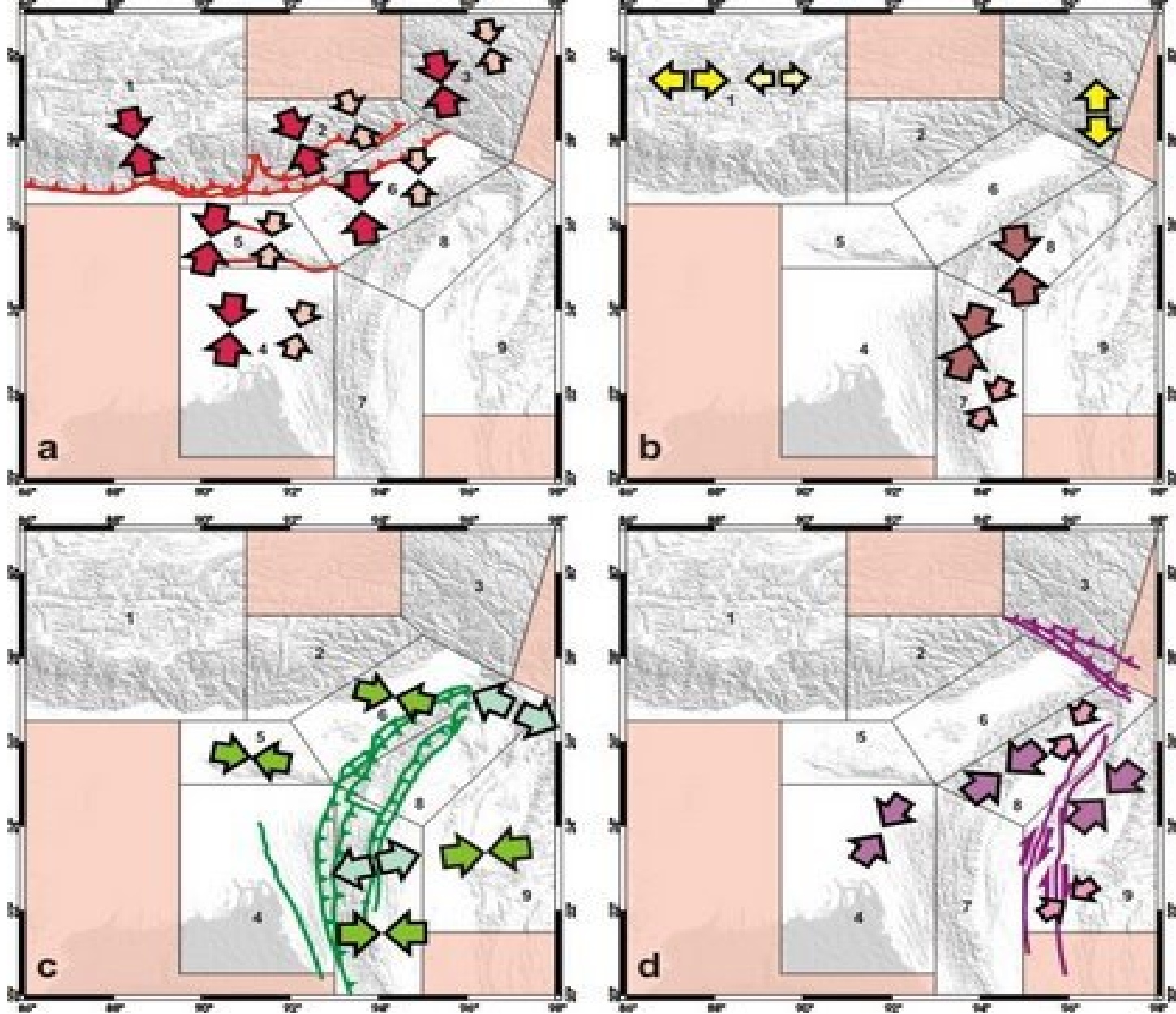


**RESULTS
OF
INVERSIONS
TO
OBTAIN
AVERAGE
STRESS
TENSOR**

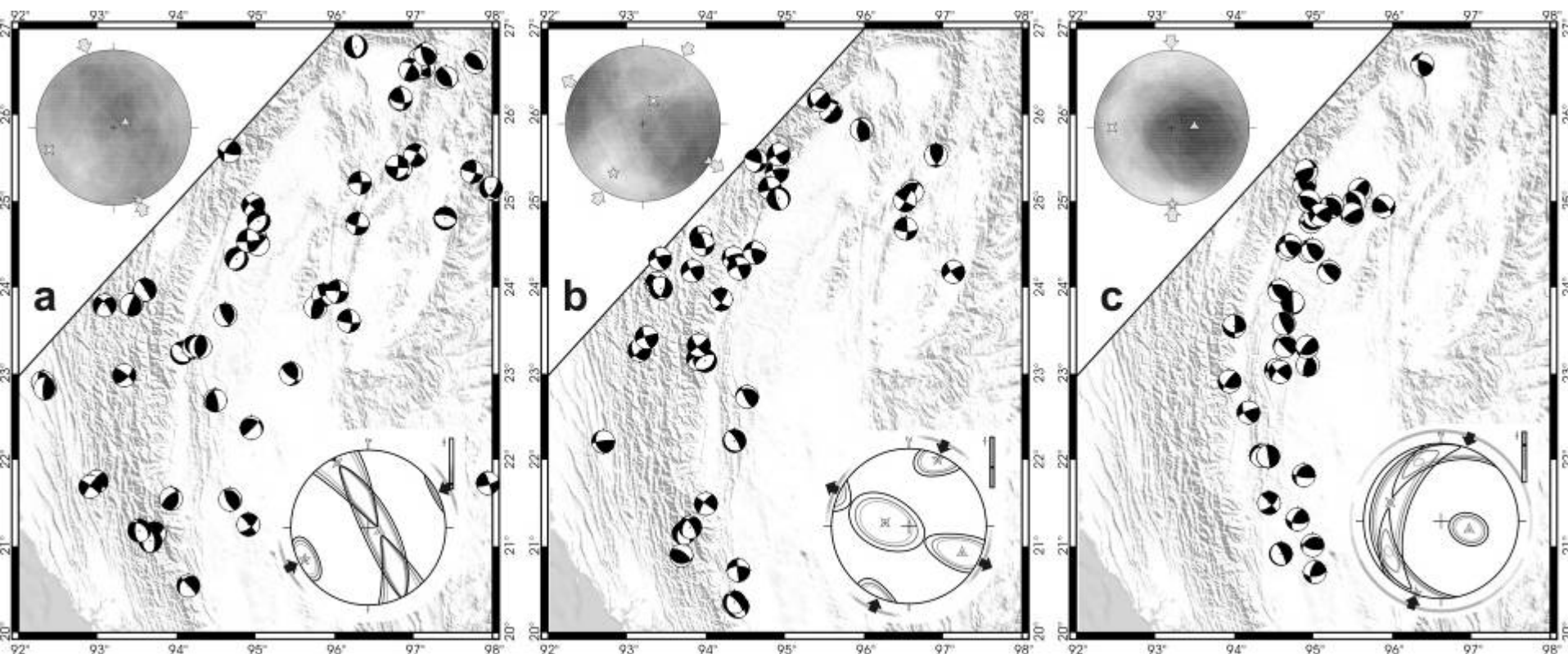
RESULTS OF STRESS INVERSIONS COUPLED WITH STRESS SEPERATION OF FOCAL MECHANISMS FOR

1. BHUTAN HIMALAYA
2. ARUNACHAL HIMALAYA
3. MISHMI THRUST
4. TRIPURA BELT
5. SHILLONG PLATEAU
6. ASSAM VALLEY
7. SOUTHERN INDO-BURMA REGION
8. EASTERN INDO-BURMA REGION AND
9. SAGAING FAULT REGION

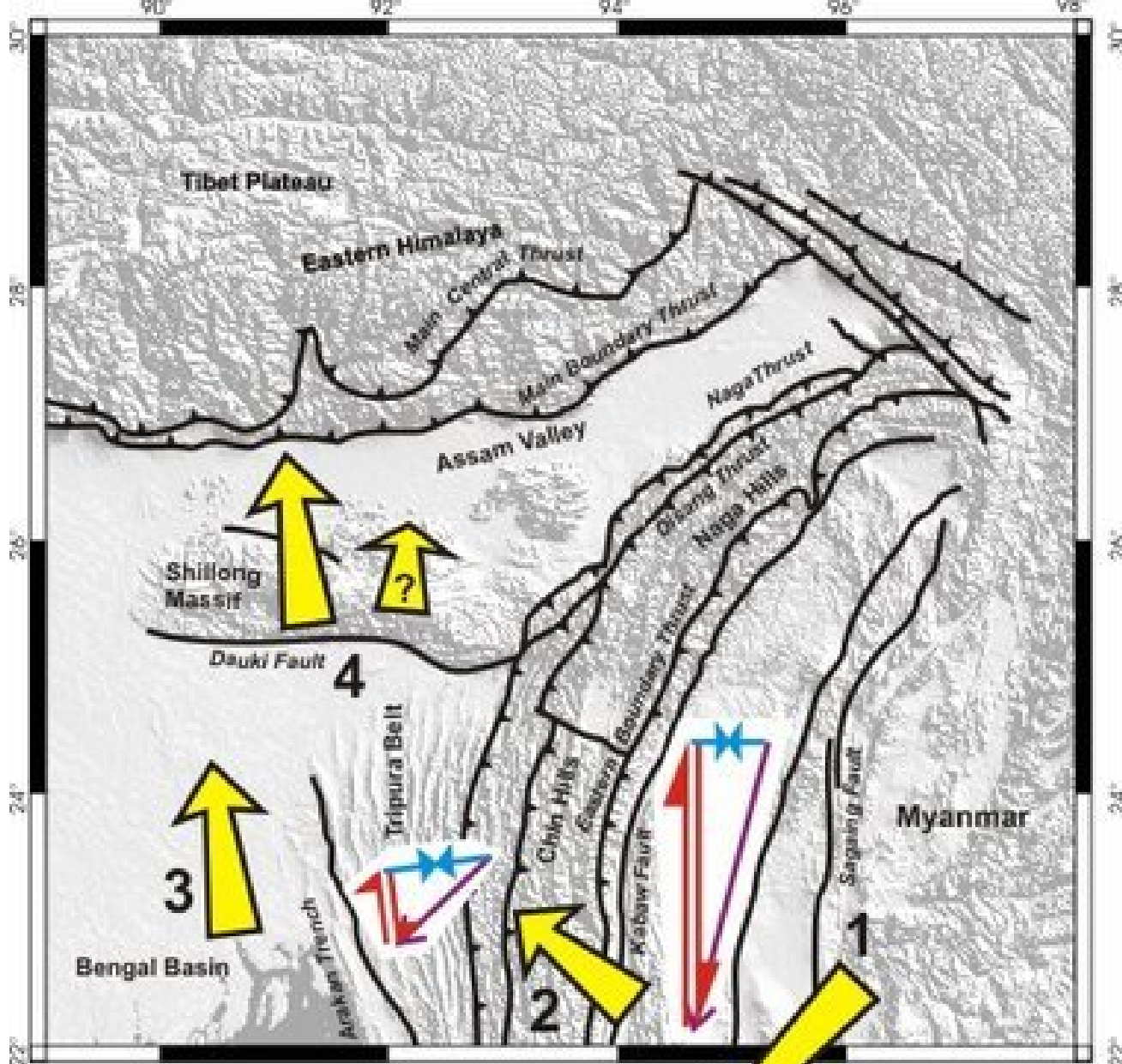




SYNTHESIS OF MAIN STRESS REGIMES. PAIRS OF CONVERGENT ARROWS FOR COMPRESSION, DIVERGENT ARROWS FOR EXTENSION.



Map of focal mechanism solutions in the Indo - Burma Ranges. The three maps correspond to different Earthquake depth ranges in the same area: (a) 0-45 km (56) ; (b) 45-90 km (39) and (c) 90-160 km (38).



MAJOR KINEMATIC FEATURES OF NORTHEAST INDIA.(a) PRESENT DAY AVERAGE VELOCITIES RELATIVE TO LASHA, TIBET: 1 WESTERN EDGE OF SUNDA PLATE 2 CENTRAL MYANMAR BASINS 3 BENGAL BASIN 4 SHILLONG-MIKIR MASSIF. (b) TO THE EAST, VELOCITY OF 1 wrt. 2 (c) MAINLY ACCOMODATED ALONG SAGAING FAULT, (d) TO THE WEST, VELOCITY OF 2 wrt. 3 ACCOMODATED ACROSS VARIOUS FAULTS OF BURMESE ARC AND TRIPURA BELT



CONCLUSION

- **North- South compression, in a direction consistent with India-Eurasia convergence, prevails in the whole area from the Eastern Himalayas to the Bengal Basin through Shillong-Mikir –Assam Valley block and the Indian Craton.**
- **The Indo-Burma ranges reveals complex stress pattern.**
- **Indo-Burma ranges are under compression as a result of oblique convergence between the Sunda and Indian plates. The maximum compressive stress rotates from NE-SW across the inner and northern arc to E-W near Bengal Basin.**

ACKNOWLEDGEMENT

I SINCERELY CONVEY MY HEARTFELT THANKS TO

PROF. CATHERINE DORBATH FOR INVITING ME to EOST

ALSO

THANKS TO THE ORGANIZER OF THE SEMINAR

THANK YOU VERY MUCH